#### Mapping ice deposits on Mars through subsurface radar sounding

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#### MARSIS and SHARAD

- MARSIS and SHARAD are syntheticaperture, orbital sounding radars, carried respectively by ESA's Mars Express and NASA's Mars Reconnaissance Orbiter. They work by transmitting a low-frequency radar pulse that is capable of penetrating below the surface, and is reflected by any dielectric discontinuity present in the subsurface.
- MARSIS is capable of transmitting at four different bands between 1.3 MHz and 5.5 MHz, with a 1 MHz bandwidth. SHARAD operates at a central frequency of 20 MHz transmitting a 10 MHz bandwidth.
- Whereas MARSIS is optimized for deep penetration, having detected echoes down to a depth of 3.7 km over the South Polar Layered Deposits, SHARAD is capable of a tenfold-finer vertical resolution, namely 15 m or less, depending on the dielectric constant of the material being sounded.







### Getting to know the data

MARSIS and SHARAD data are affected by a number of artifacts:

- Clutter: lateral reflections reaching the radar after nadir echoes, can be taken for subsurface echoes
- Ionosphere dispersion: the echoes become blurred as different frequencies propagate at different speeds
- Multi-path propagation: changes in the refraction index of the ionosphere bend the ray in unexpected ways
- Variation of propagation velocity: within media with a dielectric permittivity greater than 1, it changes the apparent shape of subsurface features





#### Clutter or subsurface detection?

SHARAD 194801 observation





#### Data vs. simulations



INAF

#### **Ionosphere** dispersion



### Multi-path propagation







#### Now there is, now there isn't...

#### 3821





### SHARAD coverage



### MARSIS coverage



#### MARSIS coverage (cont'd)

- After four years of operations, MARSIS has achieved a 42% coverage of Mars.
- Considering only the night-side orbits, coverage decreases to 30%.
- It was found that 15% (or more) of the surface has characteristics unsuitable to obtain good radar performances.
- Moreover the night coverage of the Northern Hemisphere is very sparse.





#### MARSIS and SHARAD Subsurface Detections

#### North Polar Deposits



#### **Planum Boreum**



#### Arcadia Planitia







#### **Deuteronilus Mensae**







#### **Medusae Fossae Formation**







#### East Hellas



#### **Dorsa Argentea Formation**







#### **Planum Australe**



#### Data over the SPLD



• There are areas in the SPLD where subsurface echoes are brighter than surface echoes, in spite of the attenuation resulting from propagation within a dielectric medium.





## Model of surface and subsurface reflections

- We assume that the detected surface and subsurface echoes are specular reflections from plane parallel layers.
- We also assume that dielectric properties are uniform within the layers.
- The peak power of the specular return from the surface is then given by (Porcello et al., 1974) [Eq 1]
- The reflection coefficient at the surface/atmosphere boundary is [Eq 2]

$$P_{s} = P_{t} \left(\frac{G\lambda}{8\pi H}\right)^{2} R_{s}^{2}$$

$$R_{s} = \frac{\sqrt{\epsilon_{s}} - 1}{\sqrt{\epsilon_{s}} + 1}$$





# Model of surface and subsurface reflections (2)

- The power reflected from a subsurface specular dielectric interface is given by [Eq 3]
- The subsurface interface reflection coefficient is given by [Eq 4]
- The time delay between echoes is related to the thickness and the the relative dielectric constant of the surface layer [Eq 5]

$$P_{ss} = P_t \left(\frac{G\lambda}{8\pi(H+X)}\right)^2 (1-R_s^2)^2 R_{ss}^2 \cdot 10^{-2.74\tau f \tan \delta_s}$$

$$R_{ss} = \left| \frac{\sqrt{\epsilon_{ss}} - \sqrt{\epsilon_s}}{\sqrt{\epsilon_{ss}} + \sqrt{\epsilon_s}} \right|$$

$$\tau = \frac{2X\sqrt{\epsilon_s}}{c}$$





### Radar wave propagation in the subsurface

- An electromagnetic wave reflected from the bottom of the SPLD is attenuated and scattered in several ways:
  - Surface scattering from the random rough SPLD surface
  - Attenuation within the dielectric SPLD material
  - Weak reflections within the SPLD due to the variation of dust concentration in ice with depth
  - Volume scattering caused by random variations of dielectric properties of the SPLD
  - Surface scattering from the random rough SPLD bottom.





#### Values of dielectric permittivity for natural materials

• CO<sub>2</sub> ice  $\epsilon \approx 2$ • H<sub>2</sub>O ice  $\epsilon \approx 3$ • H<sub>2</sub>O ice mixed with dust **E** < 3 Dry regolith **E** > **3**  $\varepsilon \approx 4-10$  Dry rock Water-bearing rocks or regolith ε > 10 Liquid water  $\varepsilon \approx 80$ 





Implications for the values of the dielectric permittivity at the base of the SPLD

- Strong echoes imply a large difference between the dielectric permittivity of the SPLD and that of the material beneath the SPLD.
- If the SPLD are mostly made of water ice, then  $\varepsilon \approx 3$ .
- In this case, to produce a reflection from the bottom of the SPLD that is as strong as the surface reflection,  $\varepsilon \approx 10$  for the material beneath the SPLD.





## Effects of attenuation and scattering within the SPLD

- Because of the weakening of the radar signal as it propagates through the SPLD, the real dielectric contrast between the SPLD and the underlying material is probably higher, requiring ε > 10 beneath the SPLD.
- Could this imply the presence of liquid water, or is there another <u>explanation?</u>







### Summary

- Subsurface layers have been seen by the MARSIS and SHARAD radars only in limited areas of Mars.
- Determination of the composition of subsurface layers is based on the estimate of their dielectric properties.
- A rigorous determination of such properties requires the inversion of the radar signal.
- Many factors can affect the strength of subsurface echoes, and excluding the presence of subsurface liquid water on the basis of existing analyses seems premature.



