# Io: A (geo-)physicist's playground

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#### **Presentation Structure**



- Aim: Discussion of phenomena at lo in the context of evolution.
- Basic surface characteristics and composition
- Atmosphere and escape
- Laplace resonance
- Current evolution of the system
- Sampling the internal composition of Io
  - Io Volcano Observer concept

#### **Basic Surface Characteristics**

#### • The surface itself can tell us very little about lo's history.



G7: April 1997

C10: September 1997

C21: July 1999

#### Impact Rates in the Jupiter System



This calculation gives around 1.25, 100 km impacts every 10<sup>7</sup> years

Zahnle et al., Icarus, 2003

#### **Basic Surface Characteristics**

#### The surface itself can tell us very little about lo's history.



G7: April 1997

C10: September 1997

C21: July 1999

- Re-surfacing rates are estimated to be 1-10 mm year<sup>-1</sup> (cf. Soderblom et al., 1979, Spencer and Schneider, 1996)
- 100 km diameter crater buried in <2 10<sup>7</sup> years.
- NB: 1.5-4.6% of material ejected from Io by an impact hits Ganymede.
- Sulphur/oxygen from IPT may be an even larger source (Carlson et al., 2004).

# Maui-Amirani Flowfield



# **Colourful Surface**



Colours indicating sulphur and its various allotropes. Some argument in the literature about purity (e.g. Kargel et al., 1999; see also McEwen et al., 2004).

 $S_2$  detected as a gaseous species using HST.



#### **Surface Composition**

• Dominated by sulphur and compounds (SO<sub>2</sub>).



Detection of gaseous SO<sub>2</sub> above the surface of Io by the IRIS experiment on Voyager 2

Pearl et al., 1979

#### **Surface Composition**

- Dominated by sulphur and compounds (SO<sub>2</sub>).
- Originally silicates expected to be present (to provide structural strength; Clow and Carr, 1980)

# High constructs indicating structural strength - silicates

50 km

km high

11



#### **Surface Composition**

- Dominated by sulphur and compounds (SO<sub>2</sub>).
- Originally silicates expected to be present (to provide structural strength; Clow and Carr, 1980)
- Discovery of high temperature volcanism (GLL NIMS and groundbased) consistent with silicate volcanism (E.g. Davies, 2007).
- Other species inferred to be present from atmospheric and torus emissions.





127 (22 Feb 2000)



## Atmospheric and Torus Emissions



Atmospheric CI I emission (Feaga et al., 2004) after CI II discovered by Küppers and Schneider (GRL, 2000).

NaCl detected – other compounds likely (see Moullet et al., 2010).





# [OI] in the torus at 6300 Å

Potassium in the neutral clouds at 7699 Å

Thomas , 356 B.C.

## lo Plasma Torus



#### Charge-exchanged neutral





#### **Hydrogen-Bearing Species**



- Limited evidence of hydrogen. Some evidence of H<sub>2</sub>S from HST.
- But NO clear evidence of water (see Carlson et al., 2007)

E/Q < 100 indicating protons; Frank and Paterson, 1999

#### **Current Mass Loss Rates**

- 1 tonne/s lost at present.
- Equivalent to roughly 300 m Gyr<sup>-1</sup> of loss in radius of the moon.
- Nowhere near enough to remove a water layer similar to that of Europa.

Absence of water primordial unless past conditions markedly different from today.

#### **Carbon-Bearing Species**

- Some discussion of CO<sub>2</sub> in 1990s (e.g. Sandford et al., Icarus, 1991) but probably faint (previously unknown) SO<sub>2</sub> lines.
- CO and CO<sub>2</sub> should be most abundant in volcanic gases (Schaefer and Fegley, ApJ, 2005)
- Upper limit on C/S of 10<sup>-3</sup>
- Upper limits on CS<sub>2</sub> set by HST (but not actually expected on basis of chemical equil. calculations.)



Mole fraction of carbon species as a function of the C/S atomic ratio for Pele vent conditions. CONDOR model of S+F, 2005.

#### **Composition Summary**

- We have a crude inventory of species from remote sensing.
- However, this is not really quantitative.
- Several key items missing/undetermined.

#### Internal Structure



- 3.528 g/cm<sup>3</sup>.
- Substantial Fe-rich core expected but likely to be rich in sulphur (Schubert et al., 2005).
- Size range from 650 -950 km (10-20% of Io's mass).
- Lithosphere can be thick (and probably needs to be to support topography;
  Turtle et al.). Simple models suggest <40 km (Moore et al., 2007).</li>
- Heat conduction through lithosphere not a major transport mechanism – magmatic transport required.
- Not completely obvious that lithospheric composition same as underlying mantle.
- Crust "re-cycled" in  $< 4 \ 10^7$  yr.

# Laplace Resonance $\lambda_I - 3\lambda_E + 2\lambda_G \approx \pi$

 $\lambda_x$  = mean longitudes of the satellites Io, Europa and Ganymede



Europa forces lo's orbit to be eccentric leading to motion of the tidal bulge which is raised by Jupiter.

 $C_{22}$  > 50 times larger than Callisto.

Friction generates enormous quantities of heat leading to observed volcanism.

Global heat input: 2.5 W m<sup>-2</sup> (Veeder et al. 1994) 0.4-1.2 W m<sup>-2</sup> (Marchis et al. 2005) For 4.5 Gyr that = 0.03% of Jupiter's rotational energy!

#### Synchronous Rotation

- Peale (1977) estimated despin timescale as
  - $-t_{despin} \simeq 5000 \text{ x (Q/100) yr.}$
- Q = 10 500 (Goldreich and Soter, Icarus, 1966)

 Implies extremely rapid transition to synchronous rotation unless Io underwent large migration.

#### Is the Laplace resonance stable?

• Lainey et al., Nature, 2009 claim that the system is evolving OUT of the Laplace resonance (at the 3 sigma level).



### **Evolution of Tidal Heating**

- Fischer and Spohn (1990) studied evolution through the Laplace resonance
  - There are multiple equilibria!
  - E.g. Frictional heating slows as solid matrix loses coherence.

Also, if dissipation reduces (because lo leaves the resonance inwards), the satellite will accelerate outwards again – back into resonance.

Moore et al. (2007) after Fischer and Spohn (1990)



#### The Fischer-Spohn Assumptions

- Three layer model with an Fe/FeS fluid core, a silicate mantle, and a thin (30 km) lithosphere.
- Few direct constraints on lo's rheology.

 But the time needed to reach the present state suggests Laplace resonance age of >0.5 Gyr.

## **Volcanic Plumes**



Pillan

Geissler and McMillan (2010)

#### Volcanic Plume, Ra.



Separated by 1 hour. Out of and in eclipse. Geissler and McMillan (2010)

#### **Prometheus-Type Models**



Particles included in other runs.

#### **Comprehensive 3-D DSMC Model**



## Io Volcano Observer

A Discovery mission to lo

- Proposed in last Discovery round and ranked category 2.
- Details in Adams, E. et al., IEEE, 2012.



• The aim would be ultimately to fly through and sample a plume

#### **Payload Summary**

Instrument	Mass (ex. Shield) [kg]
Narrow angle camera	8.3
Wide angle camera	4.8
Ion and neutral mass spectrometer	4.3
Plasma ion analyser	0.9
Fluxgate magnetometer	2.9
Thermal mapper	8.8

# IVO: Fly-bys

- **Goal:** minimize the total ionizing dose while achieving science objectives
  - I1 & I2 are nightside passes for plume & hot spot searches
  - I3 & I5 are dayside passes optimized for measurement of the induced magnetic signature from mantle melt
  - I4 is a 178-km alt flyby optimized for surface composition measurements by INMS & hot spot observations
  - I6 is a 200-km alt flyby over the Loki Patera





#### RDM = radiation design margin

Adams et al. (2012)

#### Conclusions

- Io's activity has erased any trace of surface evolution prior to ~2 10<sup>7</sup> years ago.
- That activity has "polluted" the surface of the other satellites.
- Absence of water (in any form) places constraints on the formation mechanism.
- Laplace resonance is probably old (>0.5 Gyr).
- Further progress can be made IF Io's composition can be investigated.
- Internal composition can be sampled by a "fly-through" mission which need not be vastly expensive.



Plot concept "stolen" from Axel Korth

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# Deviation from co-rotation in Jupiter's magnetosphere (SIII)





The Interior of lo

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#### **Satellite Formation Models**

- Accretion disc
  - Circum-Jupiter disc as solar nebula gas flows in
- Spinout disc
  - Stranded material as Jupiter cools and contracts
- Blowout or impact-generated
  - Perturbation of Jupiter's axis??
- Gas-free co-accretion disc
  - Small body collision with Jupiter's Hill sphere
    - Composition gradient??

(see McKinnon, 2007)

#### Formation?

- Following core accretion of Jupiter, the accretion of gas and dust continued but did not terminate "cleanly".
- The Galilean satellites may have formed from a gas-starved, accretion disc.
- The satellite disc was probably fed until the solar nebula was no longer present to feed it.
- Given the ages of observed nebulae, Io probably formed within 10 Myr of solar system formation and possibly before solar hydrogen ignition.
- The temperature gradient (required to explain the depletion of water at lo but not at Europa) is strongly dependent upon many parameters which are poorly known.

(see McKinnon, 2007)

## Energy Balance (Steady-State)

- Heat transfer
  - Jupiter's luminosity
  - Potential energy of infalling matter
  - Viscous dissipation within disc
  - Illumination from the proto-Sun/solar nebula
  - Radiative cooling



Cf. free sublimation temperature of  $H_2O \approx 200$  K.

(see McKinnon, 2007 after Canup and Ward, 2002)

#### Temperature Distribution in a Time-Evolving Inflow



- Alibert et al. (2005)
- Implies migration inward over large distances.
- Time scale for propagation of snow line is extremely fast.
- May be at odds with the idea that post-formation lo should evolve outwards (acceleration by Jupiter).

#### Is the Laplace resonance stable?

- Lainey et al., Nature, 2009 claim that the system is evolving OUT of the Laplace resonance (at the 3 sigma level).
  - Io moves inwards, towards Jupiter, and loses more orbital energy by dissipation of solid-body tides raised by Jupiter and by the Laplace resonance interaction than it gains from the exchange of angular momentum with Jupiter's rotational energy through tidal dissipation in Jupiter.
  - This would be surprising since it would imply we are at a preferred time but ....