

Mars photometry with OMEGA observations

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Spectral variations in OMEGA data set

- Compositional variations
 - Albedo variations
- + Variations related to:
- viewing geometry \Leftarrow surface physical properties
 - aerosols scattering \Leftarrow aerosols density,
composition,
size distribution
 - surface temperature \Leftarrow emissivity vs reflectance

Spectral variations in OMEGA data set

Variability is explored through signal modeling in a number of situations

Expected outcome:

- Signal prediction for observations
- Understanding the crucial parameters
- Defining a valid range for these parameters
- Deriving physical information

Mars spectral modeling

Computes:

$I/F = f(L_s, \text{local time, location, viewing geometry, surface tilt})$

Based on previous simulations (Erard 2001, GRL), improved

Uses new reference data:

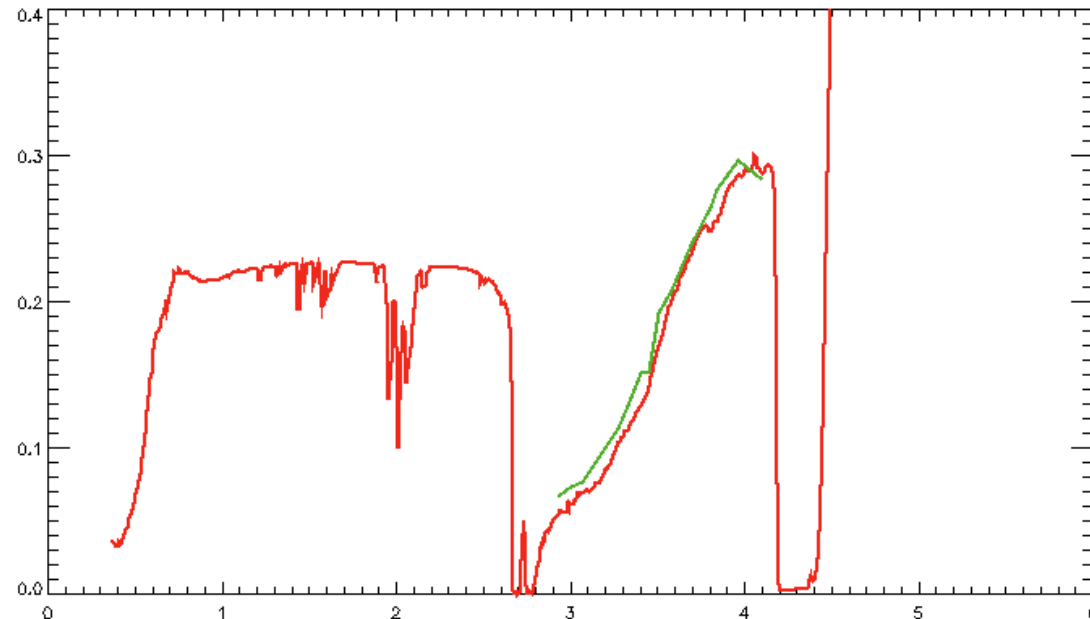
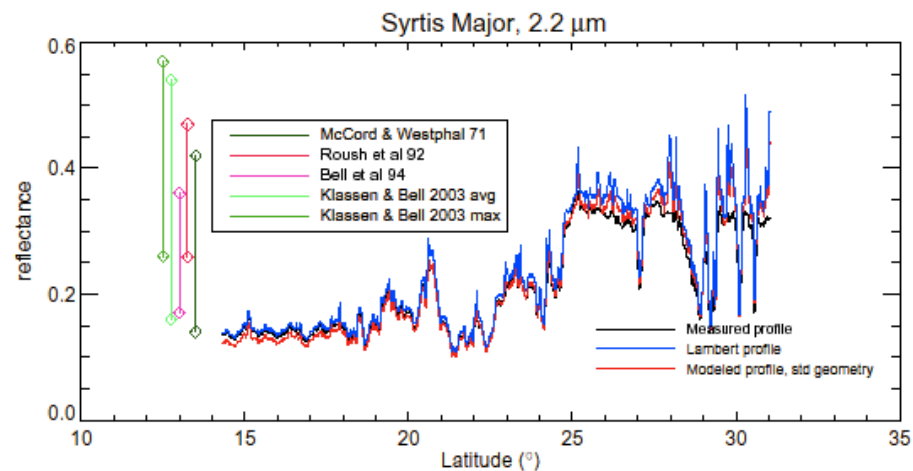
- Complete interface to European Mars Climate Database (v. 3.1, MGS scenario - Lewis et al 2001)
=> $T, P = f(L_s, \text{local time, Long, Lat, } z)$
- Interface to TES albedo map
=> surface albedo / spectral type in a 1° box
- Absolute solar spectrum (from Colina et al 1996 + Kurutz 1997)
- Sun distance and L_s (from Allison & McEwen 2000)

Model output

- 1) Spectra (here compared with Klassen & Bell 2003, IRTF)
- 2) Monochromatic profiles (modeling of OMEGA session)

Syrtis Major
Dark regions

Arabia Terra $\phi = 9.3^\circ$, $e = 25^\circ$, $i = 34^\circ$
 $L_s = 53^\circ$



Mars spectral modeling - arguments

Prime arguments:

Ls

local time

longitude / latitude

viewing geometry (e, i, φ)

From Ω observations
(Spice routines)

Optional:

surface orientation (i, e) included in the Ω data base
(computed from MOLA)

Surface pressure

Opacity

Surface reflectance

} Default values from GCM + TES,
can be changed

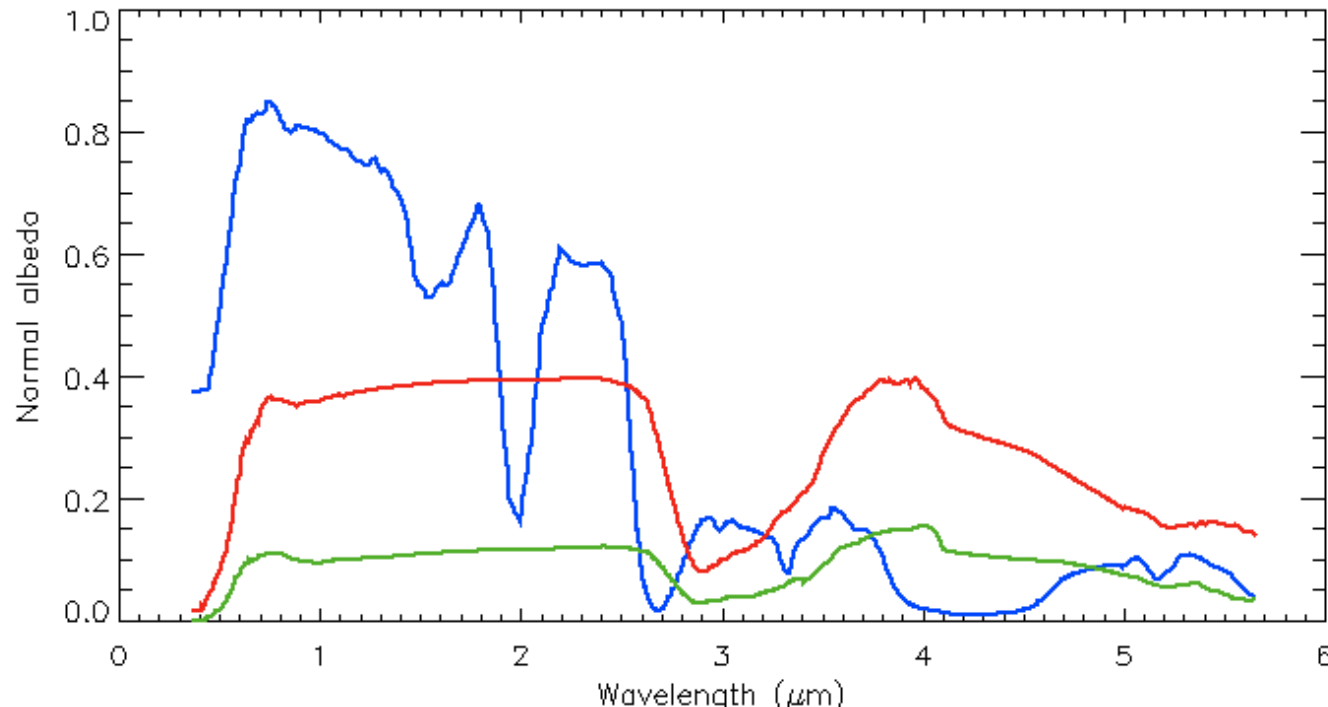
Mars spectral modeling - surface

Surface reflectance: linear combinations of 3 spectral types

0.3-5.7 μm composites (modified from Erard & Calvin 1997)

Bright & dark soils from thermal modeling of IRS at longer λ
(derived under Lambertian assumption)

Polar caps spectra poorly known before OMEGA



⇒ Check absolute levels
⇒ Check cap spectrum

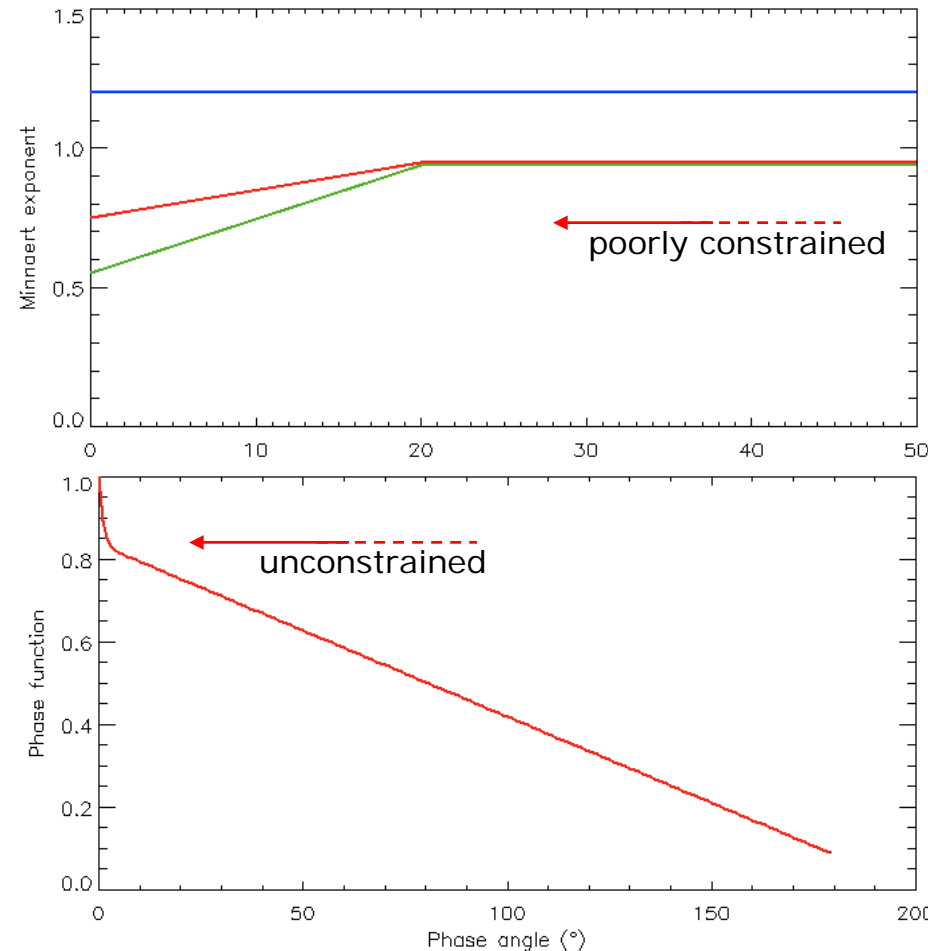
Mars spectral modeling - surface

Surface photometry

1) Minnaert + phase function

Minnaert exponent from a synthesis of observations:
Binder & Jones, 1972, Thorpe 1973,
Pleskot & Kieffer 1977,
de Grenier & Pinet 1995, Erard 2000

Phase function from ISM bright regions
(Kirkland et al 1997)
+ other solutions (Heynyey-Greenstein)
No spectral dependency
Pb: Minnaert breaks down at phase $> 60^\circ$



Mars spectral modeling - surface

Surface photometry

2) Hapke model

Parameters from visible measurements:

Arvidson et al 1989 (Viking lander red filter)

Johnson et al 1999 (Pathfinder bright soils)

$w = 0.7$ $\Theta = 5^\circ$
(but isotropic particles)

=> 5 parameters per surface type

No observational constraint on dark regions

Mars spectral modeling - atmosphere

Model atmosphere

Very simple modeling, but not required to be very accurate here.

Two options:

- 1) 6 mbar CO_2 (downsampled from MODTRAN band model)
+ simple log scaling with path length
- 2) OMEGA a posteriori simulations (Melchiorri et al 2004)
 - Exact geometrical path length computed for high angles
<= Neglects refraction
 - Rayleigh scattering added
 - Atmospheric emission added (single isothermal layer)

Mars spectral modeling - aerosols

Aerosols scattering / emission

Used mainly for continuum shape (spectral slope):

- Decoupled from gaseous absorption, single scattering
- Q_{back} , Q_{ext} and phase function from Mie theory:
 - ~ neutral dust (mineralogical) $n = 1.55$ $k = 0.01$
 - Modified-gamma size distribution (Drossart et al 1991)
- Scaled with opacity from GCM [$f(L_s, \text{Lat}, \text{Pressure})$]
- Backscattered and emitted light added to surface reflectance, filtered by upper atmospheric layers => reasonable approximation at low phase / i , e angles

Mars spectral modeling - assessment

Validity domain:

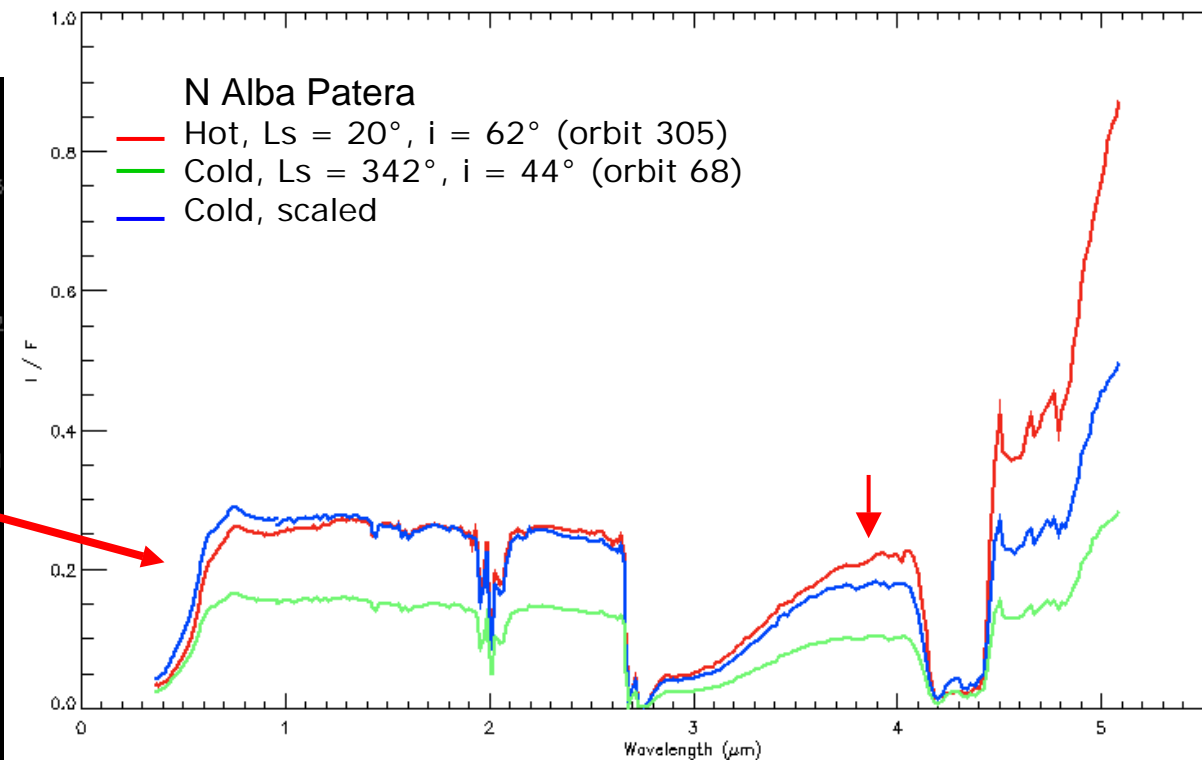
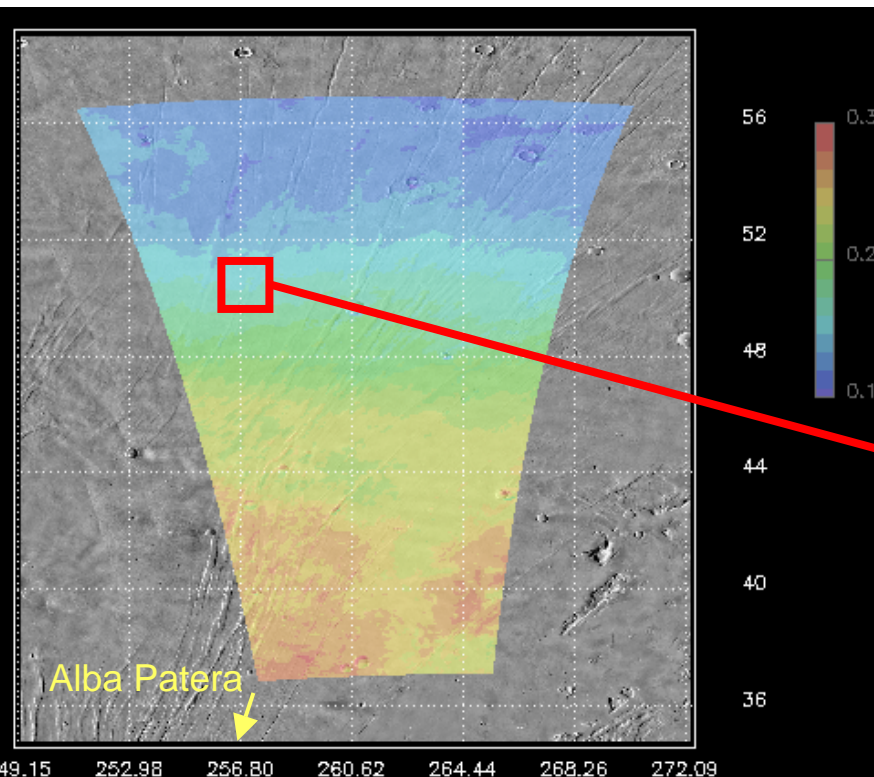
- 0.3-3 μm range
- Moderate phases ($<30^\circ$) / near-normal geometry ($<40^\circ$)

What needs to be better constrained:

- Surface reflectance at long wavelengths
- Surface phase function at large phase
- Aerosol phase function at large phase
- Aerosol scattering at large angles / large opacity

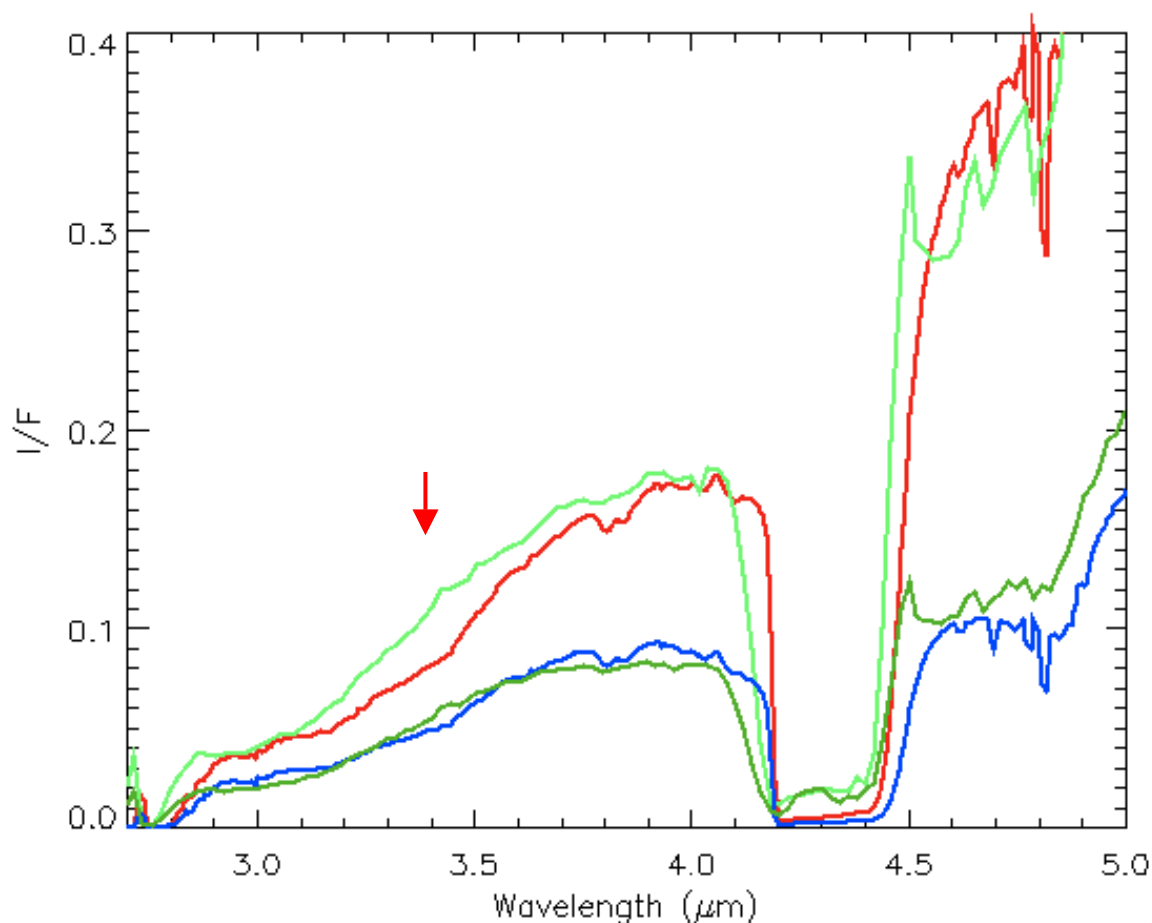
Reflectance at longer wavelengths

OMEGA: allows observation of reflectance at long wavelength
Cold terrains (winter, early morning): low thermal contribution



Reflectance at longer wavelengths

Observations vs simulations



N Alba Patera

- Hot, Ls = 20°, i = 62°
- Hot, simulated (Tsurf = 246K)
- Cold, Ls = 342°, i = 44°
- Cold, simulated (Tsurf = 219K)

~ correct fit of level
and shape

⇒ reflectance estimates
~OK at long wavelengths?
(slightly overestimated
3.2-3.6 μm range?)

Angular variations

Pb: Surface photometry & aerosols scattering always express together (at least in the 0.7-2.7 μm range)

Methods:

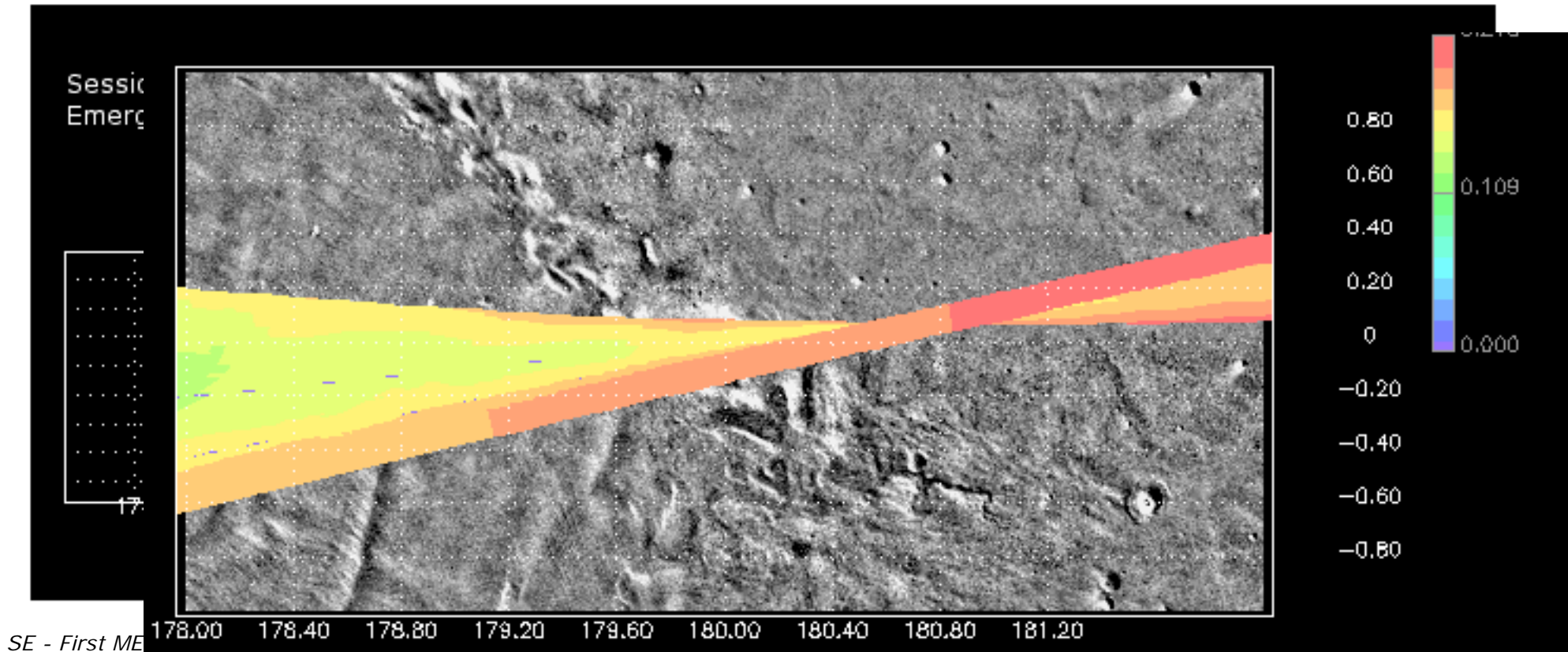
- EPF sequences
 - Pb: very limited coverage with OMEGA so far
- Observations of same areas under different opacities
- Observations of same areas under different geometries
 - Overlapping observations
 - (e.g. presentation by Baratoux yesterday with HRSC)
 - Several instruments
 - Pb: Amplifies remaining calibration issues

OMEGA EPFS

Bright uniform region (Lucus Planum, lat= 0°, long ~ 180°)

Scaled TES albedo (1 μm) = 0.31 - 0.34

Ls = 58.8° (approaching aphelion) => low expected opacity ~0.19 in visible
Morning (7:50) — 30 min acquisition => possible dissipating fogs
see poster by Pinet et al.



OMEGA EPFS

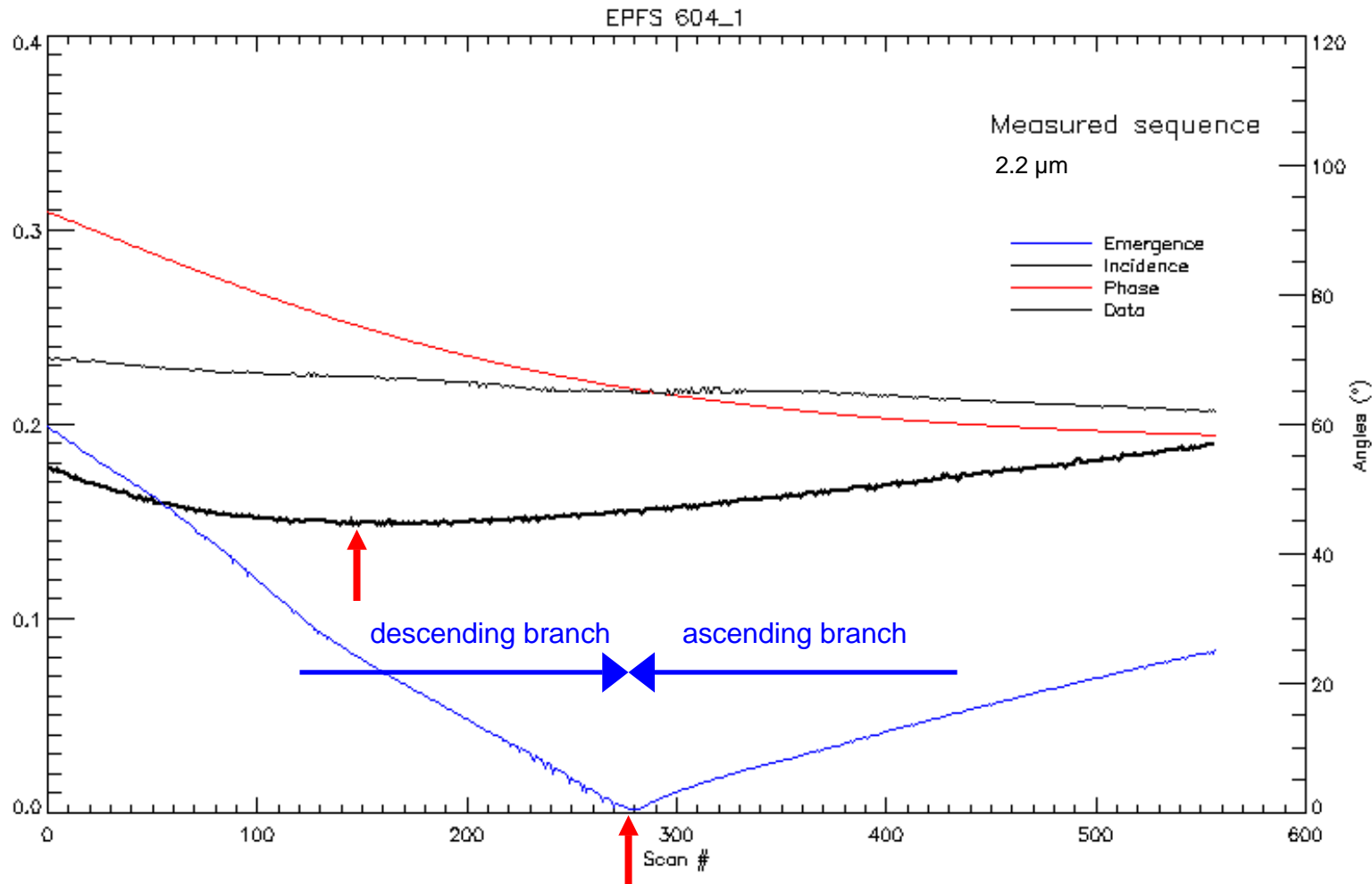
Properties:

Large variations in e (60° - 0° - 25°)

Slow decrease in incidence (70 - 62°)

Large phase angles (58° - 93°)

- Lower level does not correspond anything special

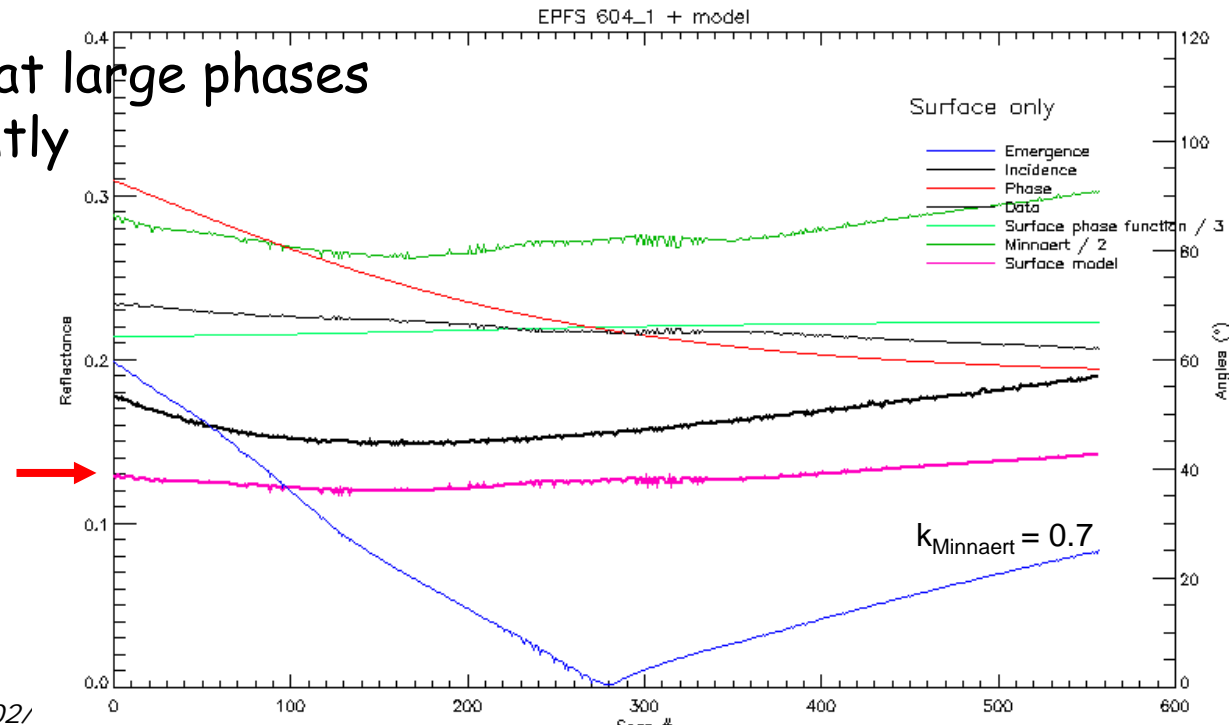


EPFS fit: surface model

Best fit: $R_{surf} = P_N * \mu_0^k / \mu^{(1-k)} * f(\varphi)$ (Minnaert + phase fct)
Minnaert $k=0.7$ (value at low phase) + less steep phase function

- 1) Angular variations ~ follow measurements
- 2) Measured signal = 1.3 x surface model

- Not Lambertian
- Brighter than expected at large phases
- Visible behaves differently



EPFS fit: aerosols scattering

Previous surface model + Mie scattering

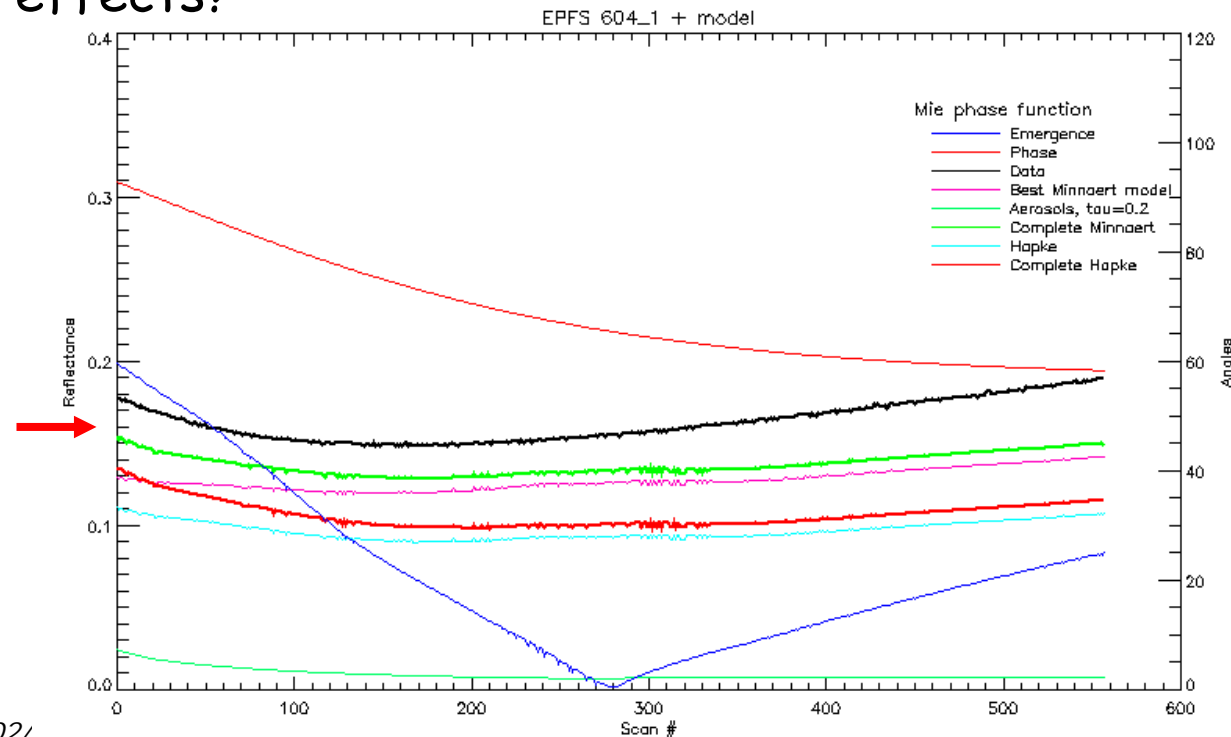
1) \pm Correct angular shape in the NIR

Mie phase function too steep - Overestimated at large phases?

Color effects? \Rightarrow Non-spherical particles

2) opacity ~ 0.2 does not brightens spectra

\Rightarrow Multiple scattering effects?

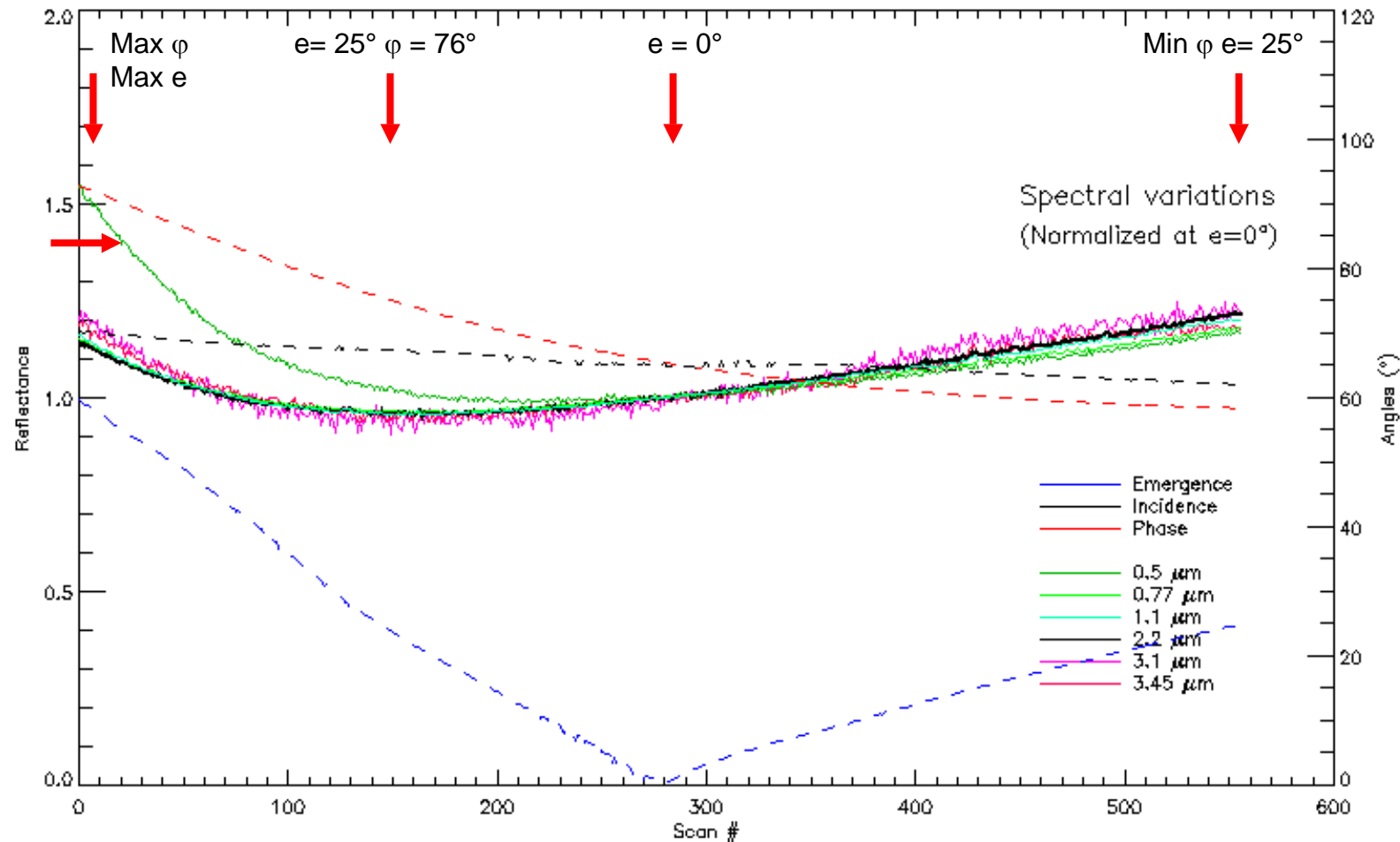


OMEGA EPFS - Spectral variations

Very homogeneous in the NIR (0.7-2.5 μm)

Larger NIR difference occurs at lower phases

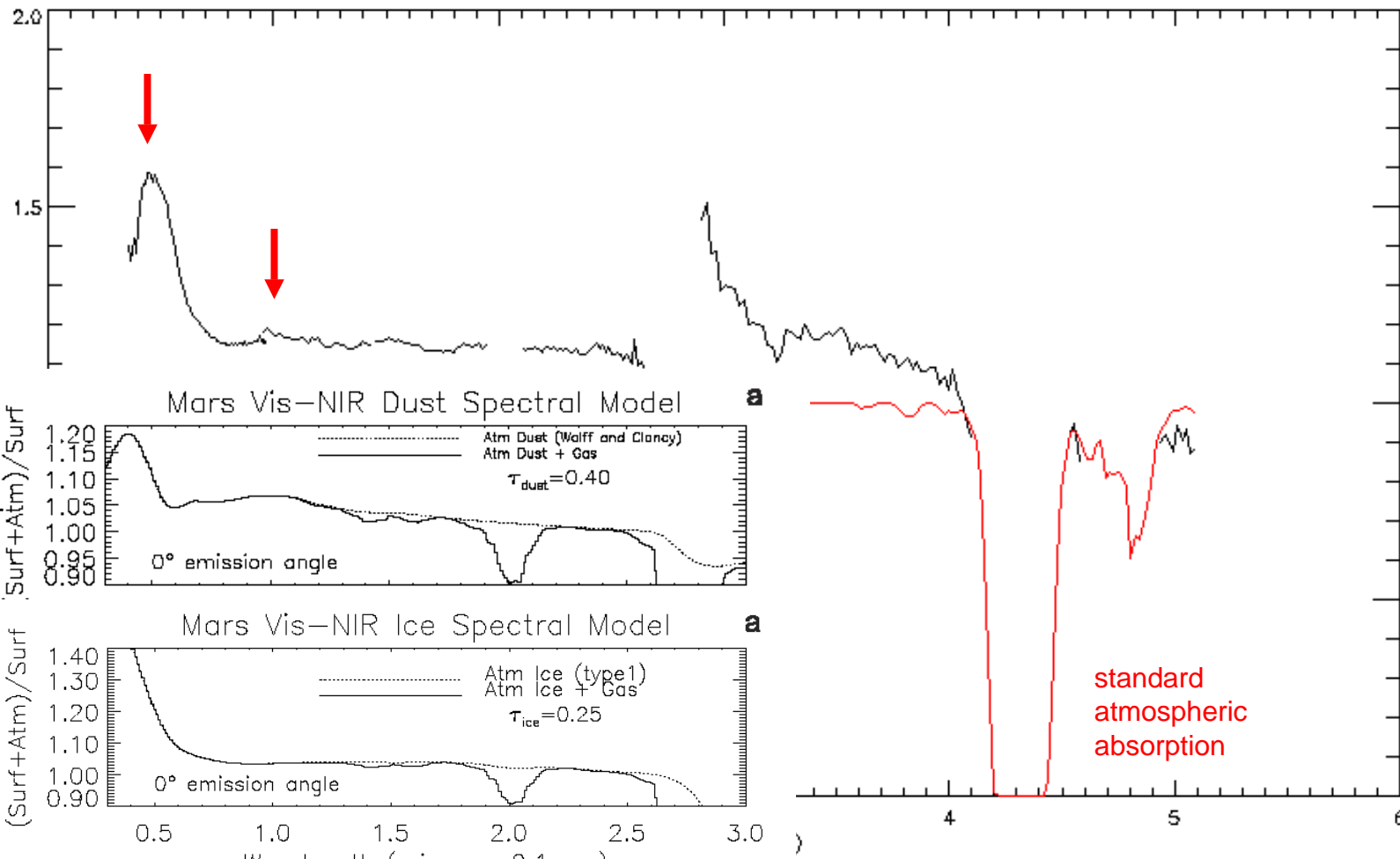
Stronger dependency on e or ϕ in the visible



Aerosols composition?

Ratios are similar to Wolfe & Clancy 2003 dust model (not ice)
Broad maximum at $\sim 1 \mu\text{m}$ associated with $1.5 \mu\text{m}$ mean particle size

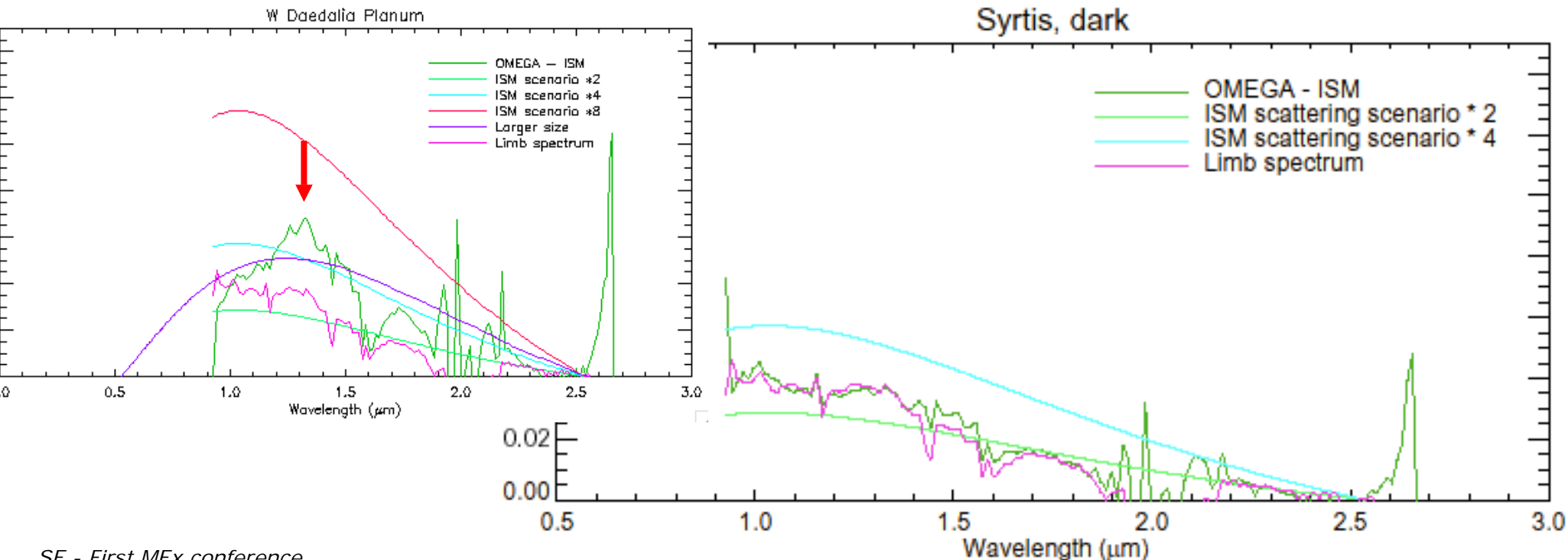
Ratio of
maximum e and $\varphi / e = 0^\circ$



Scattered spectrum

OMEGA-ISM in similar areas + Limb spectra

- Level fit with reasonable opacity ($\tau \sim 0.3-0.5$)
- Occasional maximum at $1.3 \mu\text{m}$ fit by slightly larger (30%) effective radius



Conclusions

What can be improved from OMEGA analysis:

- Surface phase function at large phases
- Surface phase function in the visible
- Mie phase function for scattering (level, spectral dependency)
- Variations in aerosol size?

Prospects

- More accurate atmospheric/scattering modeling required, in particular at large phase angles
(e.g. Blecka 2002, Blecka & Erard 2004)
- Systematic study of cold areas to assess surface reflectance in the 3-5 μm range
- More systematic study of aerosol scattering at large phase