

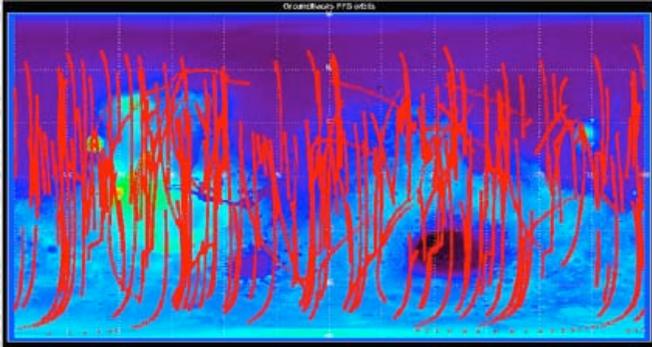
METASTABLE METHANE CLATHRATE PARTICLES AS A SOURCE OF METHANE TO THE MARTIAN ATMOSPHERE

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Why a seasonal cycle of methane?

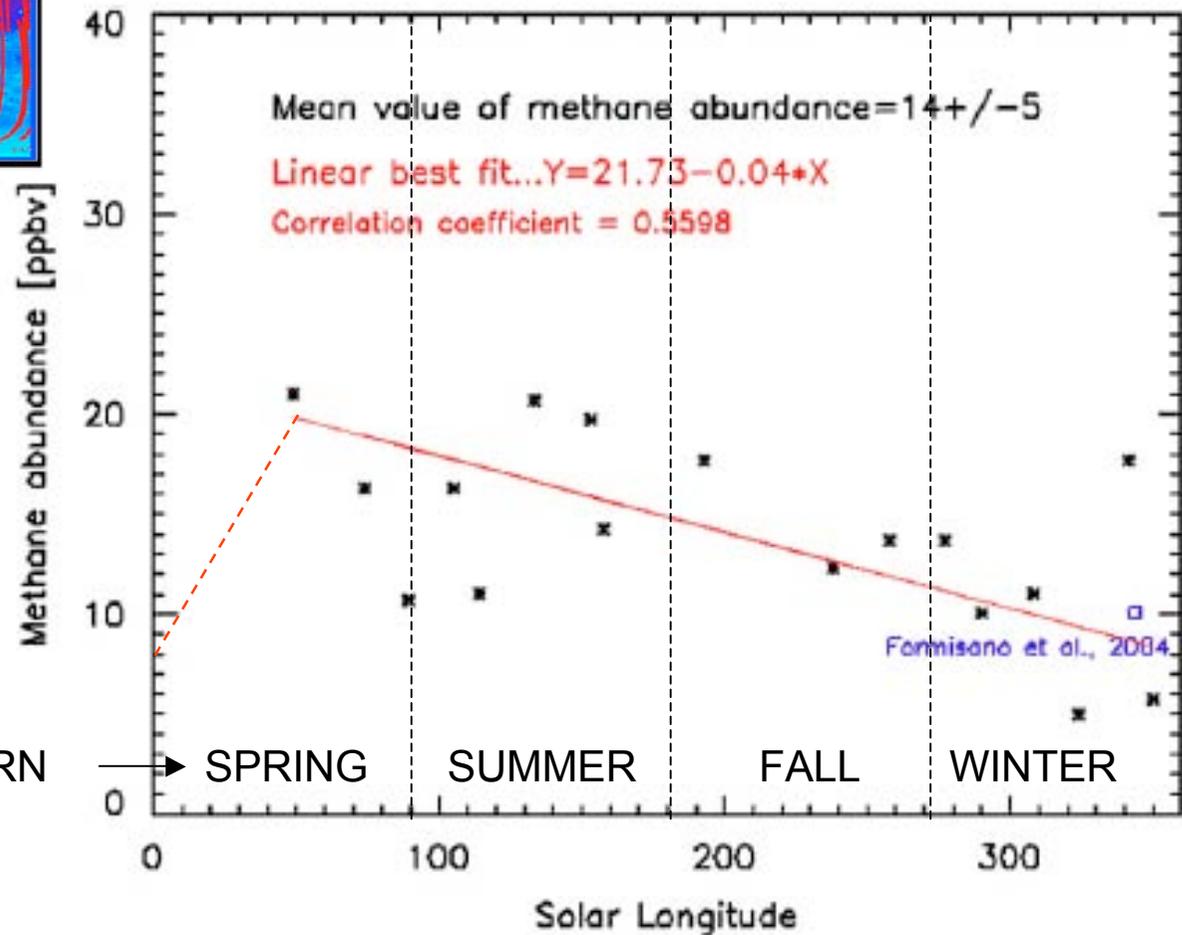


Seasonal variation of methane abundance according to PFS

Formisano et al,
2004

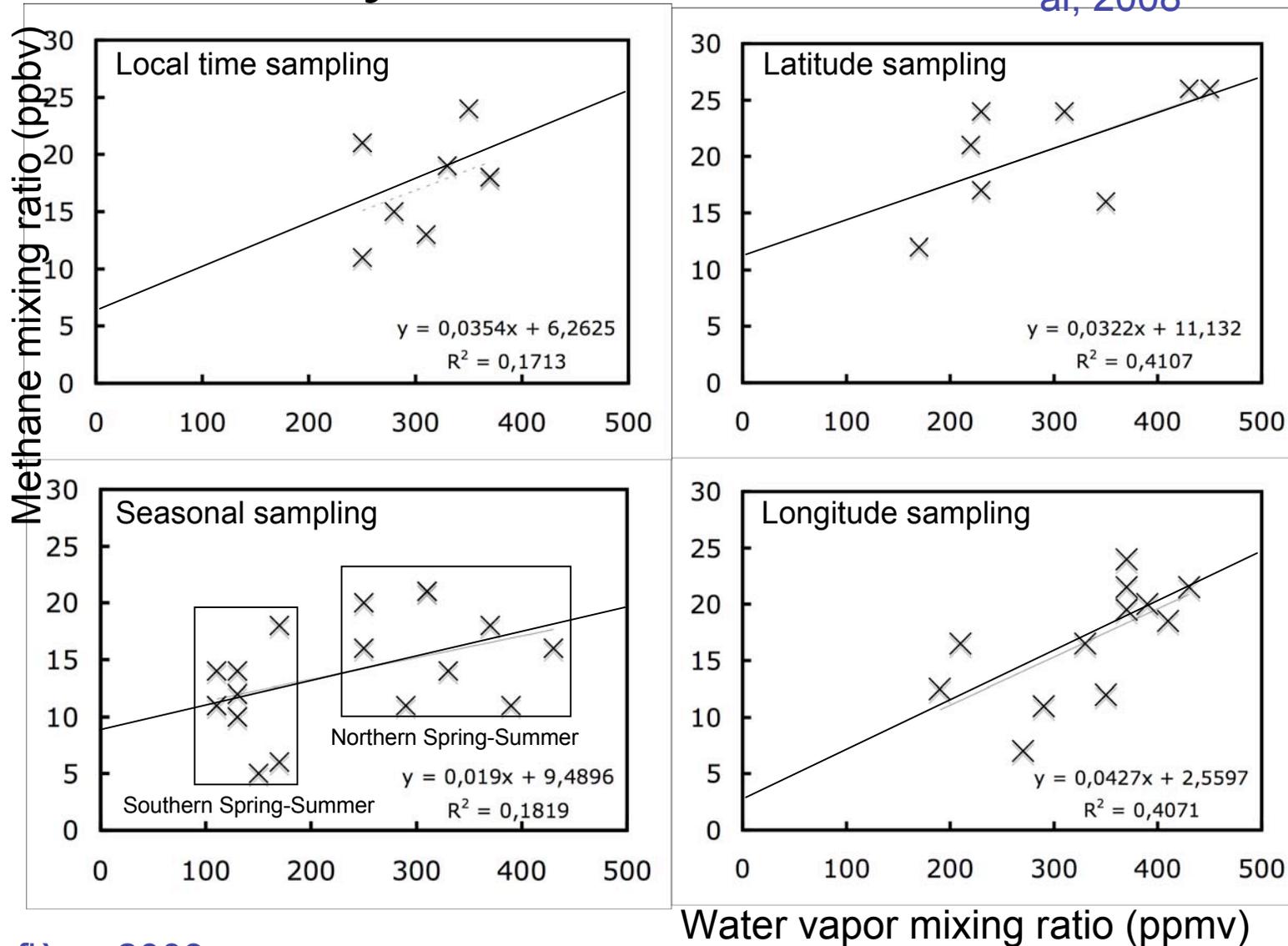
Geminale et al,
2008

NORTHERN

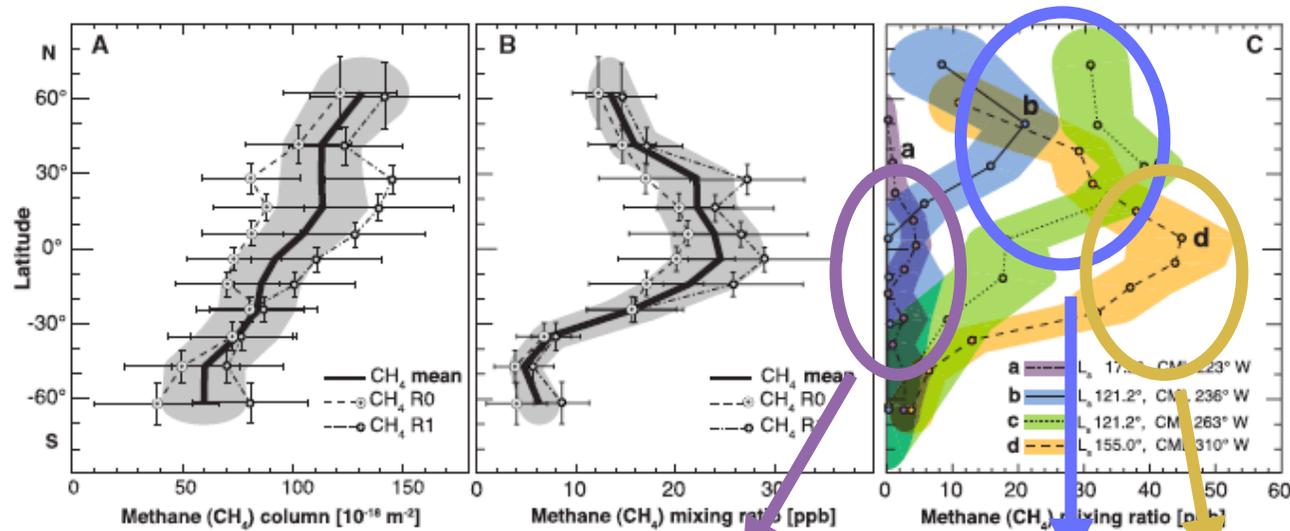


Why a general H₂O/CH₄ correlation, as seen by PFS from orbit...

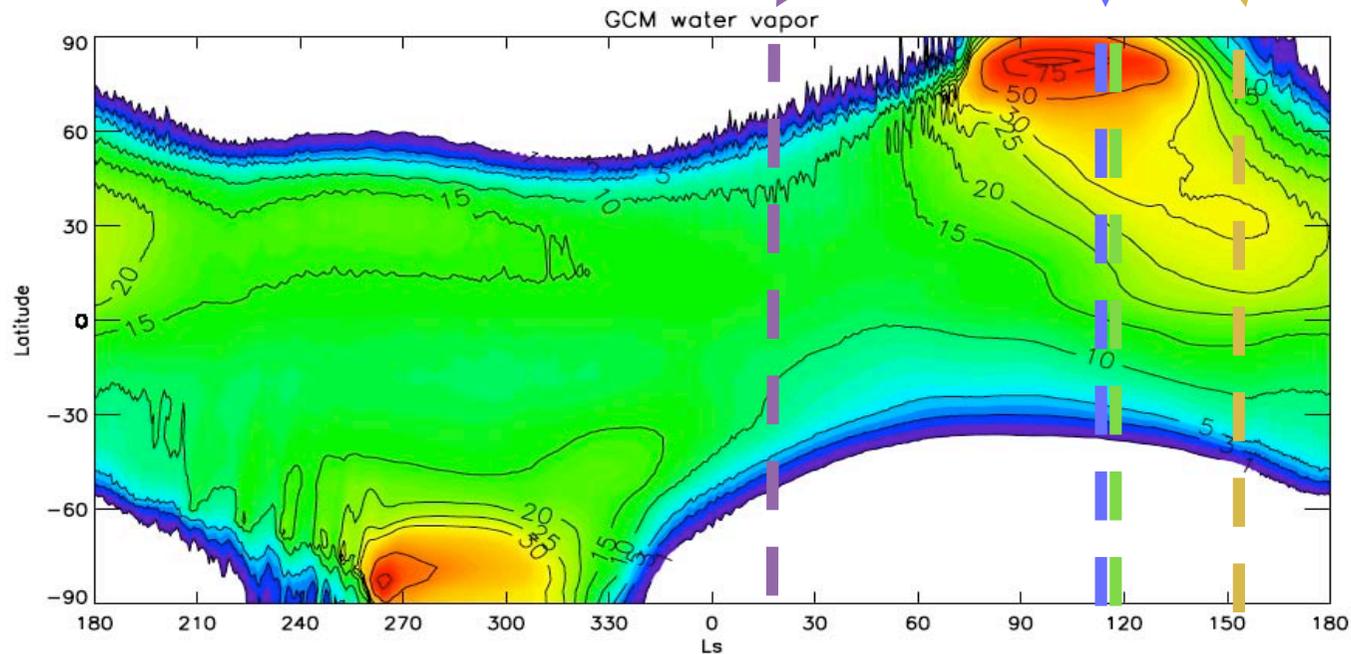
Data from Geminale et al, 2008



... and not contradicted by recent Earth-based measurements?



Mumma et al., 2009



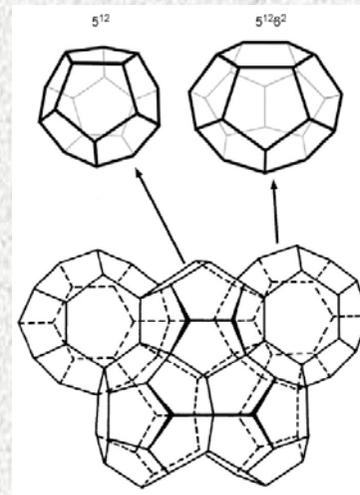
Montmessin et al., 2004

Our hypothesis to solve the mystery

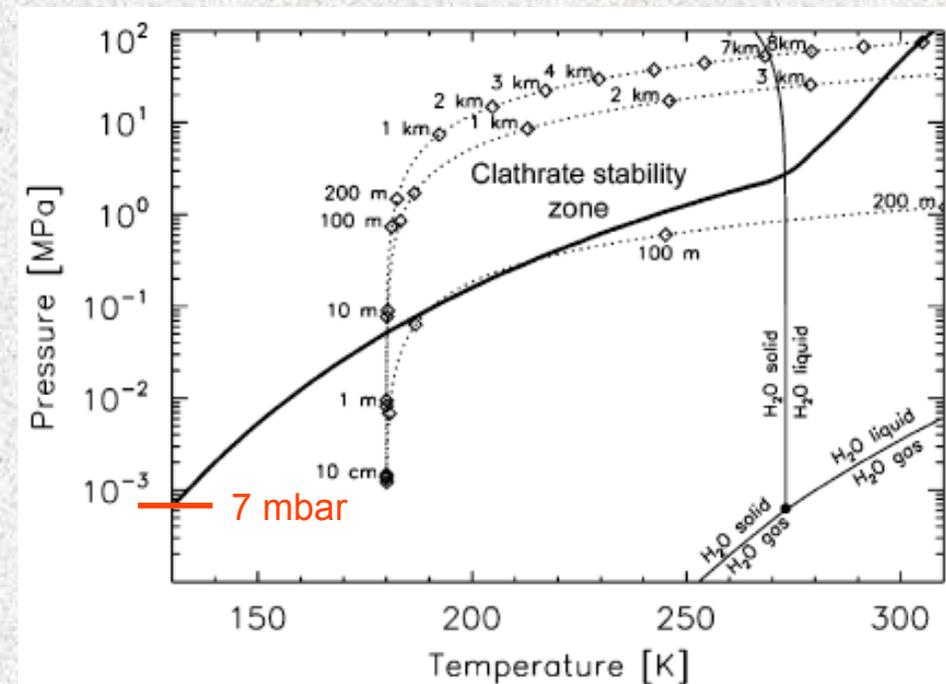
- **How can the atmospheric H₂O distribution, controlled by meteorology, be correlated with the CH₄ distribution if CH₄ is produced by fixed sources at the surface?**
- **One possible solution : gaseous CH₄ is produced in the atmosphere by a process involving H₂O.**
- **If so, the source is necessarily solid :**
 - Methane ice? NO
 - Adsorbed methane? NO
 - Clathrates? POSSIBLE IF STABLE

Methane clathrate hydrate

- CH₄ molecules trapped in water molecules cavities : CH₄-5.85 H₂O.
- Form at high pressure (e.g. ocean groundfloor on Earth), **but NOT in Mars atmosphere.**
- May form at depth in Mars crust, at p>1 bar.

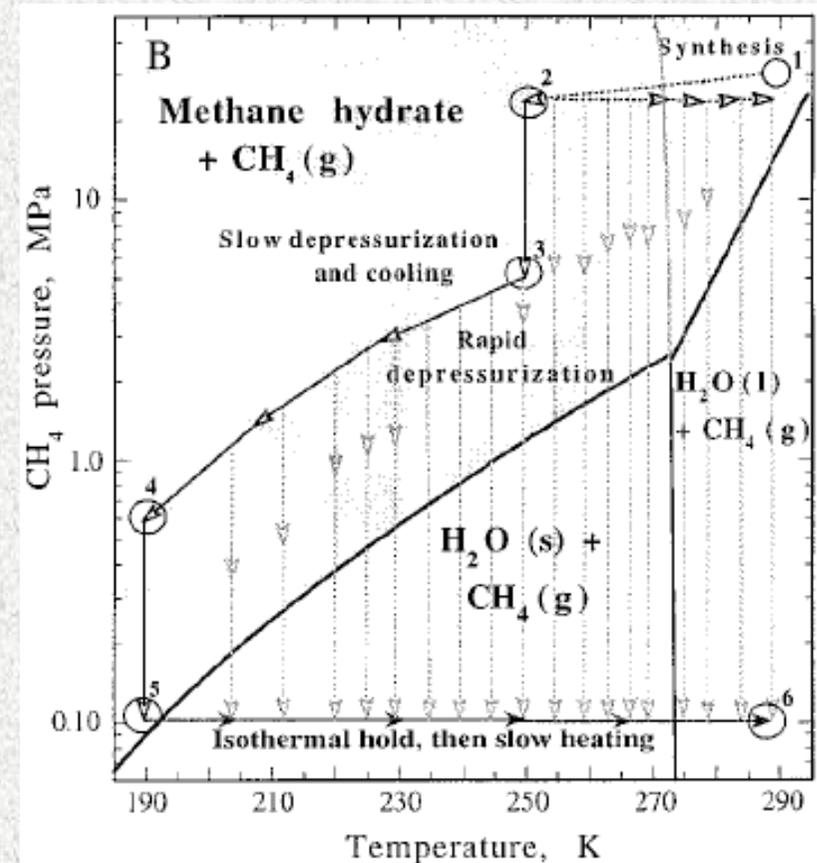
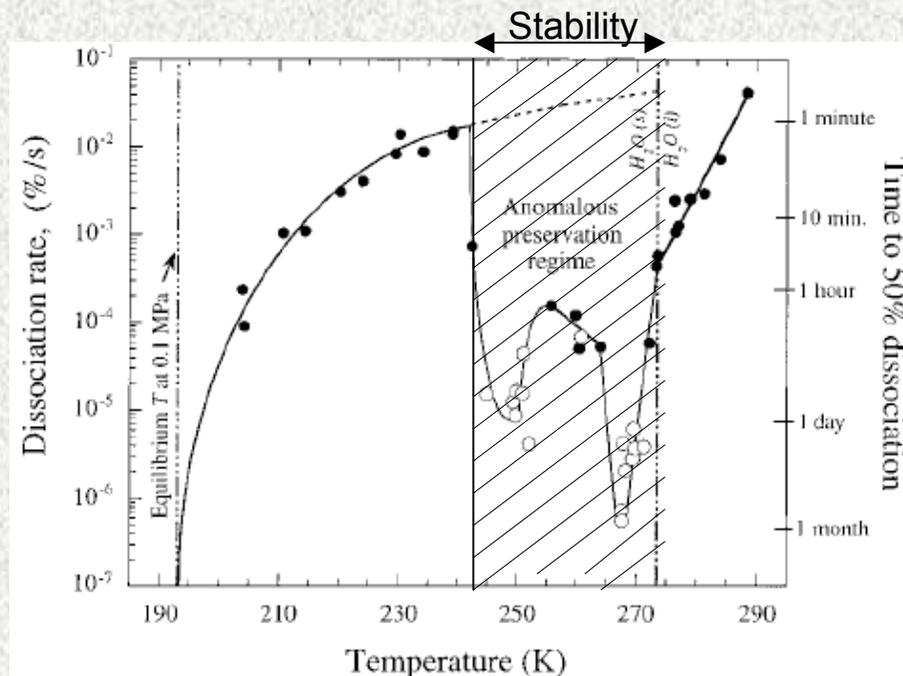


Chastain and
Chevrier, 2007



« Anomalous preservation » of clathrates

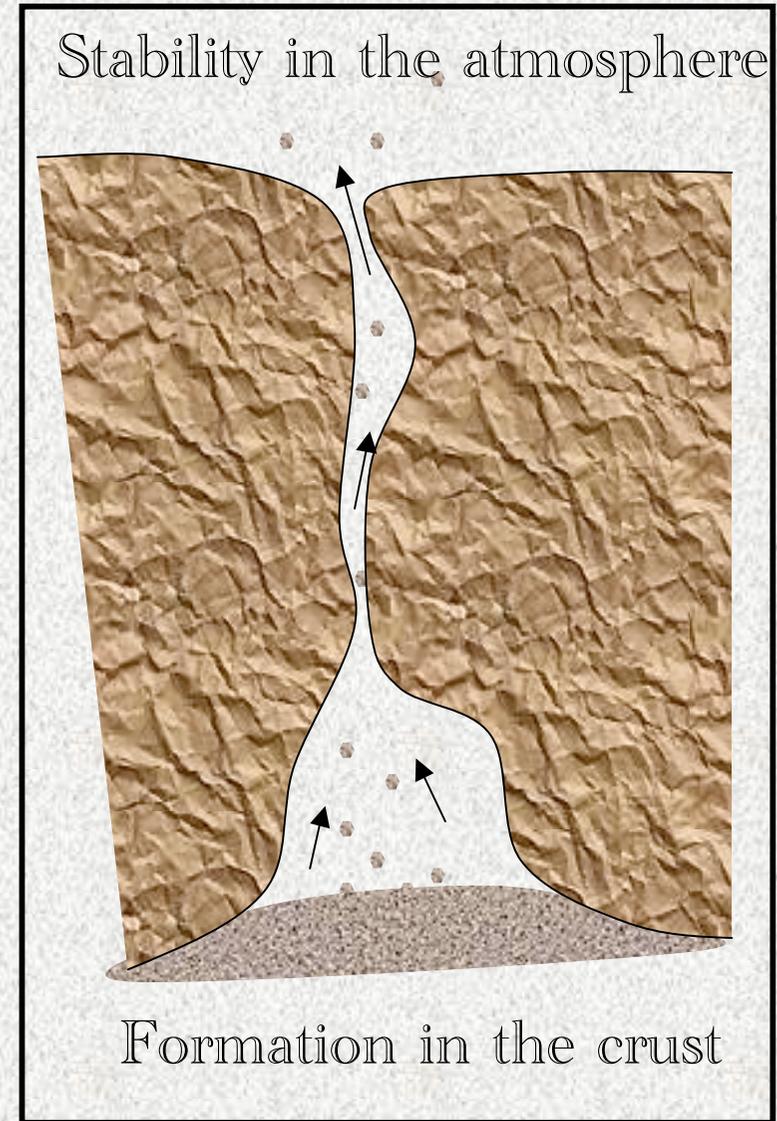
- Anomalous preservation of clathrates (Stern et al. 2001 and ref. therein)



- Ultra-stability of very pure clathrate crystal aggregates (Zhang and Rodgers, 2008): <1% dissociation at 1 bar/268 K during 10 days.

Release of methane clathrate particles

- 1- Formation of clathrate in the crust at high (lithostatic) pressure.
- 2- Fracturation of subsurface by tectonics, yielding local macroporosity.
- 3- Erosion of (friable) clathrate host framework and uplift of particles.
- 4- Release to the atmosphere of (ultra-stable) clathrate particles.

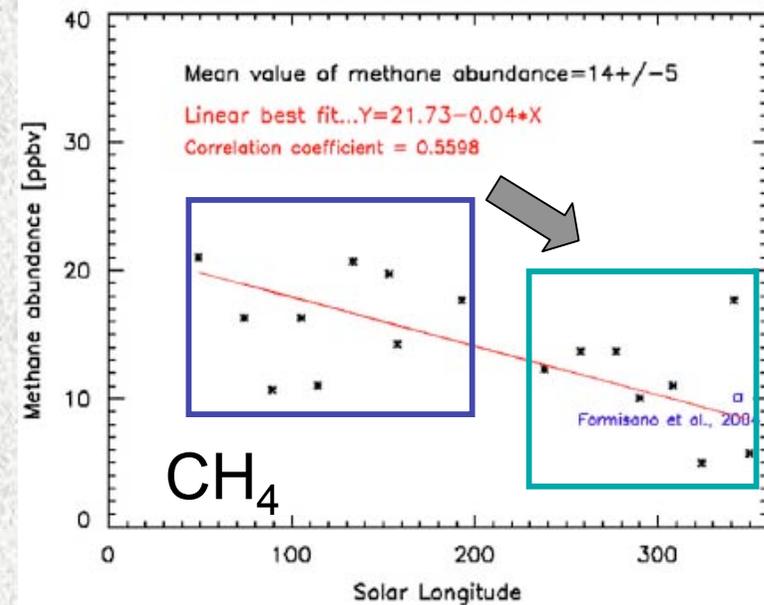
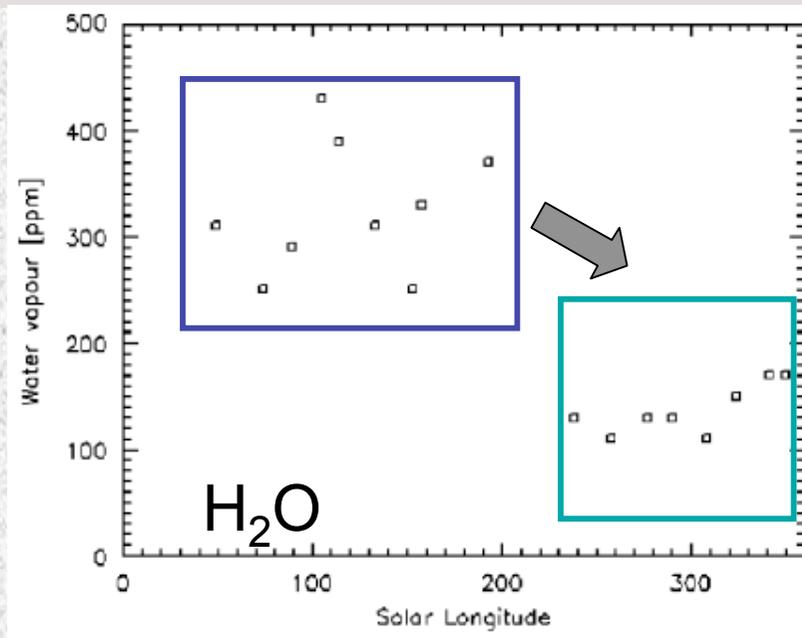


Possible destabilizing effect of water condensation

- Small clathrate particles ($<1 \mu\text{m}$), lifted up to ≈ 50 km, may serve as condensation nuclei (*second mode observed on Viking images by Montmessin et al, 2002?*)
- **Hypothesis : Condensation of H_2O on clathrate particles above the hygropause ($z > 20$ km) \rightarrow destabilization of clathrate to gas phase (by which mechanism : latent heat release, disruption of crystalline structure?)**
- Needs to be studied and validated by experimental works in Martian conditions (never done).

What would our hypothesis explain?

- The seasonal cycle of methane (low level during fall/winter, followed by sharp increase at spring)
- The general H₂O/CH₄ correlation
- A mesospheric source for gaseous methane (as suggested by limb PFS results)



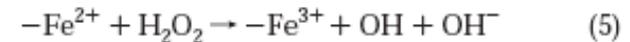
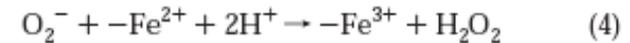
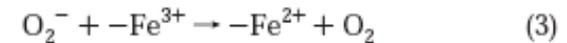
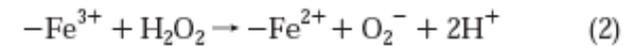
What would our hypothesis imply?

- First observational evidence of clathrates on Mars.
- Some correlation between gaseous CH₄ (*controlled by dynamics*) and mesospheric H₂O (*controlled by both dynamics and condensation*). Not straightforward (**Help, GCM!**)
- A partial or total loss of the memory of the (clathrate) source distribution at the surface due to atmospheric redistribution of clathrates.

Sink of gaseous methane

- Strongly inhomogeneous CH_4 : Lifetime < 200 days (Lefèvre and Forget, 2009).
- The oxidant could be H_2O_2 (or grains oxidized by H_2O_2) in the subsurface.
- Example of Haber-Weiss iron-catalyzed reaction, able to produce OH (but does it work in the absence of liquid water?)

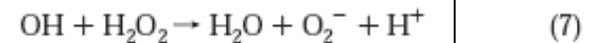
The decomposition of H_2O_2 in soils may proceed through a number of mechanisms, the first of which can be most simply described by reactions 2–5:



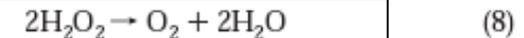
Here, $-\text{Fe}^{3+}$ represents iron in a liganded form or occupying a site at an oxide surface.^{8–10} The sum of reactions 2–5 represents the Fe-catalyzed Haber-Weiss reaction



If the OH generated in reaction 6 reacts exclusively with H_2O_2



the net reaction becomes



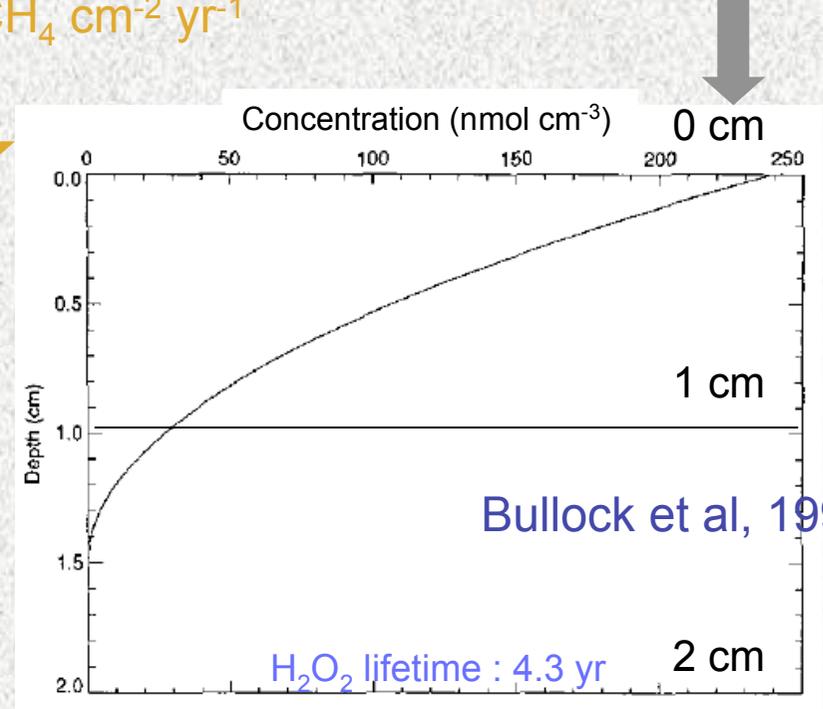
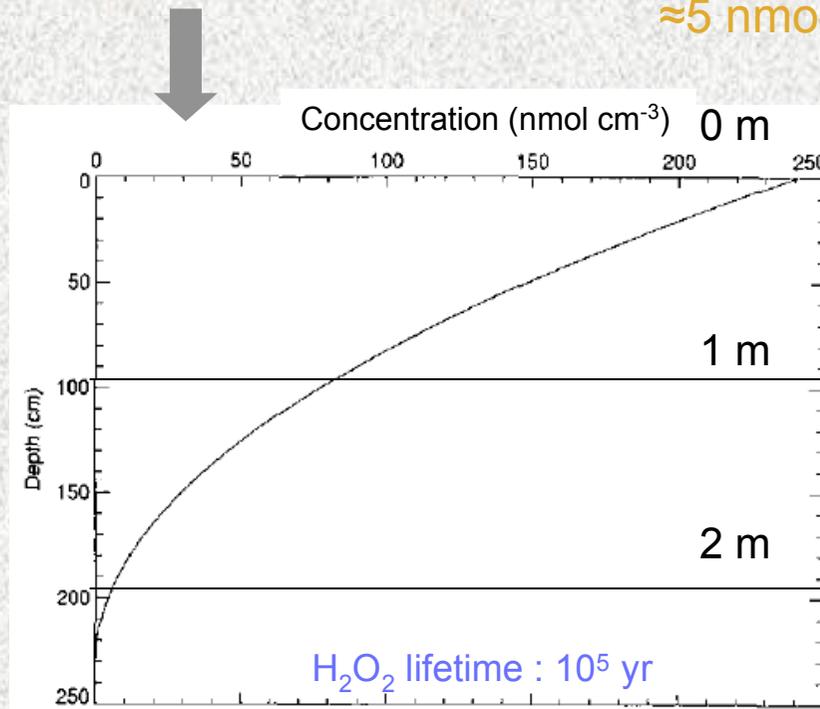
Petigara et al,
2002

Thickness of the regolith oxidation layer

0.1 nmol H₂O₂ cm⁻² yr⁻¹

≈10 ppb CH₄/200 days
 ≈5 nmol CH₄ cm⁻² yr⁻¹

10 nmol H₂O₂ cm⁻² yr⁻¹



Bullock et al, 1994

- Oxidation budget : CH₄ + 4H₂O₂ → CO₂ + 6 H₂O
- 1 CH₄ removes 4 H₂O₂ : ≈20 nmol H₂O₂ cm⁻² yr⁻¹ → 2 yr lifetime for H₂O₂, less than 1 cm depth.
- Centimetric (not metric) oxidation layer

What could be done in the future?

⇒ Laboratory measurements :

- Check and characterize clathrate metastability in Martian conditions.
- Study the decomposition of clathrates in presence of condensing H₂O.
- Study the kinetics of methane oxidation by adsorbed H₂O₂ in presence of iron-rich minerals (Haber-Weiss-type reaction)

⇒ Modelling :

- Decomposition of clathrate particles in presence of condensing water.
- GCM simulations of methane cycle (coupled with water cycle).

⇒ Space observation :

- Search for and characterize from future orbiters the second mode particle population (<1 μm) in Martian atmosphere.
- Characterize from future landers the oxidation state of the superficial regolith, with a sub-centimetric vertical sampling capability