# Laboratory investigation of the role of heterogeneous processes in Martian methane variability



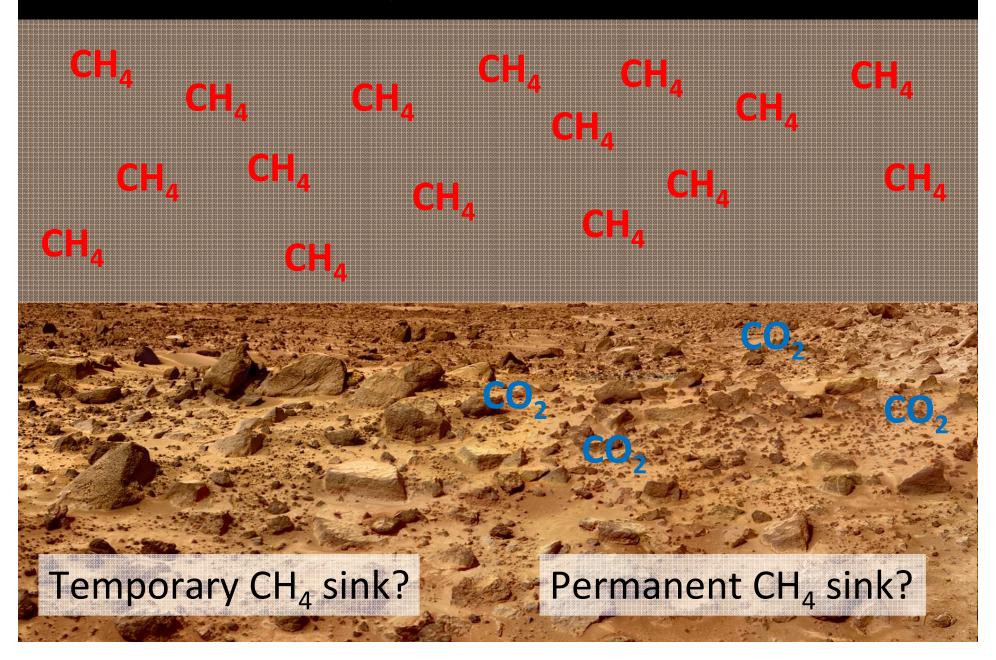


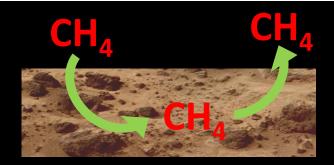
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Workshop on Methane on Mars, Frascati, Italy 25 November, 2009



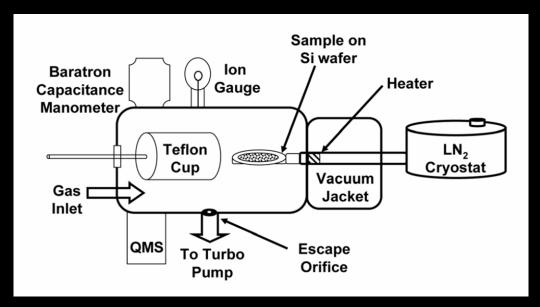
#### Possible heterogeneous chemistry on Mars?





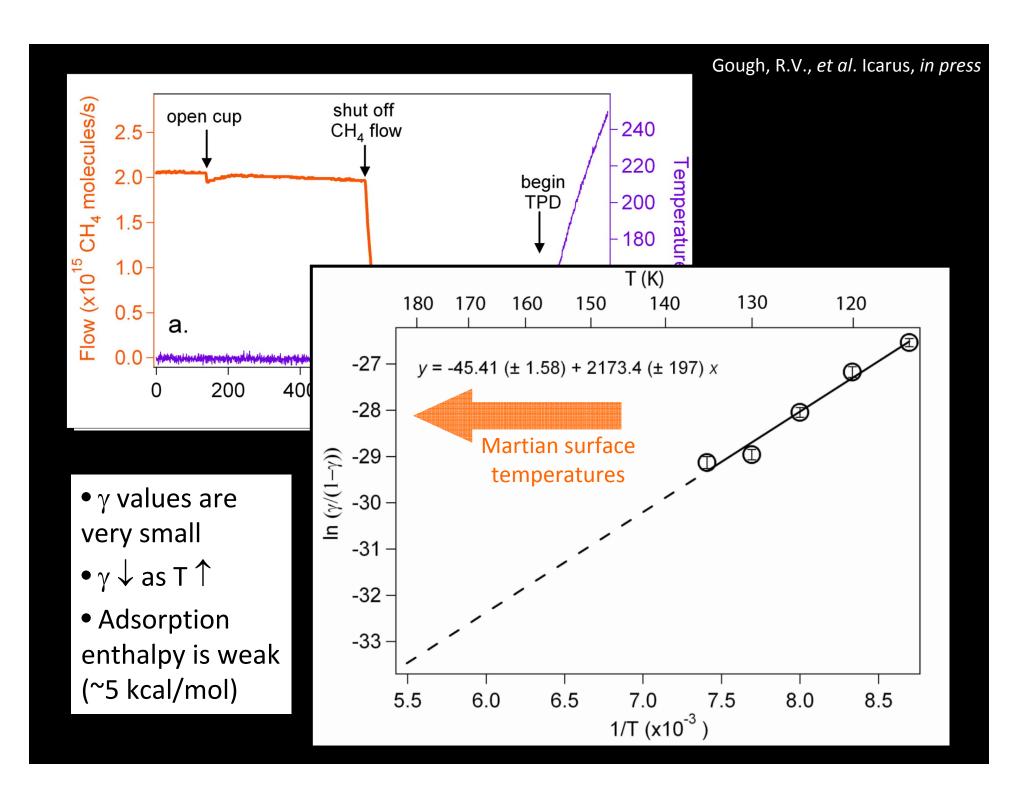
## Temporary CH<sub>4</sub> sink?

 Physisorption of CH<sub>4</sub> on JSC-Mars-1 was studied using a high-vacuum chamber (Knudsen cell)

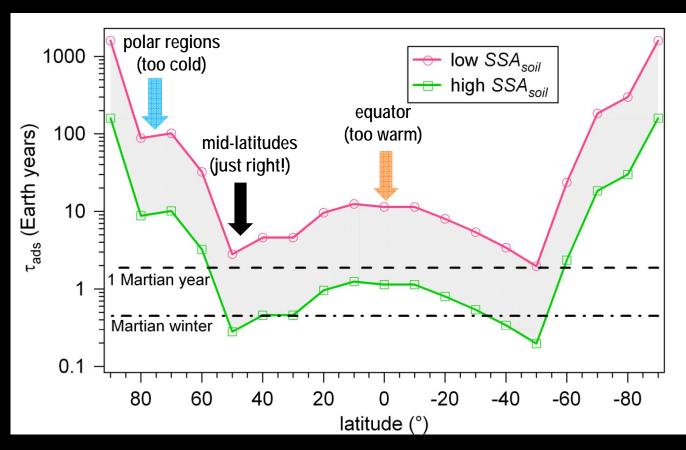


• Uptake coefficient ( $\gamma$ ) values were determined as a f(T).

$$\gamma = \frac{\text{# of molecules that stick}}{\text{# of collisions with the surface}}$$

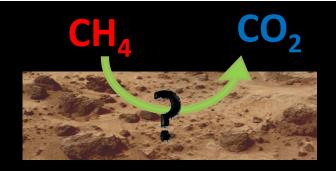


## Implications for Mars?



(Using annual mean and min. subsurface T's from Tokano et al.; ice table depths from Mellon et al. 2004)

- Mid-latitudes are most likely to be a temporary CH<sub>4</sub> sink/source
- More extensive simulations accounting for subsurface variability and atmospheric transport are warranted (Meslin et al., this meeting)



## Permanent CH<sub>4</sub> sink?

TiO<sub>2</sub>•H<sub>2</sub>O<sub>2</sub> (Quinn and Zent, 1999)

JSC-Mars-1 +  $H_2O_2$ (Levin and Straat, 1981)

Na<sup>+</sup> and Mg<sup>2+</sup> perchlorate



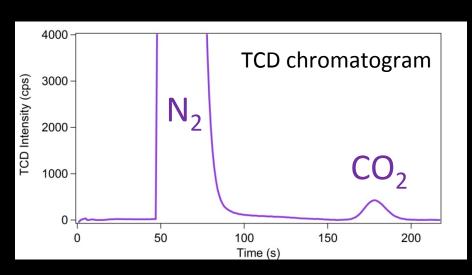




TiCle + H<sub>2</sub>O -> TiO les, we determined the reactivity toward the "Viking organics" as well as CH<sub>4</sub>

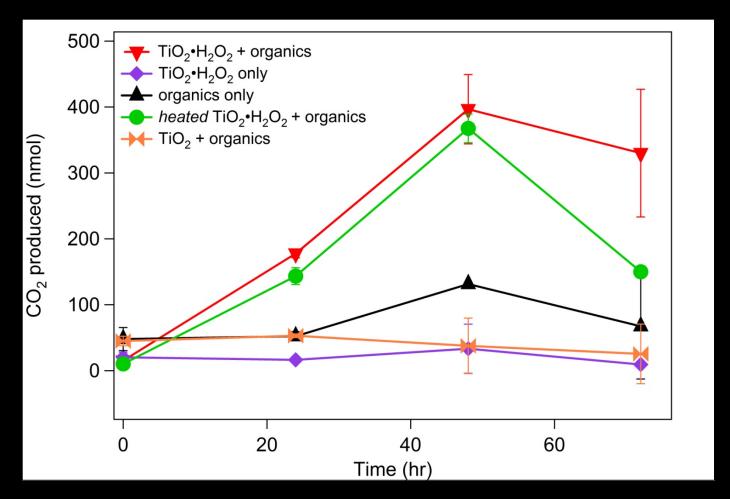
#### Experimental technique

- Samples in vials with N<sub>2</sub>
  atmosphere were kept at 2°C
- Headspace was sampled, analyzed with gas chromatography (GC)



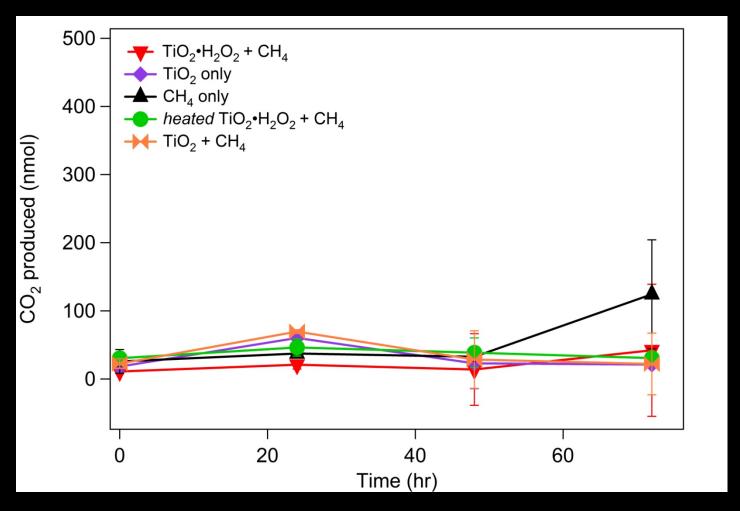
- After initial (t=0) measurement, organics or methane were added
- GC measurements taken at 24, 48, 72 hrs
- Several controls were used to rule out contamination

# Results: TiO<sub>2</sub>•H<sub>2</sub>O<sub>2</sub> + organics



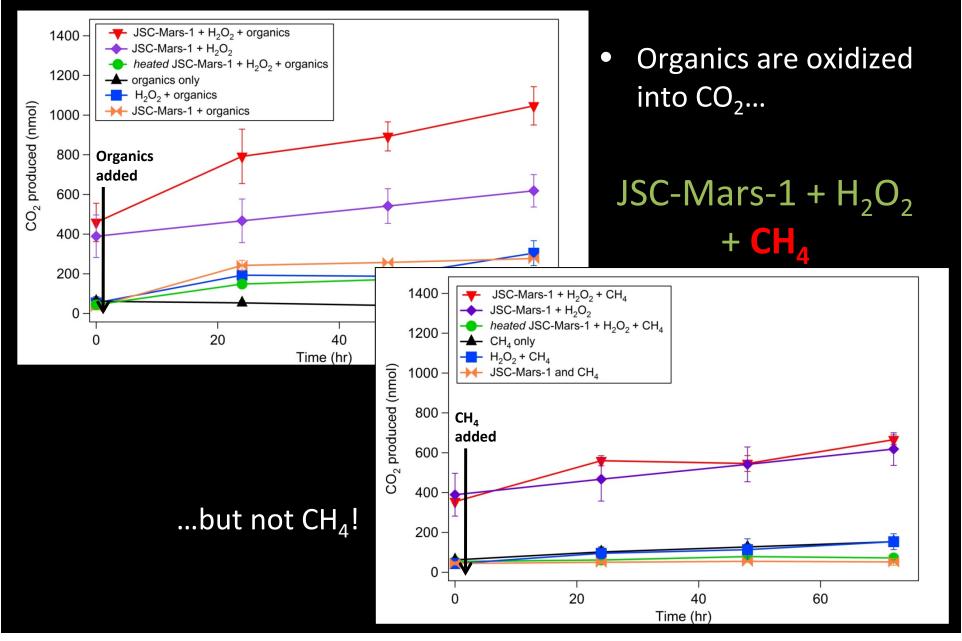
- Material is able to oxidize organics
- Results similar to Quinn & Zent (1999) (and Viking LR exp)

# Results: TiO<sub>2</sub>•H<sub>2</sub>O<sub>2</sub> + CH<sub>4</sub>

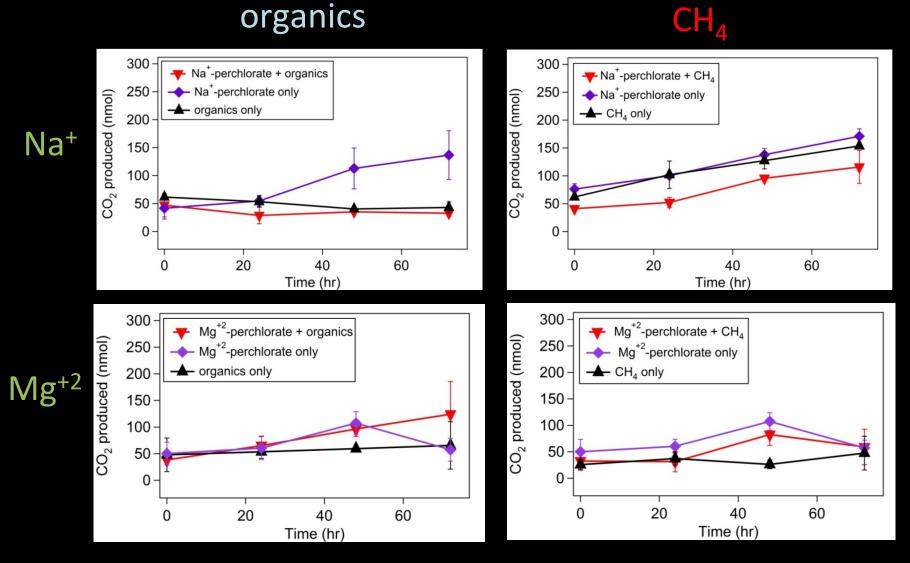


No formation of CO<sub>2</sub> from CH<sub>4</sub>

#### JSC-Mars-1 + $H_2O_2$ + organics



## Na<sup>+</sup> and Mg <sup>2+</sup> perchlorate salts



Nothing is oxidized by perchlorate salts!

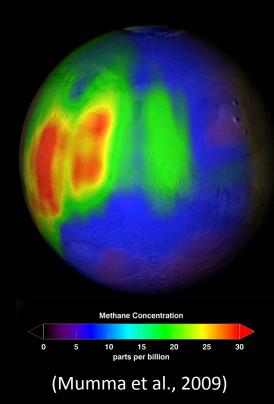
#### Martian implications

$$CH_{4(g)} \xrightarrow{k_{ox}} CO_{2(g)}$$

$$\frac{-d[CH_4]}{dt} = k_{ox}[CH_4]$$

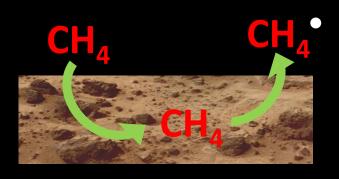
- Detected no CH<sub>4</sub> oxidation by any of the 3 sample types studied
- Can determine **upper limit rate constant**  $(k_{ox})$  as we know  $CH_4$  conc. in vials,  $[CH_4]$ , as well as upper limit of oxidation rate,  $d[CH_4]/dt$ .
- For all 3 "oxidizing surfaces",  $k_{ox} < 4.94e-10 s^{-1}$ .

## Martian implications



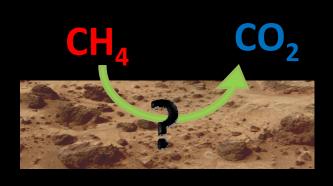
- What size  $k_{ox}$  would be needed to explain recent CH<sub>4</sub> observations? (Mumma et al., 2009)?
  - Global average 6 ppbv in 2003 → 3 ppbv in 2006
- If loss caused by oxidation, rate constant  $(k_{ox})$  must be at least ~3e-9 s<sup>-1</sup>
- Compare to experimental upper limit value:  $k_{ox} < 4.94e-10 s^{-1}$
- Levevre & Forget (2009): surface loss must be very fast (3-6 hr) to explain observations!
- Heterogeneous loss of CH<sub>4</sub> probably not important on the 3 surfaces studied

#### Conclusions



#### **Temporary CH**<sub>4</sub> sinks:

- Temperature-dependent adsorption of CH<sub>4</sub> in the regolith could play a role.
- Could explain several recent observations

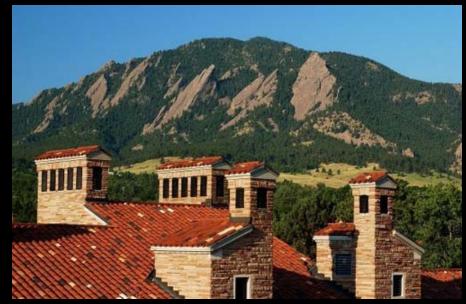


#### Permanent CH<sub>4</sub> sinks:

- Several proposed oxidative surfaces were studied
  - $TiO_2 \bullet H_2O_2$ , JSC-Mars-1 +  $H_2O_2$ , perchlorate
- No CH<sub>4</sub> oxidation was detected; rxn on similar surfaces on Mars is likely very slow
- Observed atmospheric loss possibly to other processes.

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