

Methane and the Martian Habitability

Sushil Atreya

Paul Mahaffy

Chris Webster

Patrice Coll

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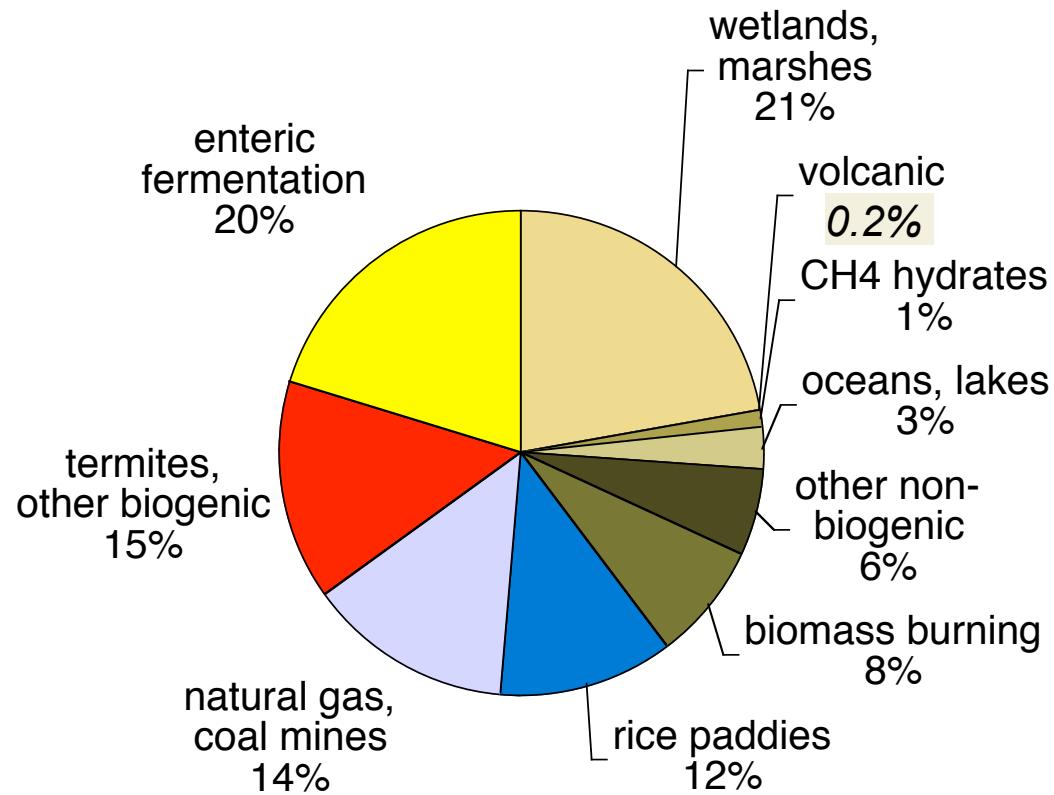
map

- *Formation* (i.e., origin, not source)?
- *Destruction*: how, where, and how fast?
- *Supply*: venting flux small or big?
- *Habitability*: tests?
- *Measurements*?

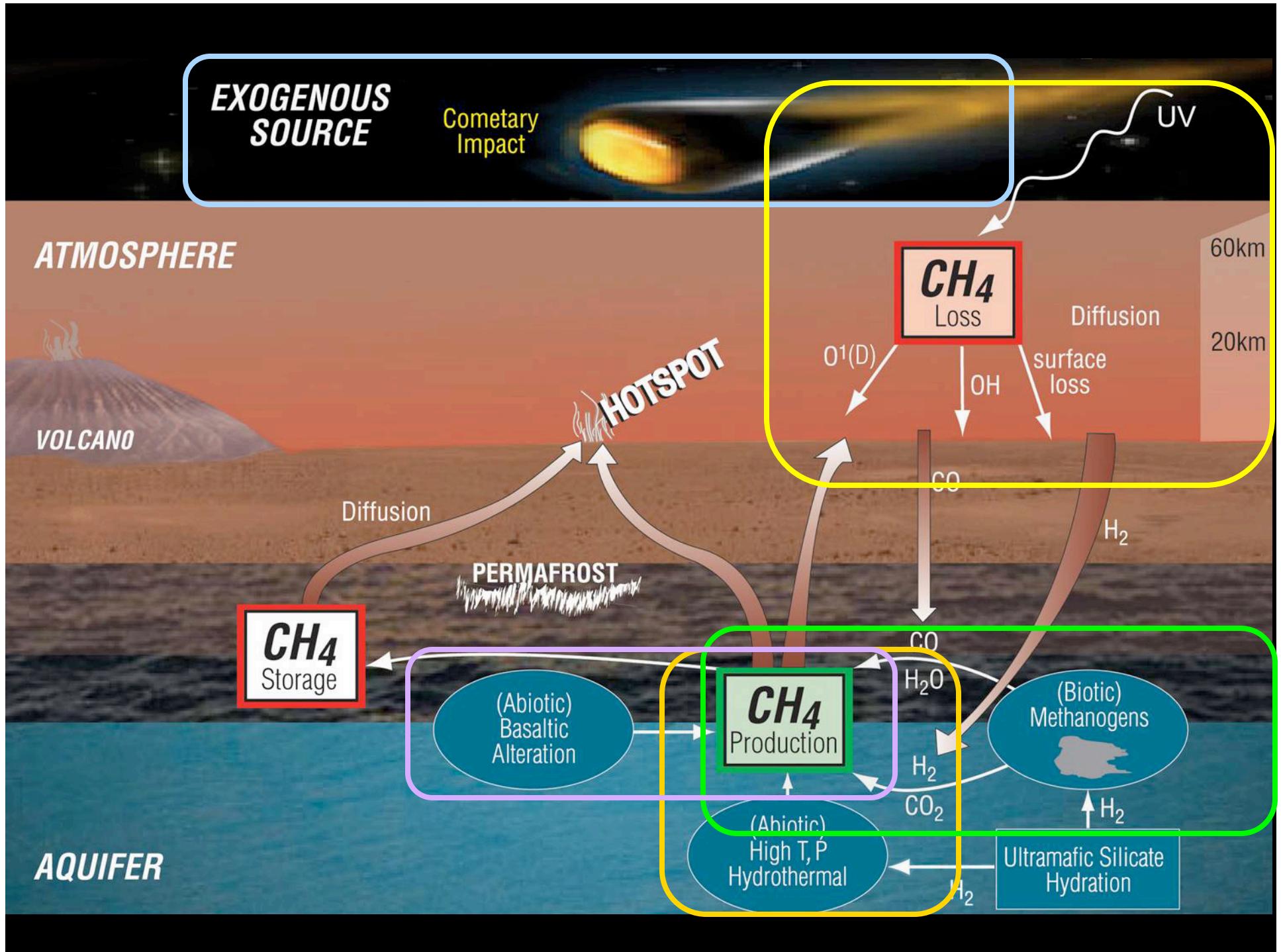
mars atmosphere composition

species	mixing ratio	minor species	<i>upper limit</i>
CO_2	95.32 %	N_2O	1 (-7)
N_2	2.7 %	NO_2	1 (-8)
^{40}Ar	1.6 %	C_2H_2	2 (-9)
O_2	0.13 %	C_2H_4	5 (-7)
CO	0.07 %	C_2H_6	4 (-7)
H_2O	1 - 100 pr- μm	SO_2	1 (-9)
O_3	0.01 – 0.8 ppm	H_2S	1 (-7)
H_2	15 ppm	HCl	2 (-9)
2003 H_2O_2	20-40 ppbv (<i>atmos.</i>)		
2003 CH_4	15 ppbv (global)		surface pressure 6-10 mb

(Earth CFC's ~1 ppbv)

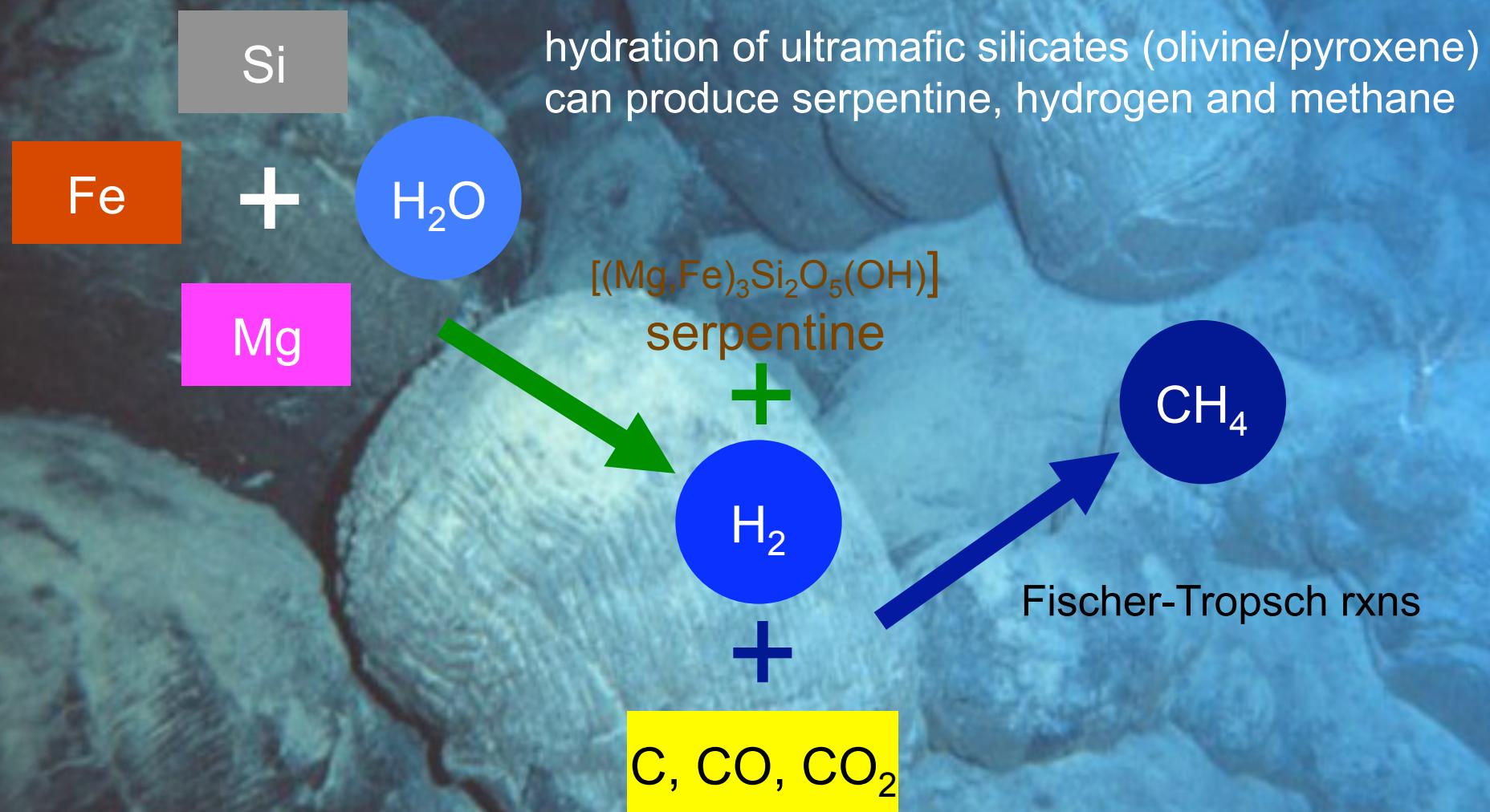


methane sources on Earth (1750 ppbv)

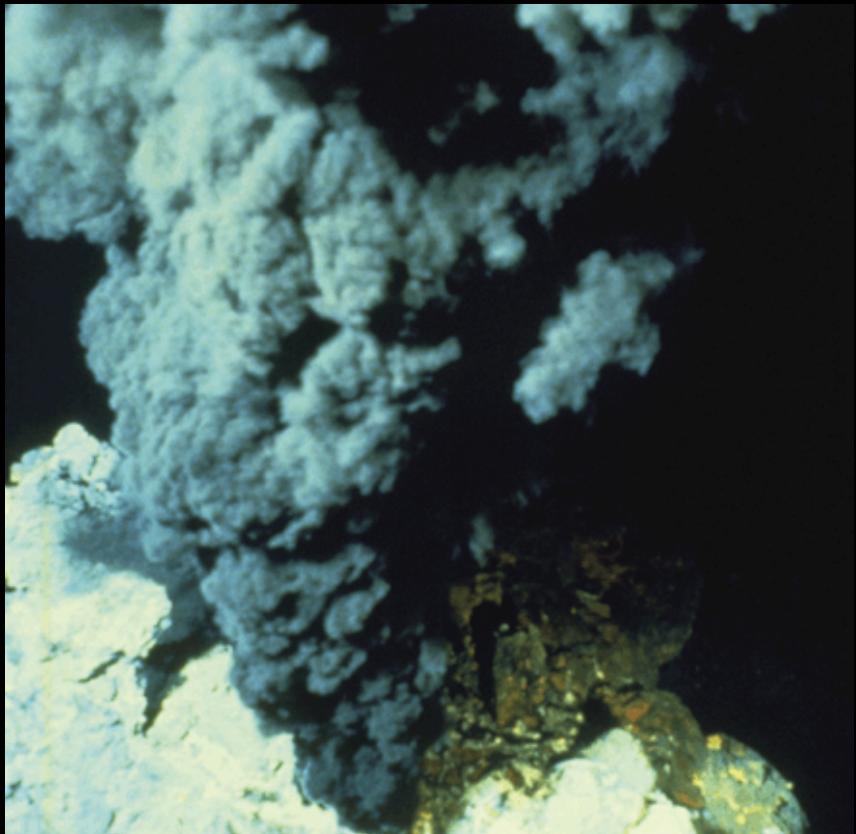


how: serpentinization

liberate hydrogen, mix carbon, make (abiotic) methane



where: Black Smoker hydrothermal vents
abiotic CH₄ by high temperature serpentinization



Mid Atlantic Ridge

300-500 C, sulfur, highly acidic (~lemon juice)

Juan De Fuca Ridge
depth 2222 m
exit temp 342 C
pressure 200 bar
chimney ht. 10 m

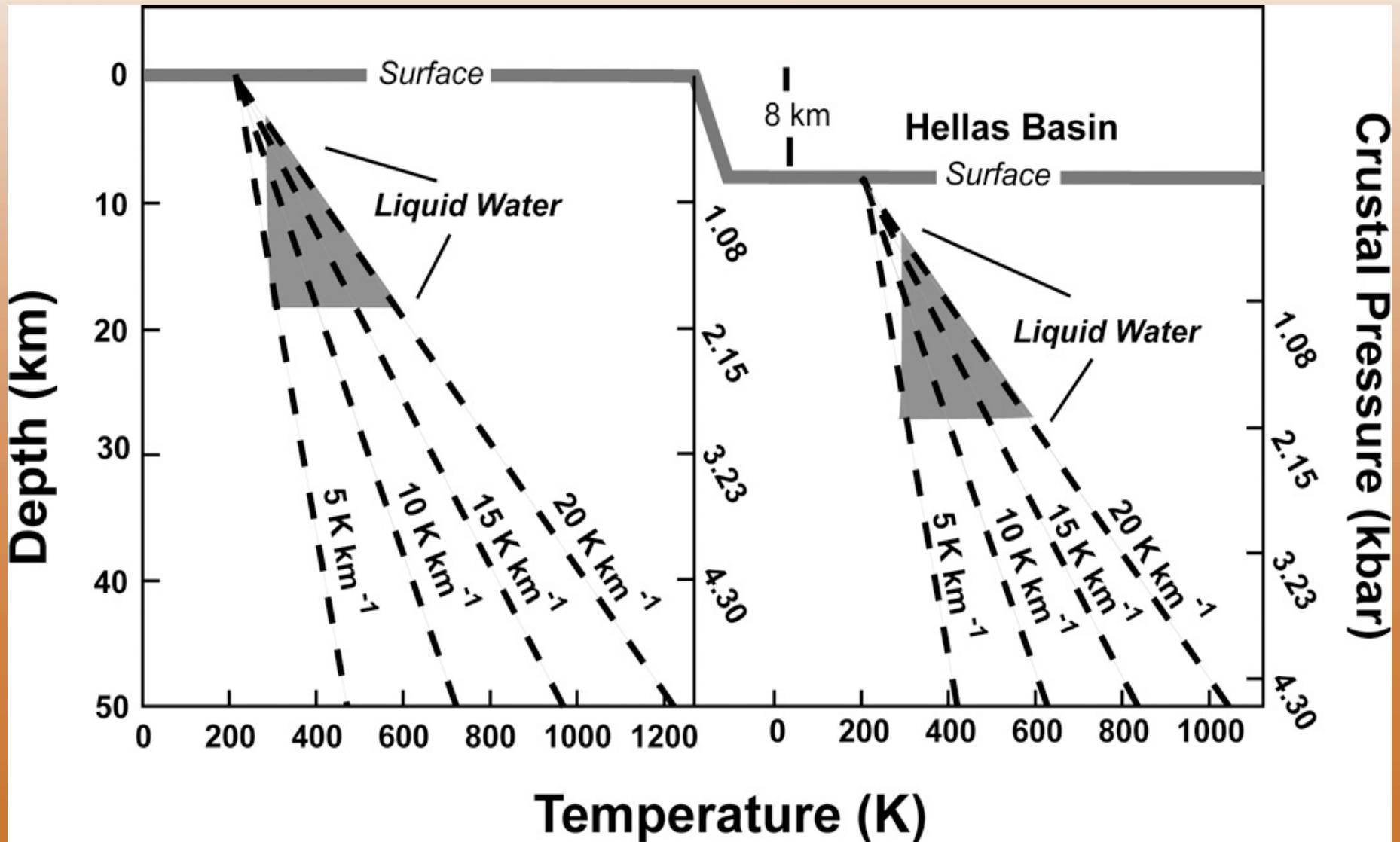


where: Lost City
abiotic CH₄ by low temperature serpentinization



- 15 km from Mid_Atlantic Ridge
- 30-90 C
- highly alkaline (ammonia, m.o.m.)
- 20m high carbonate towers
- little sulfur minerals

mars geotherms: liquid water stable in 2-25 km, 20-300 C



(Oze, Sharma, 2005)

methane origin, now or in the past

- **Geologic** (or hydrogeochemical)
i.e. water-rock chemistry, involving serpentinization and metal-catalyzed Fischer-Tropsch type reactions, or
- **Biogenic** (microbial methanogenesis)
i.e. by enzymatic reduction of CO₂, or fermentation of organic matter by microbes in the metabolic process

*Liquid water essential to both
(subsurface aquifers, if CH₄ produced recently)*

source strength: shorter life, bigger flux

- **Photochemical loss:** long life, small source flux
 - 300-600 yrs, ~200 ton/yr → uniform methane
- **Reactive surface:** faster loss, i.e. ~one year lifetime, requiring much bigger source flux
- **Surface loss due to oxidants**
 - **peroxide** (very reactive), i.e.
resulting hydroxyl, hydroperoxy (HO_2), superoxides
 - **perchlorates** (very stable)

organics and oxidants don't mix

- ▣ **Viking (1976) Life Sciences Experiments (LSE)**
 - ▣ GCMS: no organics found in Martian soil --
 - ▣ Indigenous; or externally delivered by meteorites, IDPs, comets
 - ▣ GEX: nutrients & H₂O added, O₂ released rapidly:
surface oxidant ; required amount **20-250 ppm H₂O₂**
(Oyama, 1977; Hunten, 1979; Huguenin, 1982)
- ▣ In 2003, **20-40 ppb H₂O₂** in **atmosphere** detected, but is **too low** by a factor of **100-1000** compared to LSE requirement

[H₂O₂ detection: submm 362.156 GH,
Clancy et al.; IR (8.04 – 8.13 µm) Encrenaz, et al.]



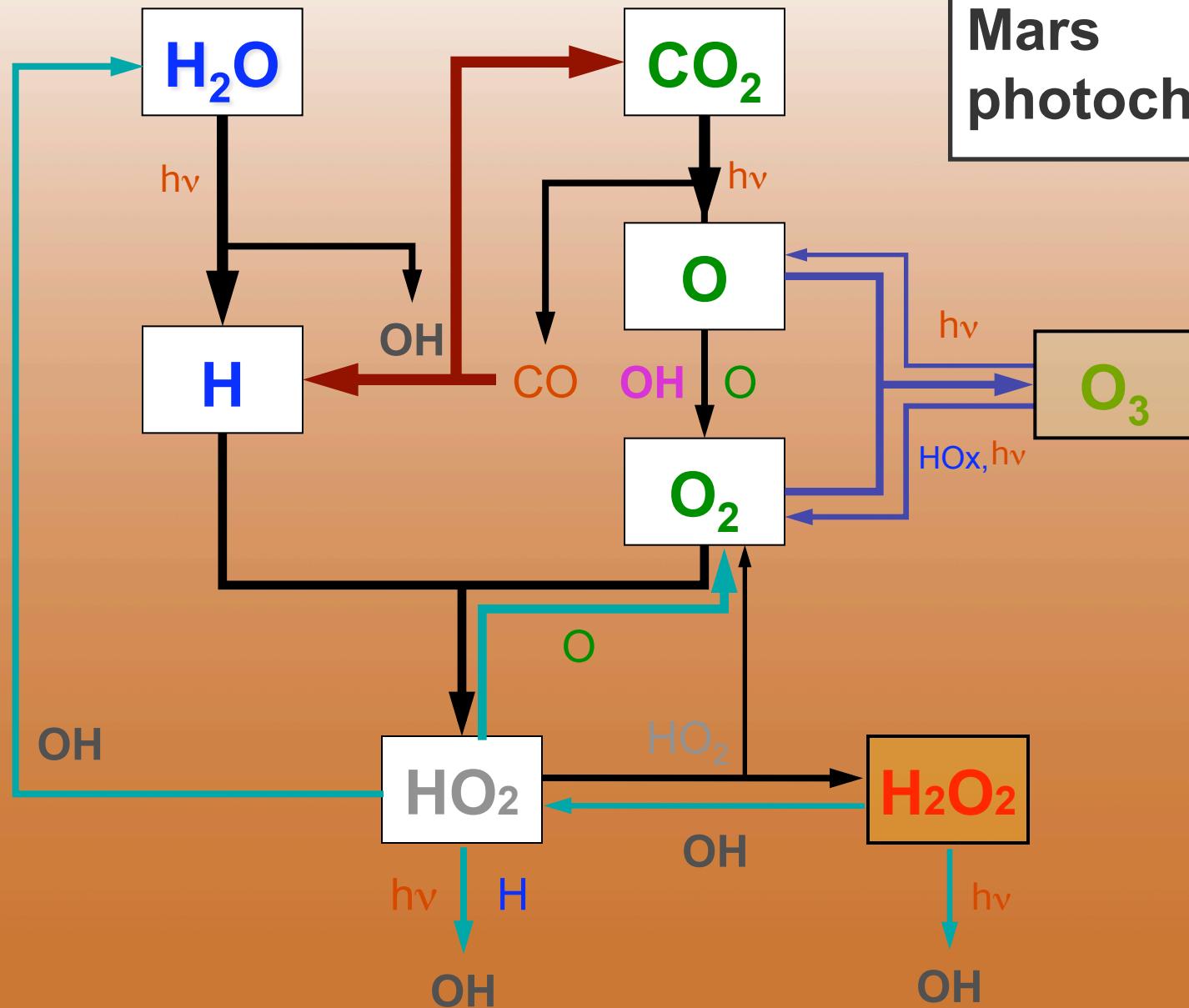
storm electric fields (triboelectricity) change near-surface chemistry

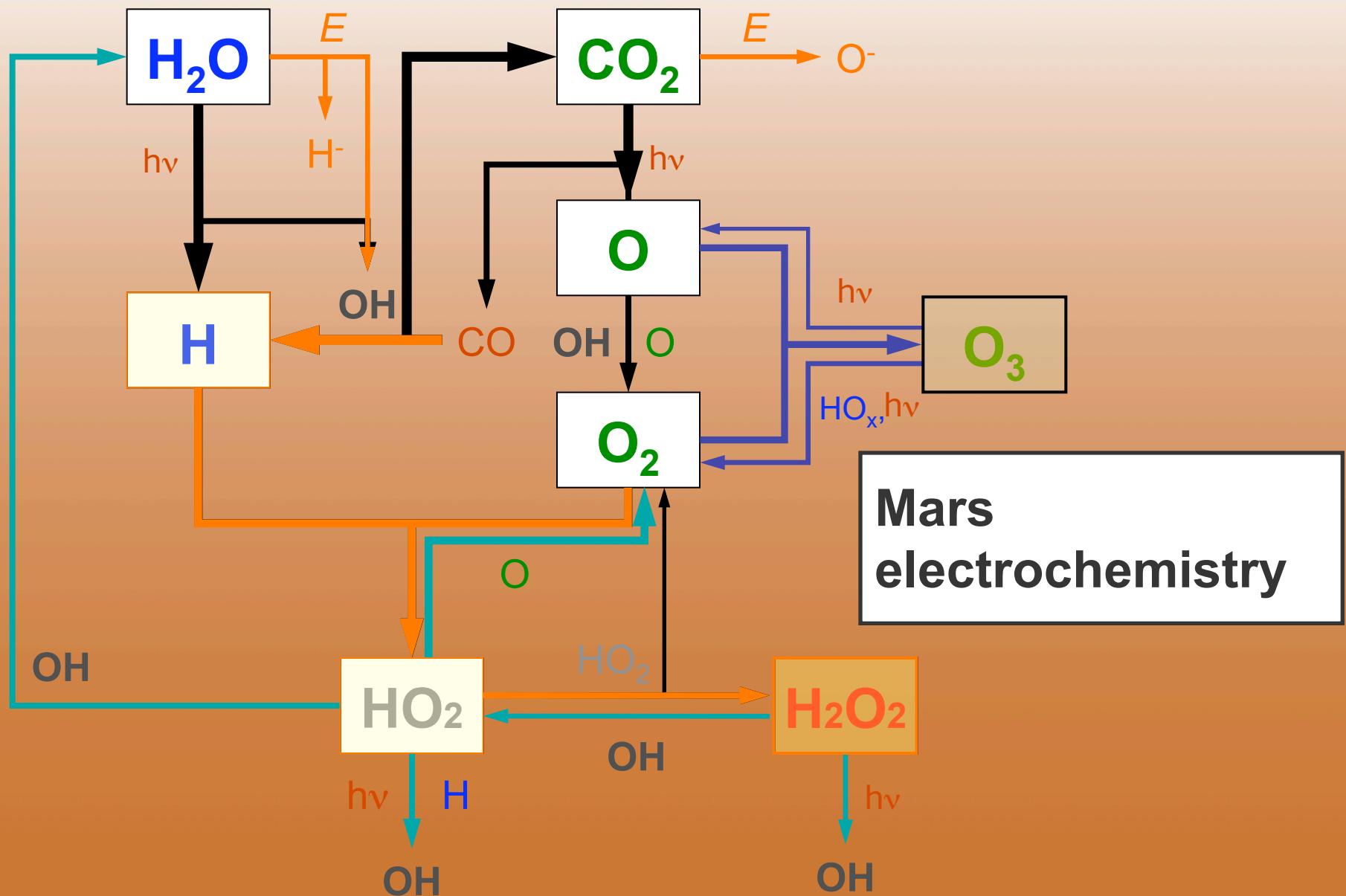
- 5-25 kV/m* electrostatic fields generated in Martian aeolian processes, i.e. in convective dust storms and dust devils
- CO/O-, OH/H- ion pairs* produced→
- H₂O₂ enhanced substantially by electrochemistry over photochemical source**

*Delory, et al. 2006, Farrell et al. 2007, Jackson 2009

**Atreya, et al. 2006

Mars photochemistry





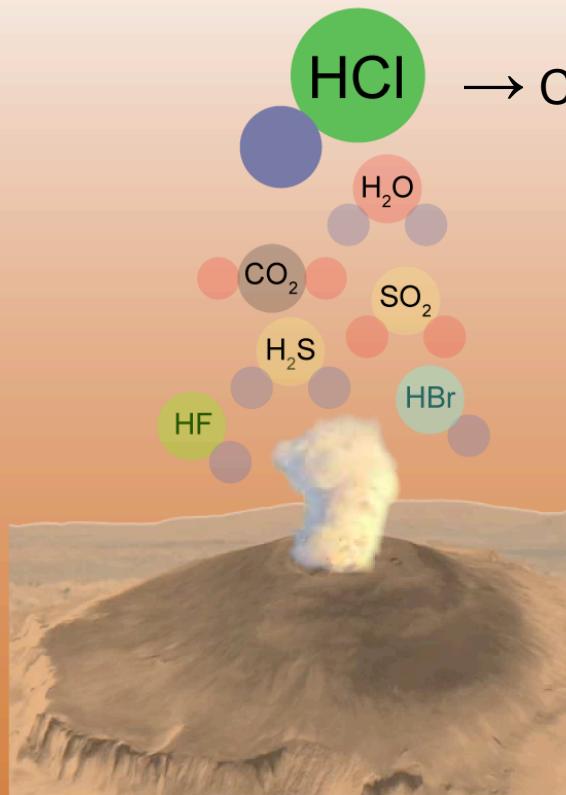
storm electric fields make peroxide; peroxide & products destroy surface organics and methane

- 10,000x enhancement of H_2O_2 over photochemical source ** →
- H_2O_2 “snow” → H_2O_2 into regolith
- H_2O_2 in regolith lives long, up to millions of years (M. Bullock, 1996), compared to *less than* one day in atmosphere
- *superoxides/ hydroxyl/ hydroperoxy (HO_2) generated in soil*
- **rapid destruction of surface organics (directly), and of**
- **methane, by OH , HO_2 and superoxides from H_2O_2**
- **excess CO recycles the CO_2 lost to triboelectricity**

Radiolysis of ice also produces H_2O_2 , as on Europa, Ganymede, Callisto

**Atreya, et al. 2006

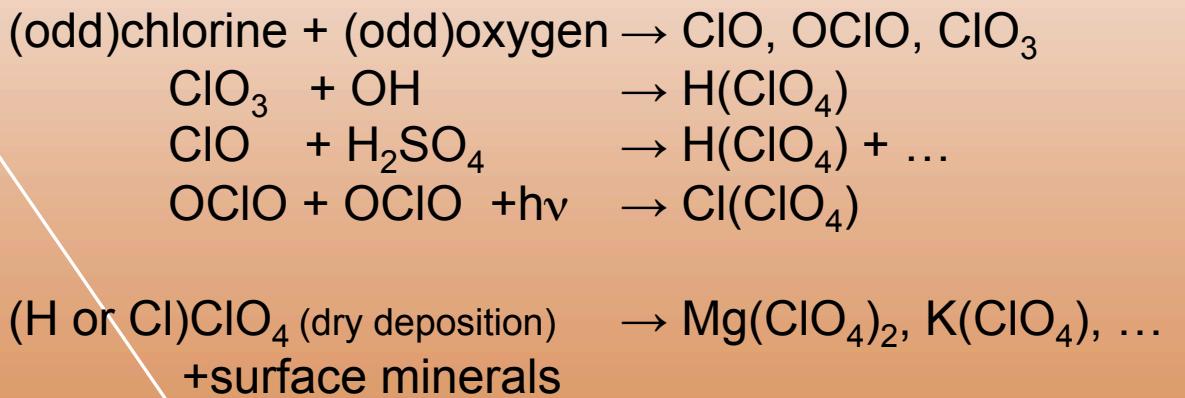
perchlorate (ClO_4^-) on Mars and Earth



volcano

where
how much
dry period

Mars
north polar region
0.3 – 0.6 wt%
?



brine

Earth
Atacama, Bolivia, Texas
0.03 – 0.6 wt% (Atacama)
5–15 Myr (Atacama)

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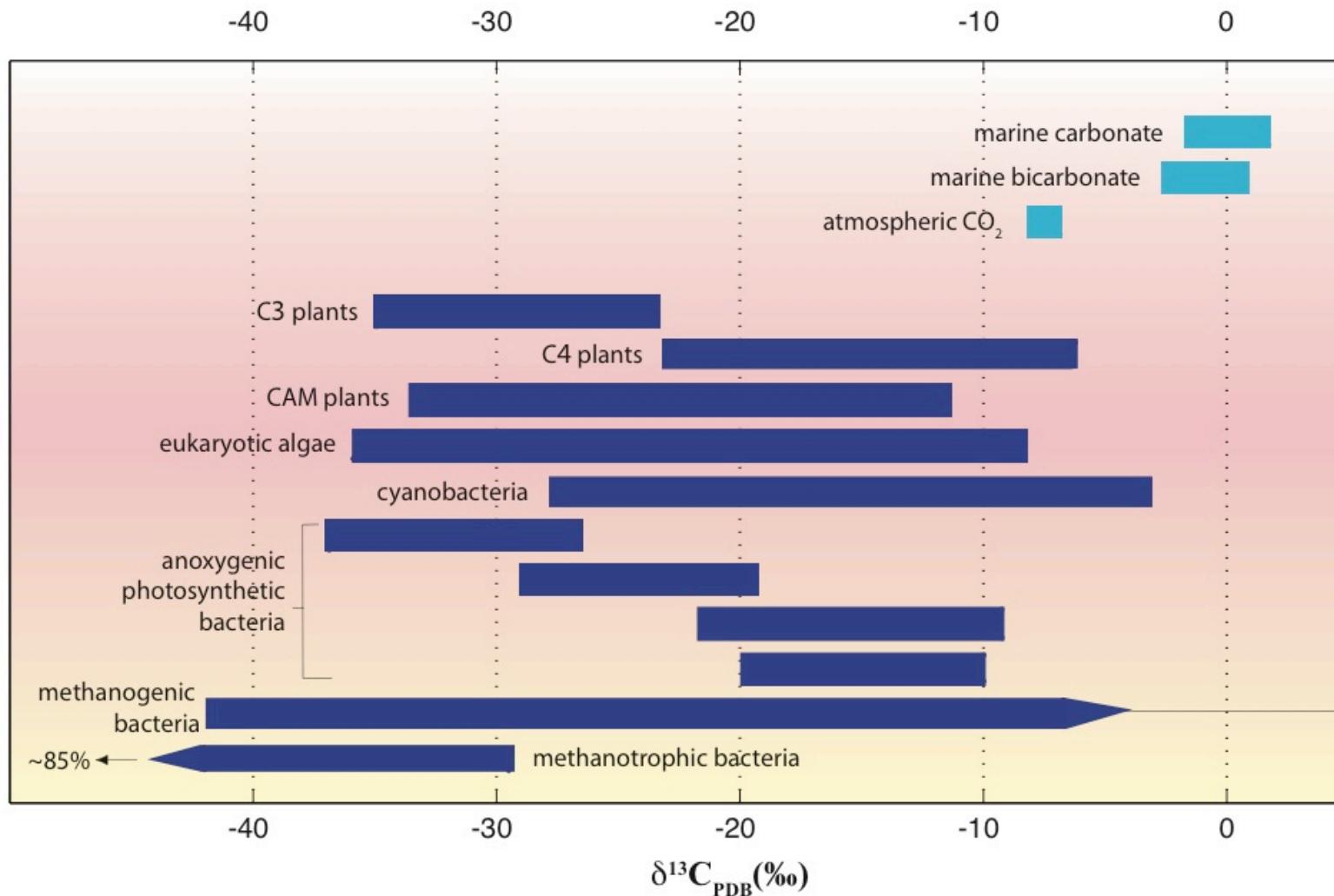
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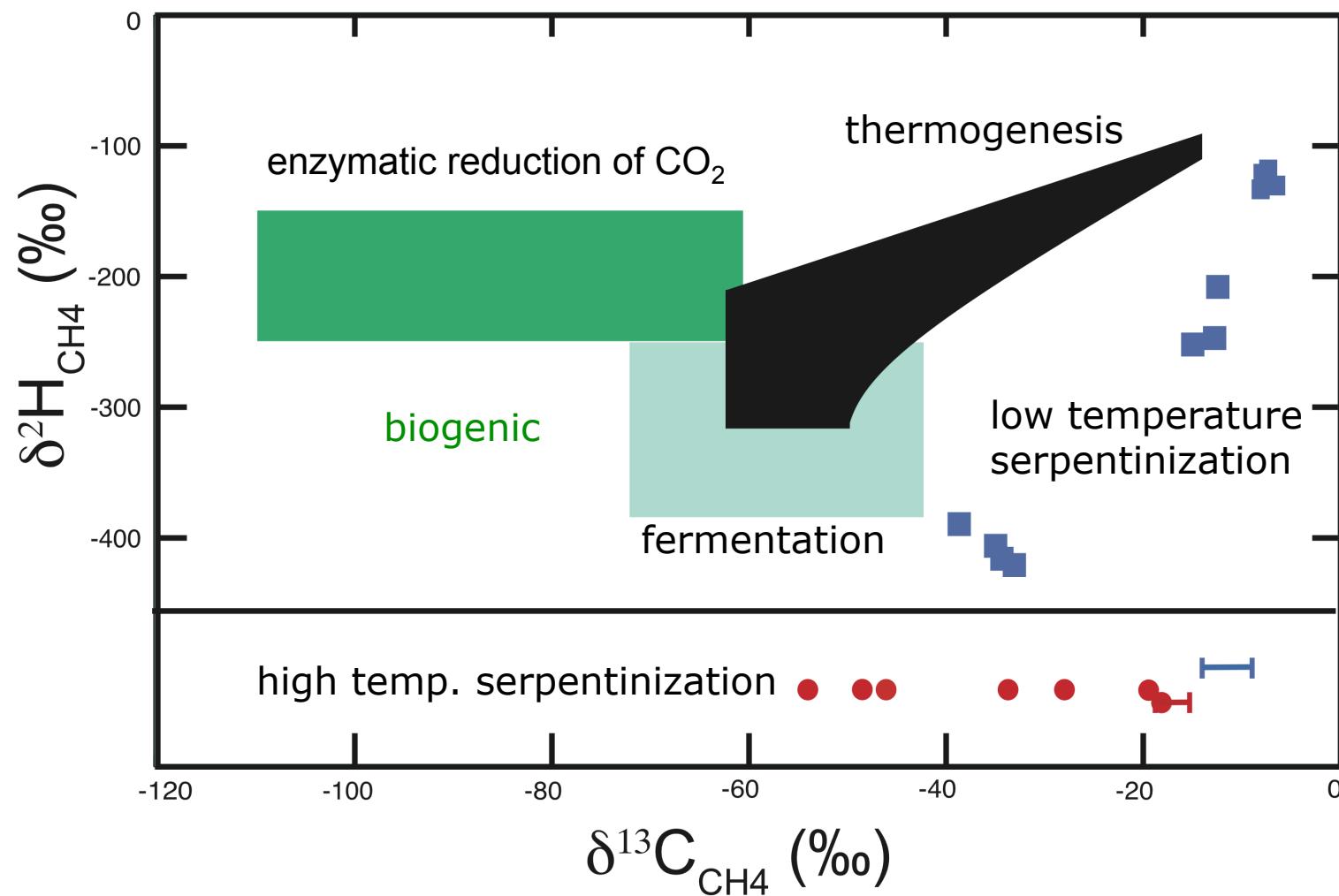
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Organic (living) vs. Inorganic ^{13}C carbon isotope

$$\text{V-PDB } ^{12}\text{C}/^{13}\text{C}=89.4$$

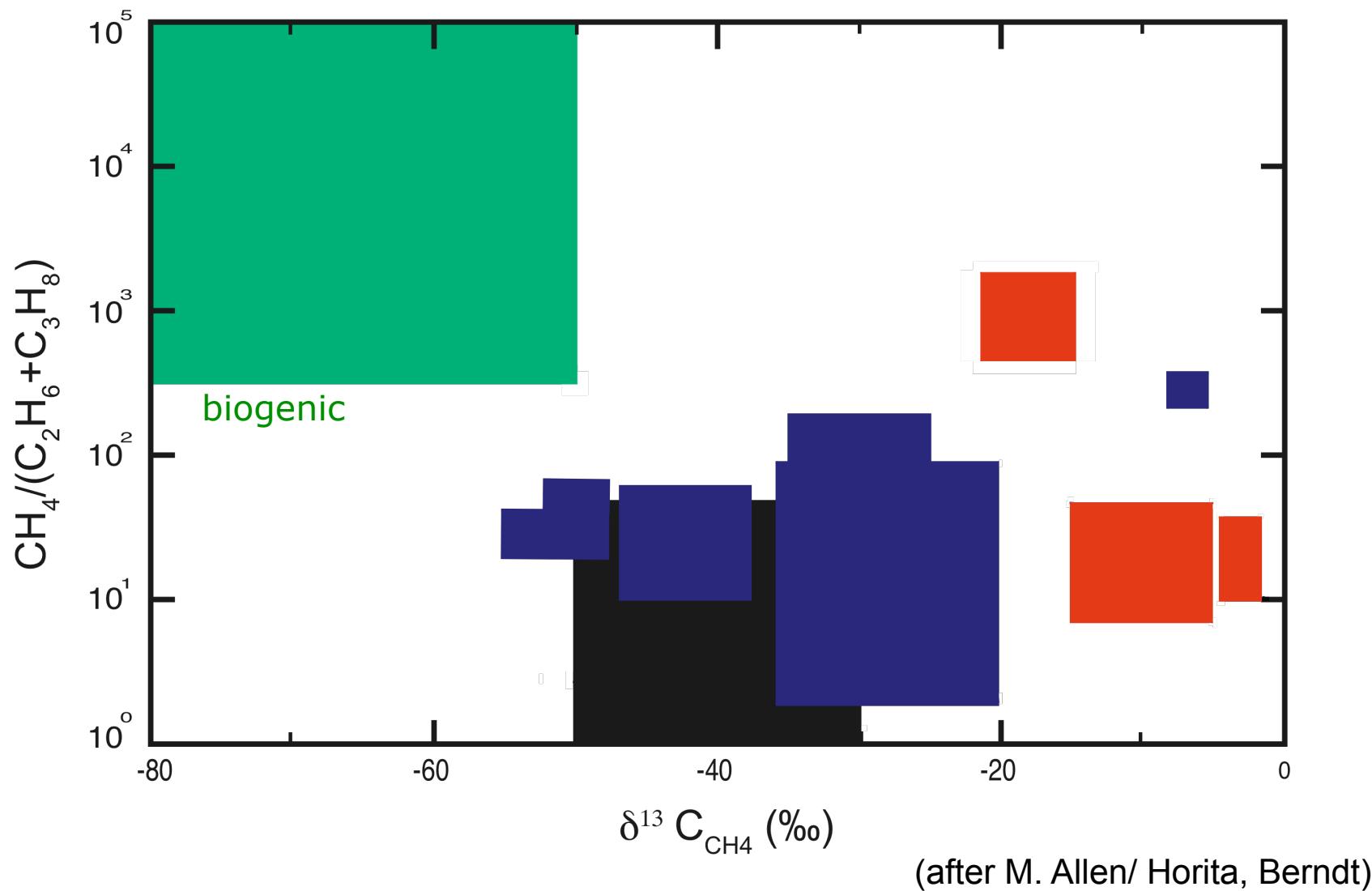


$\delta^2\text{H}$ vs $\delta^{13}\text{C}$ in CH_4



(adapted from M. Allen/ T. Onstott/ Sherwood-Lollar)

ratio of methane/(ethane + propane) with $\delta^{13}\text{C}$ (methane)



habitability: what must be measured

- **Methane** to 1 ppbv, and its *map*,
- **Localize** the vent/s to within a few tens of kilometers
- **C-isotopes** in methane
- **Reference** (inorganic) C-isotope (e.g. CO₂ in air)
- **Alkanes**: ethane, propane, butane
- **Organics**: find them, their chirality, ¹²C/¹³C in evolved CH₄
Context : geochemical, geological, environmental
- *plus more*: D/H, ¹⁸O/¹⁷O/¹⁶O, ³⁴S/³²S, trace gases, etc.

take home

- ***Methane*** is necessary, but not sufficient, for proving habitability, or extinct or extant life on Mars
- ***Geology and biology*** are equally plausible, and both require
- ***Liquid water***, to produce methane
- ***Related data*** on multiple isotopes and trace gases in solid and gas samples, mineralogical/environmental context, etc., essential to nail the origin of methane and habitability of Mars

MSL and ExoMars are only a start

?

- atreya@umich.edu
- pdf's of pubs: my website (google sushil atreya)
- Scientific American article, “Mystery of Methane on Mars and Titan”: ask or email me