



ROYAL OBSERVATORY
OF BELGIUM

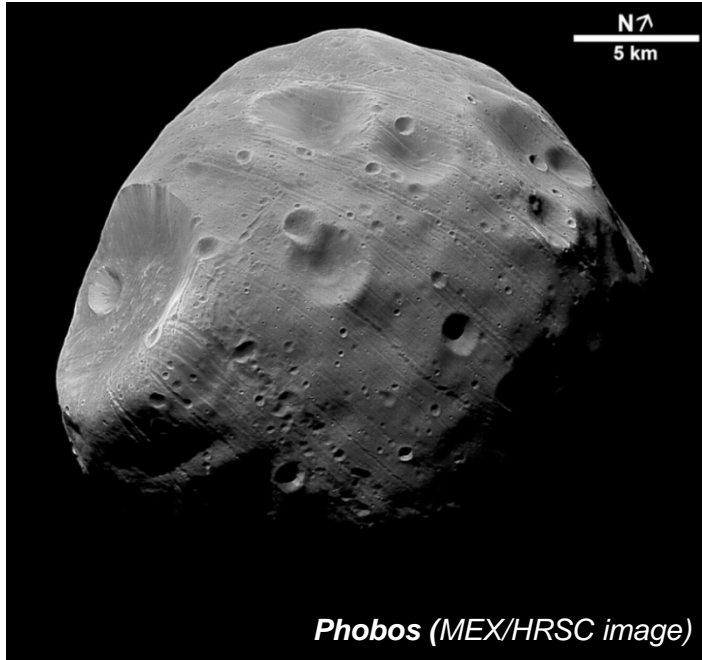


On the formation of the Martian moons from a circum-Mars accretion disk

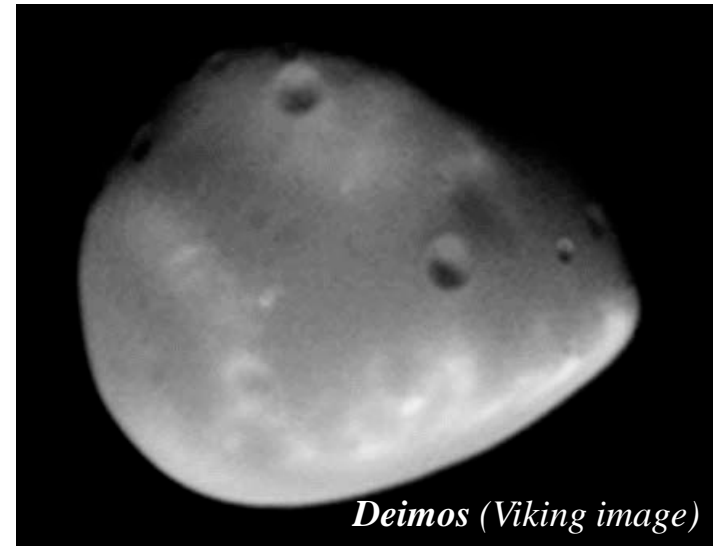
Rosenblatt P. and Charnoz S.

*46th ESLAB Symposium: Formation and evolution of moons
Session 2 – Mechanism of formation: Moons of terrestrial planets
June 26th 2012 – ESTEC, Noordwijk, the Netherlands.*

The origin of the Martian moons ?



Size: 13.0km x 11.39km x 9.07km



Size: 7.5km x 6.1km x 5.2km

- *Unlike the Moon of the Earth, the origin of Phobos & Deimos is still an open issue.*
- *Capture or in-situ ?*

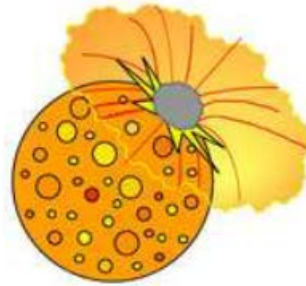
1. Arguments for in-situ formation of Phobos & Deimos

- **Near-equatorial, near-circular orbits around Mars**
(→ formation from a circum-Mars accretion disk)
- **Mars Express flybys of Phobos**
 - Surface composition: phyllosilicates (*Giuranna et al. 2011*)
=> body may have formed *in-situ* (close to Mars' orbit)
 - Low density => high porosity (*Andert, Rosenblatt et al., 2010*)
=> consistent with gravitational-aggregates

2. A scenario of in-situ formation of Phobos and Deimos from an accretion disk

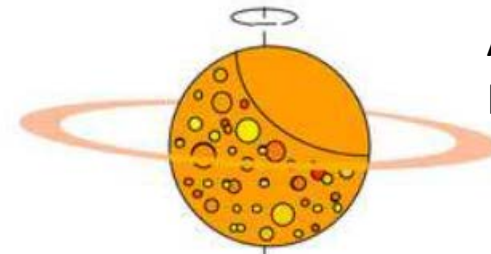
Adapted from Craddock R.A., *Icarus* (2011)

Giant impact
on Proto-Mars



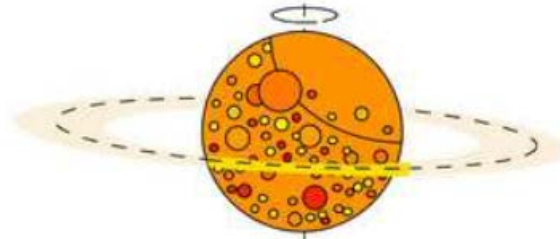
A

Accretion disk
Mass: $\sim 10^{18}$ - 10^{19} kg

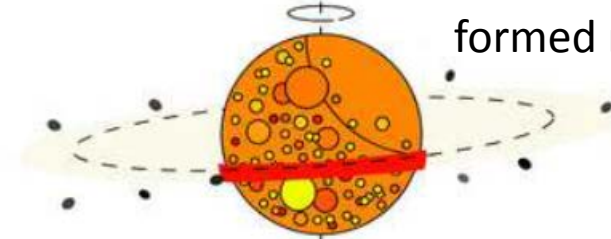


B

Gravitational instabilities
formed moonlets

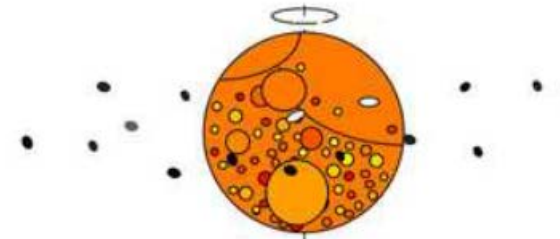


C



D

Moonlets fall
back onto Mars
→ elongated
craters



E

Two moonlets
survived 4 Gy later



F

Physical modelling never been done.
What modern theories of accretion
can tell us about that scenario?



3. Purpose of this work:

Origin of Phobos & Deimos consistent with theories of accretion?

- ✓ Modern theories of accretion : 2 main regimes of accretion.
- (a) **The strong tide regime** => close to the planet ~ Roche Limit

Background :

Accretion of Saturn's small moons (Charnoz et al., 2010)

Accretion of the Earth's moon (Canup 2004, Kokubo et al., 2000)

- (b) **The weak tide regime** => farther from the planet

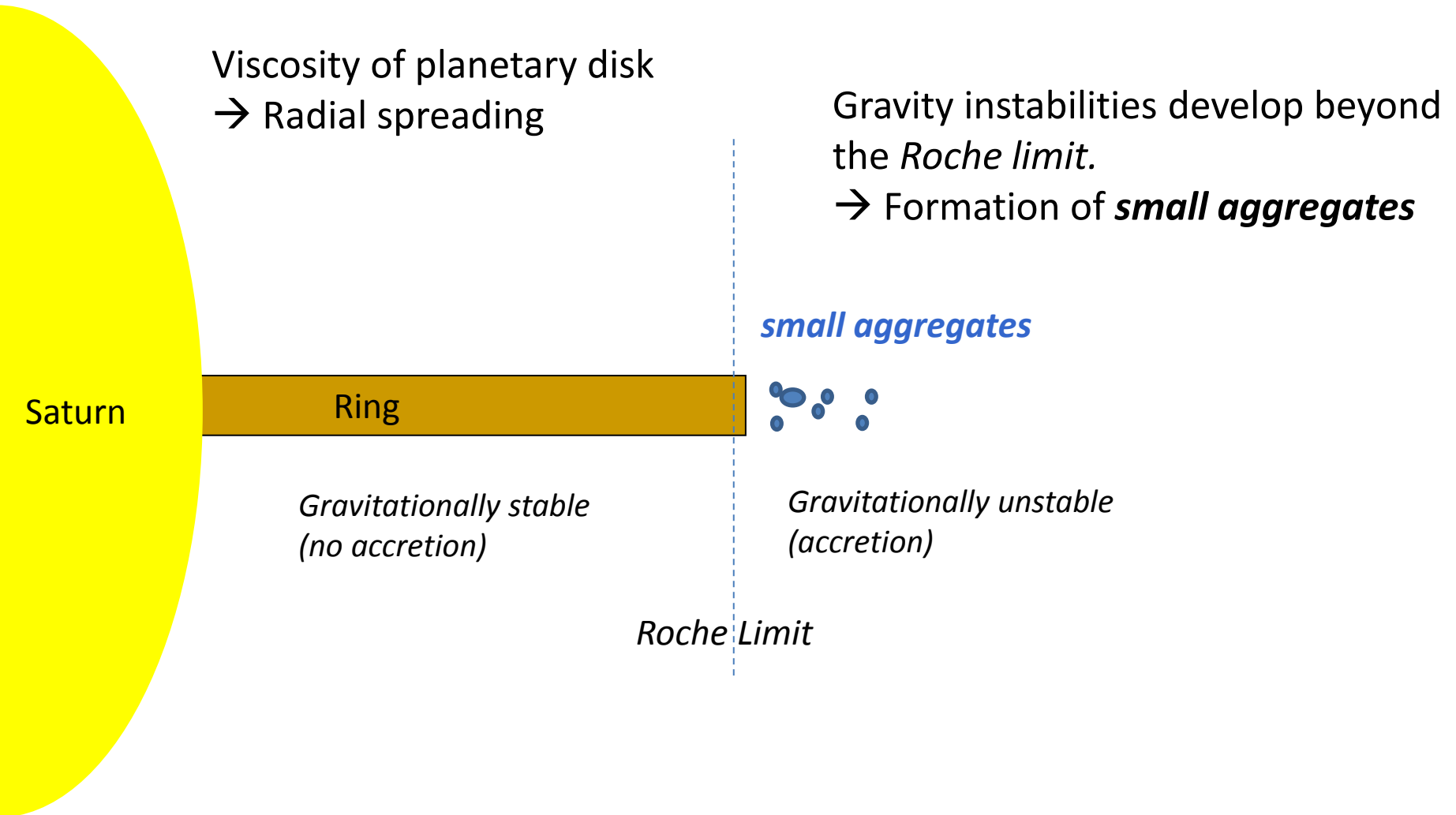
Background :

Accretion of big satellites of Jupiter & Saturn

*Accretion of planetary embryos (Lissauer 1987, Kokubo 2007,
and works from G. Wetherill, S. Weidenschilling, R. Greenberg)*

- ✓ We explore the consequences of these **regimes of accretion** for the *in-situ* formation of the Martian moons.

1. Accretion in a strong tide regime: basic physics of the disk

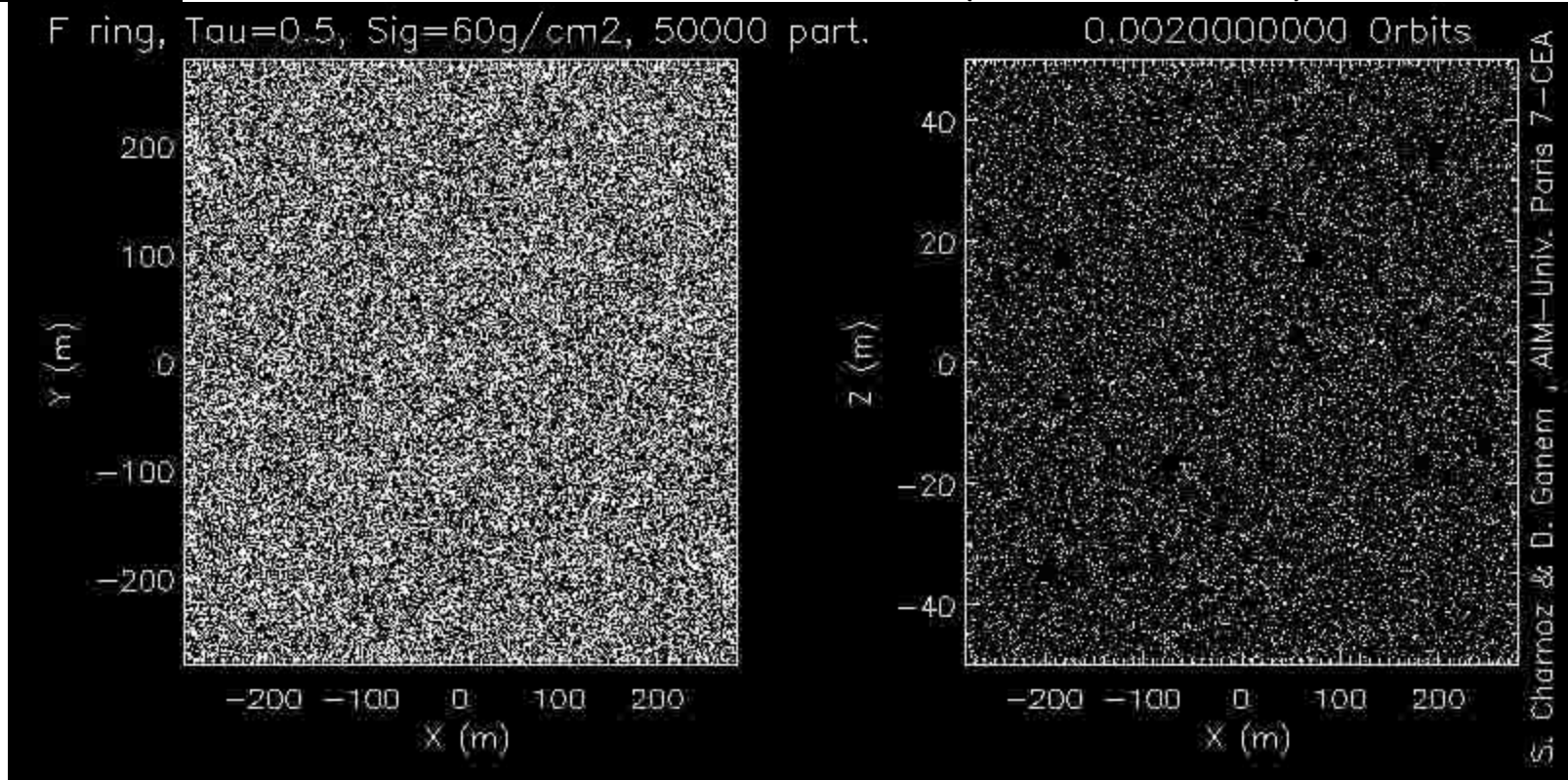


✓ This model explains the origin of Saturn's small Moons (*Charnoz et al., 2010*)

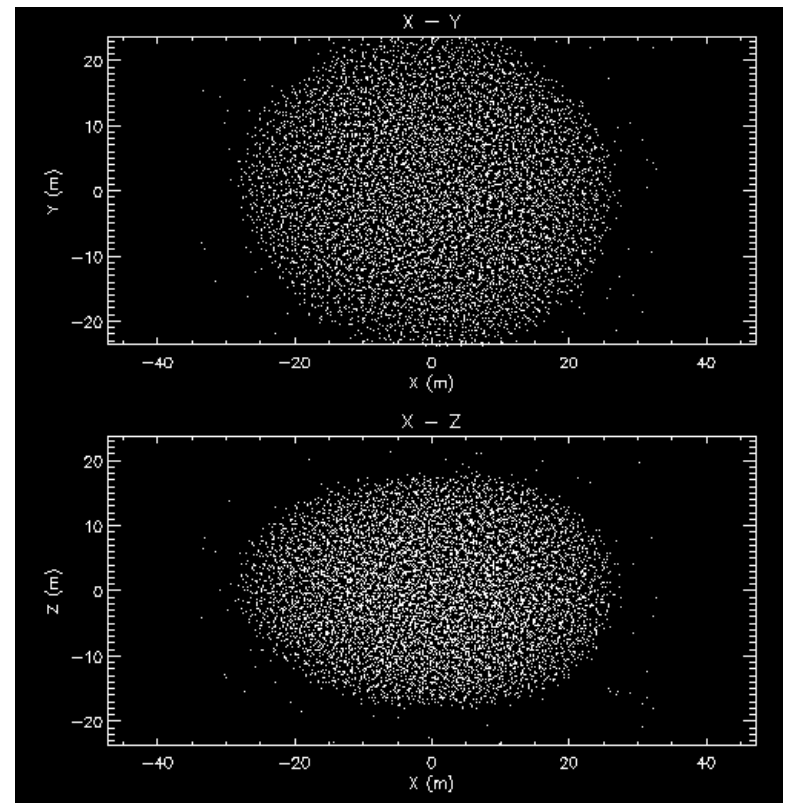
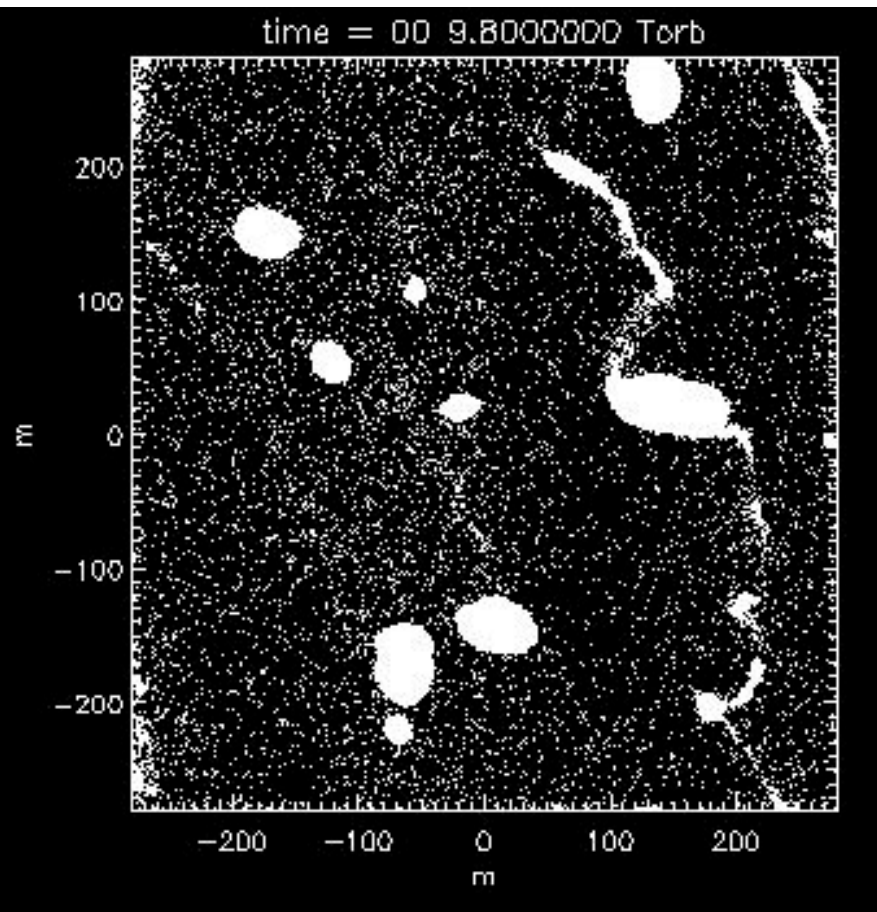
2. Accretion in a strong tide regime: Formation of gravitational instabilities

Accretion at the Roche Limit occurs on a short timescale \sim Orbital period

Example : moon formation at Saturn's Roche Limit (S. Charnoz, 2008)



3D local particle simulation (500 m x 500 m)
(Charnoz S., 2008)



Accretion of disk particule in the area of gravitational influence
(or Hill Sphere) of the instability:

→ Creation of elongated moons

With shape ~ Hill Sphere shape (Triaxial ellipsoid with axes ratio : 3:2:2)

3. Morphological properties of moons accreted at the Roche Limit

- ⇒ Growth of the gravitational aggregate is limited by its Hill Sphere
- ⇒ A gravitational-aggregate fills ENTIRELY its Hill sphere

The HILL SPHERE :

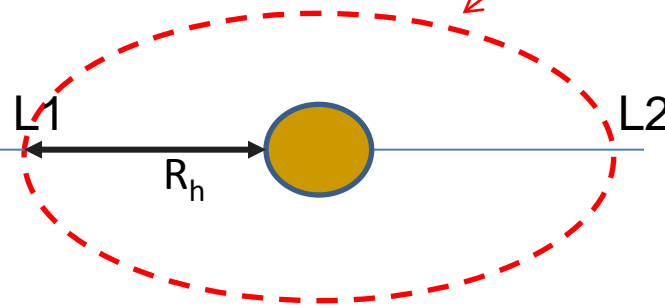
Tri-axial ellipsoid of gravitational influence

$$R_h = a (M_s / 3M_p)^{1/3} \quad \text{Volume : } \sim 1.59 R_h^3$$
$$\text{Axes ratio 3:2:2} \quad \text{Density : } \sim M_p / (1.59 a^3)$$

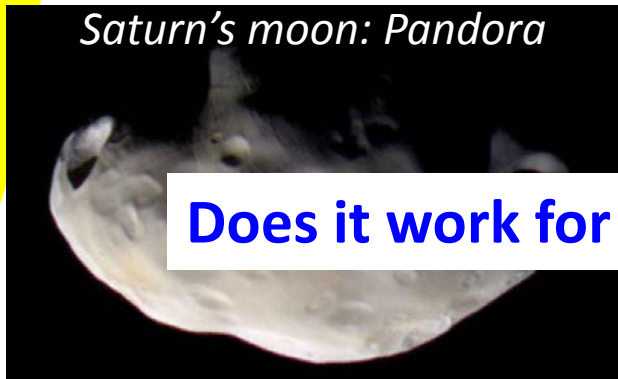
→ It works for Saturn's small moons.

« zone of accretion » of a gravitational-aggregate

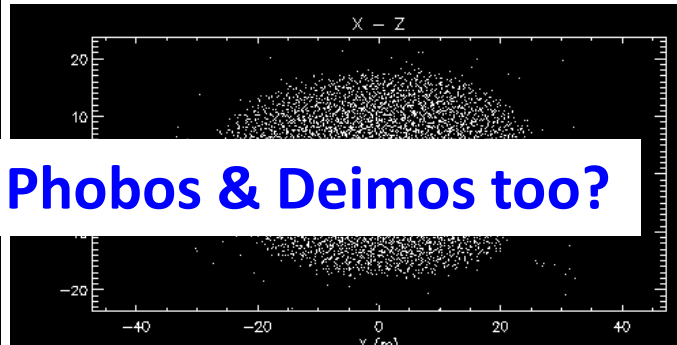
Planet



Saturn's moon: Pandora



Does it work for Phobos & Deimos too?



Simulated
Grav. Aggregate

(S. Charnoz, 2008)

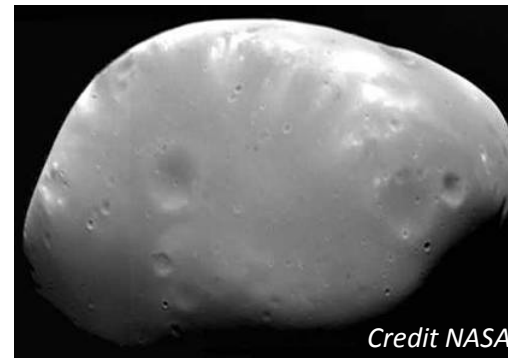
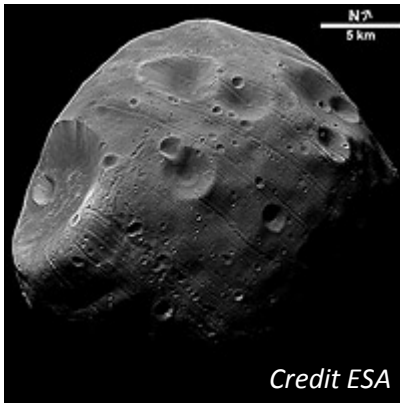
4. Morphological properties of moon accretion at the Roche Limit: Shape of gravitational aggregates

In high-tide accretion regime a gravitational-aggregate fills ENTIRELY its Hill sphere

⇒ Shape of Hill Sphere ⇔ Shape of Satellite

⇒ **COMPARISON WITH PHOBOS & DEIMOS AXES RATIO:**

Phobos	Deimos
Triaxial ellipsoid: 13.0 km x 11.39 km x 9.07 km Axes ratio: 2.86:2.5:2 (~3:2:2)	Triaxial ellipsoid: 7.5 km x 6.1 km x 5.2 km Axes ratio: 2.88:2.34:2 (~ 3:2:2)



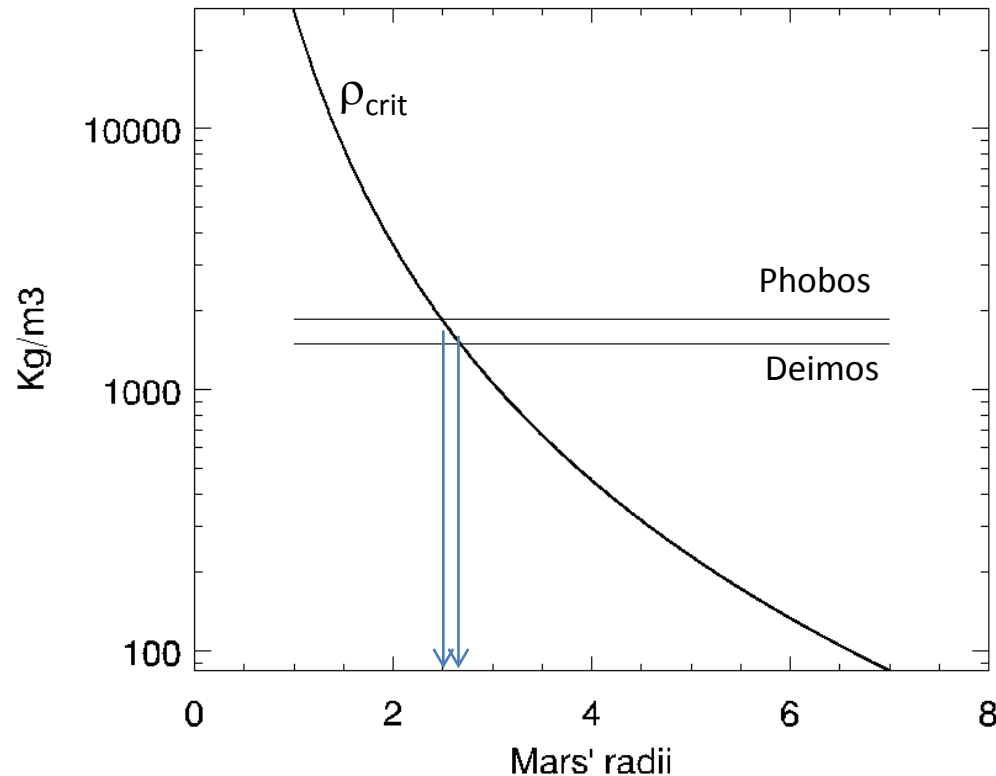
⇒ **Axes ratio 3:2:2 roughly respected**

5. Morphological properties of accretion at the Roche Limit: Density of gravitational aggregates

In a high-tide accretion regime a gravitational aggregate fills ENTIRELY its Hill sphere

=> Density of Hill Sphere \Leftrightarrow Density of Satellite

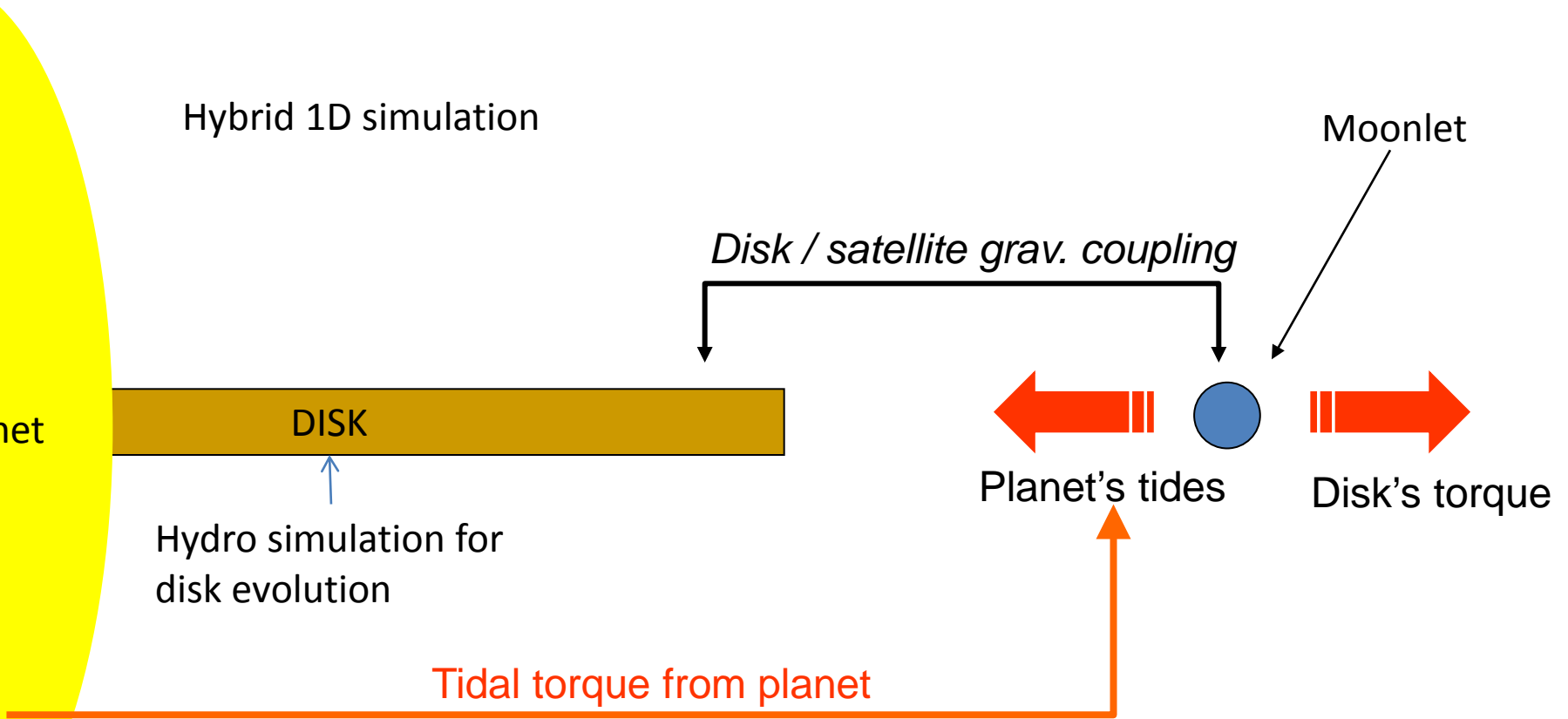
$$\rho_{crit.} \approx \frac{M_p}{1,59 a^3}$$



✓ Comparison with Phobos/Deimos density:

Density of the Martian moons are compatible with the expected critical density at Mars' Roche limit ! (=2.5 Mars' radii)

6. Putting all things together in a (simple) large scale numerical simulation:



Disk's viscous evolution:

$$\frac{\partial \sigma}{\partial t} = \frac{3}{r} \frac{\partial}{\partial r} \left[r^{1/2} \frac{\partial}{\partial r} (v \sigma r^{1/2}) - \frac{1}{3\pi(GM)^{1/2}} r^{1/2} T \right].$$

Viscous torque

Satellite torque

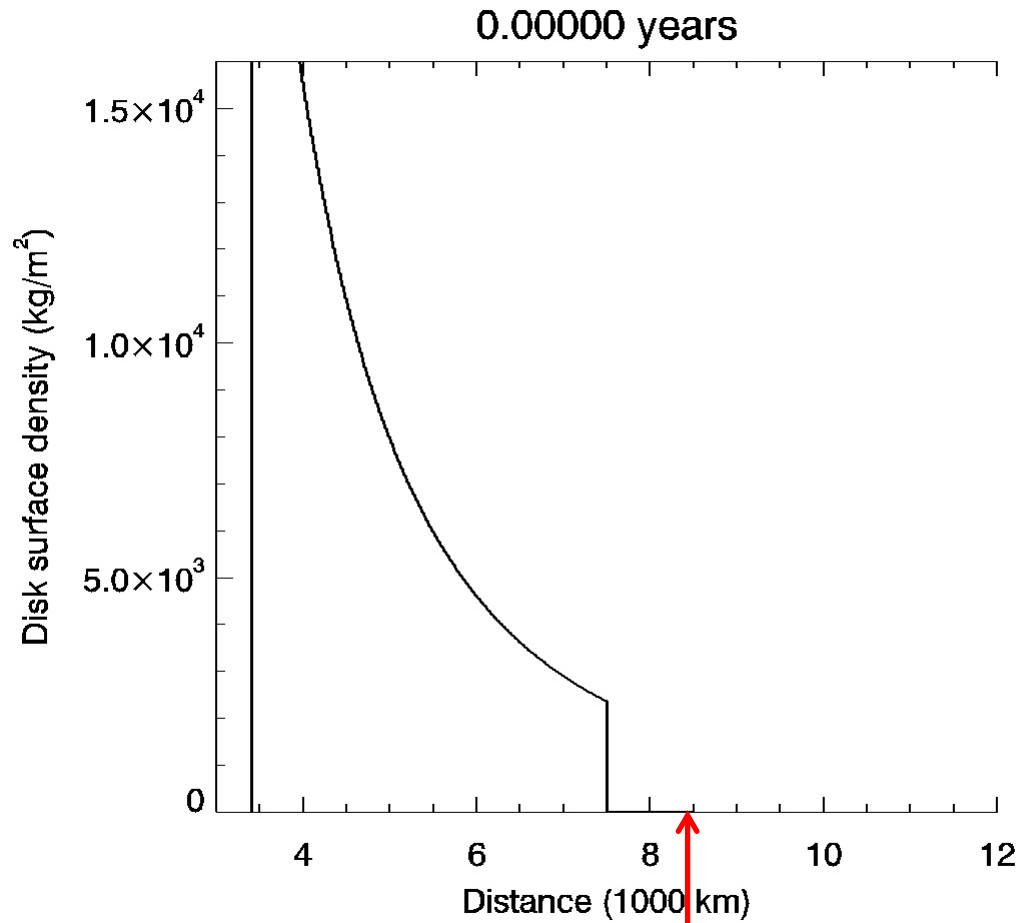
Satellite's orbital evolution
(distance to Mars - semi-major axis):

$$\frac{da_i}{dt} = \frac{3k_2^J n_i m_i (E_r)^5}{Q_J a_i^4 m_0} \left[1 + \frac{51e_i^2}{4} \right] + \frac{2a_s^{1/2} \Gamma_s}{m_s (GM)^{1/2}}$$

Tidal dissipation in Mars

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

Initial mass of the disk = 10^{18} kg (Craddock, 2011)



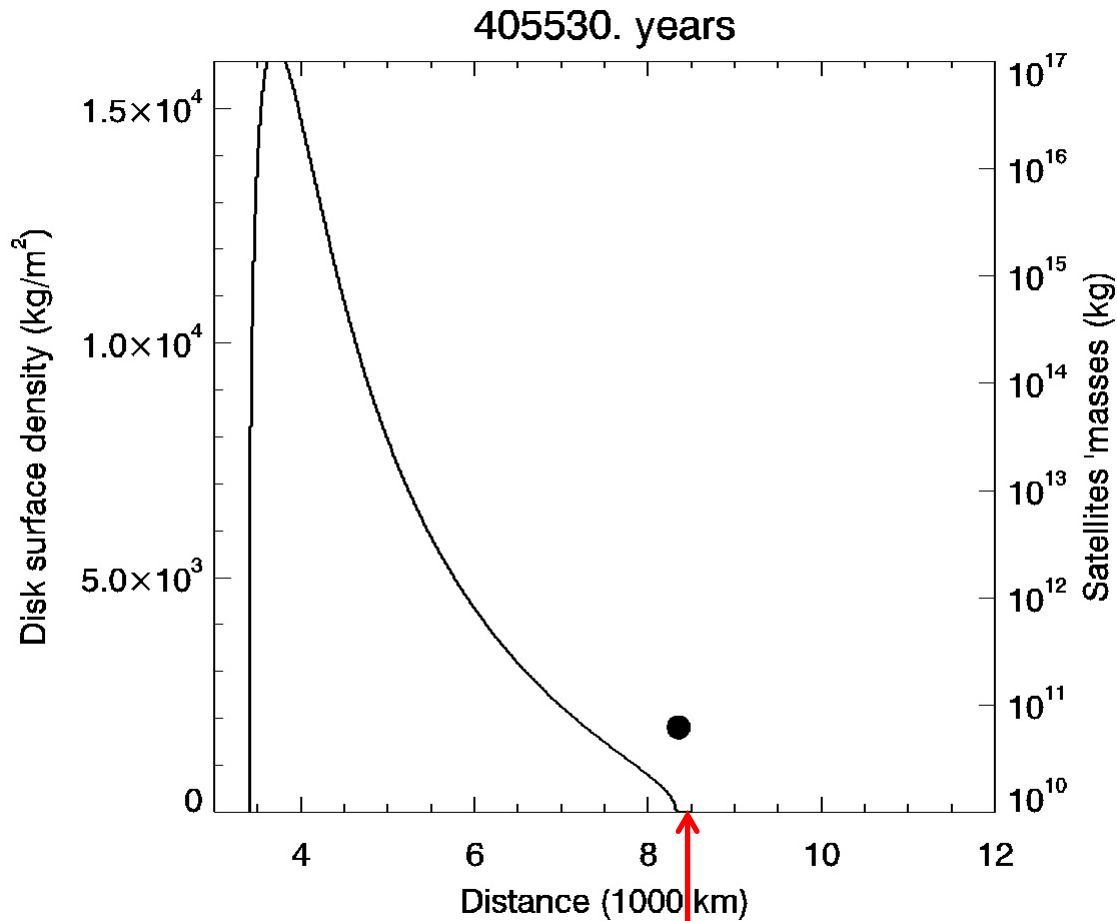
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

Viscous spreading of the disk makes it crossing the roche limit



Moonlets form at the Roche limit with a relatively low mass (10^{11} kg)



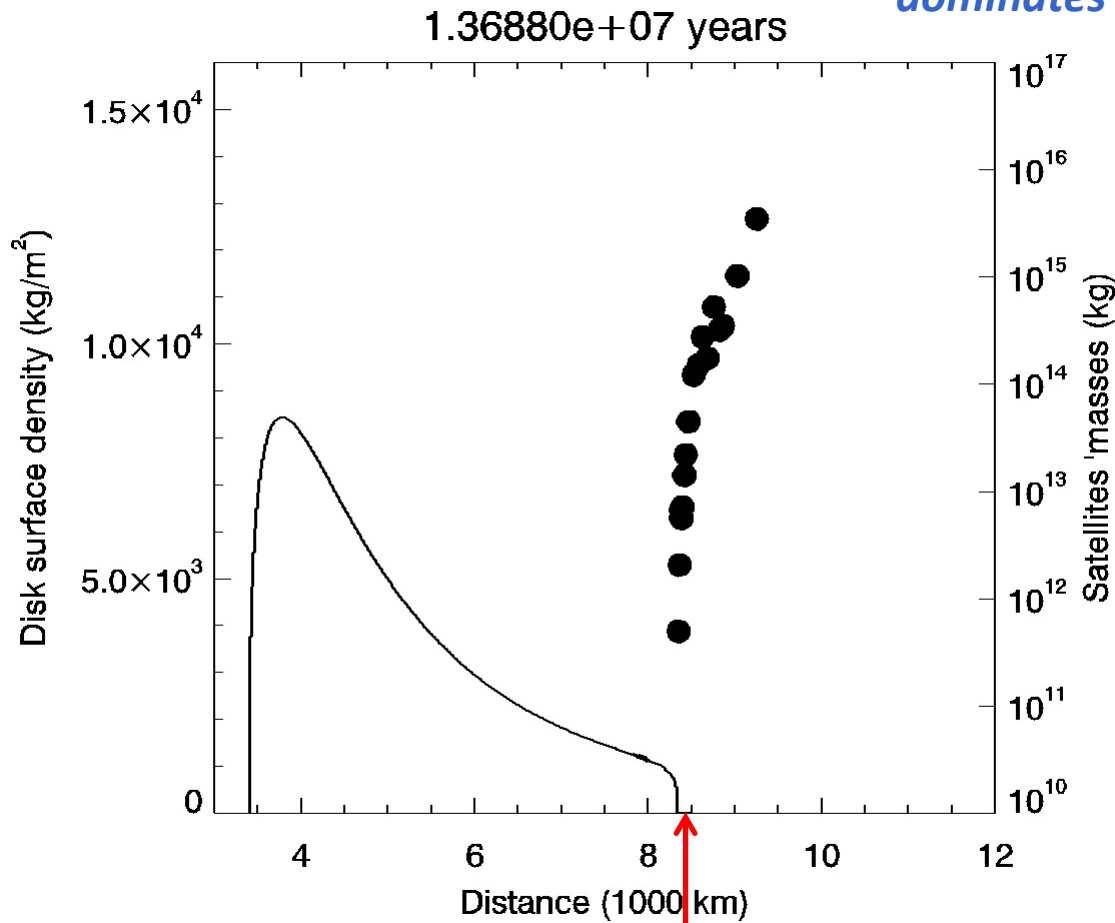
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

As viscous spreading continues several tens of moonlets are formed after about 10 Ma



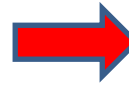
*Moonlets can reach a mass up to Deimos' mass (10^{15} kg).
Disk gravitational torque dominates the orbital evolution.*



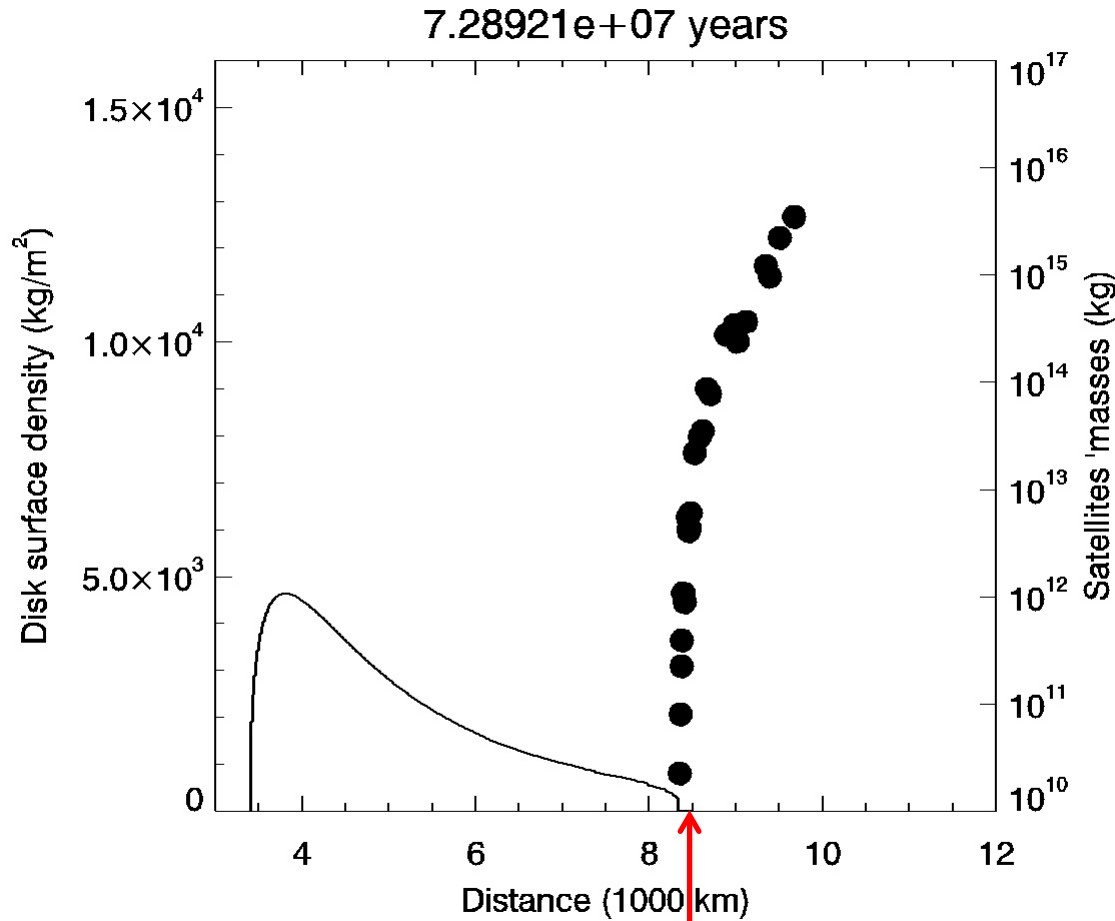
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

As the disk loses more and more mass, it is pushed below the Roche limit, stopping the accretion.



Moonlet orbits cannot reach Deimos' orbit beyond Mars synchronous limit (20000 km)



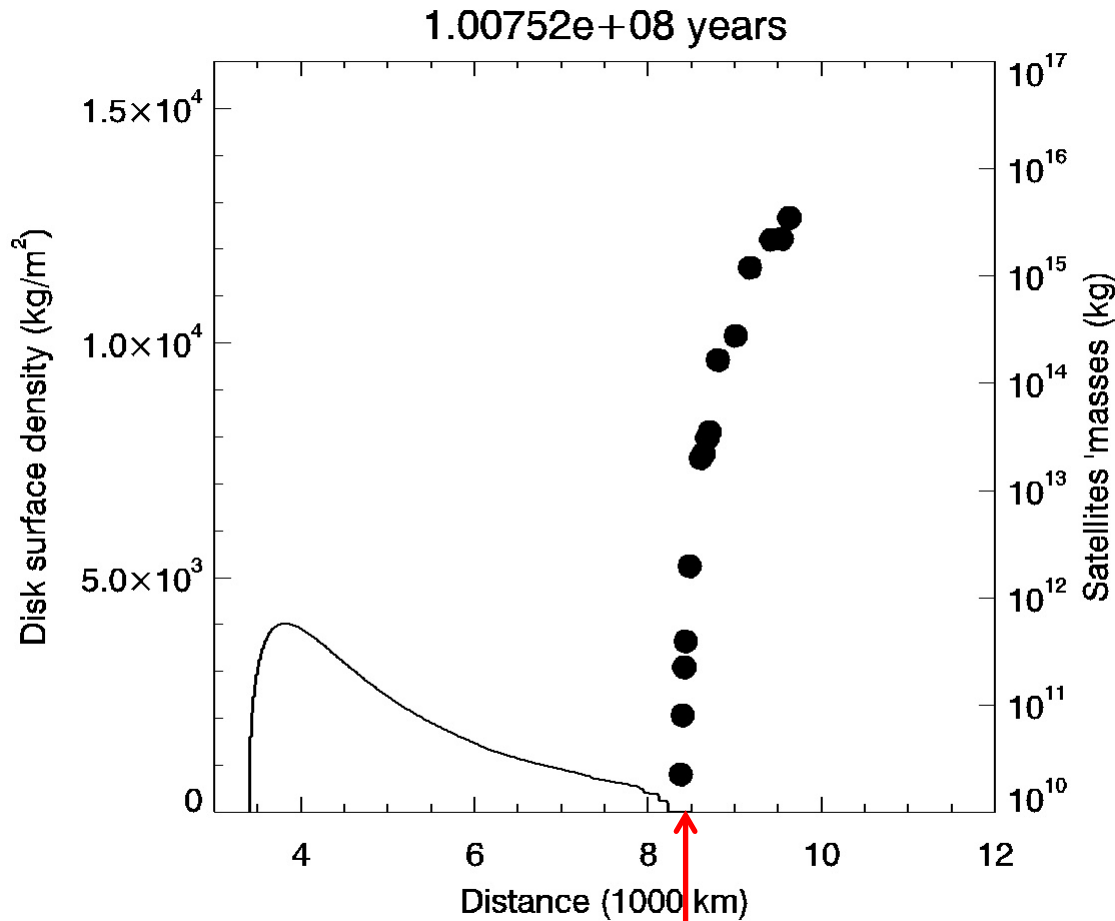
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

After 100 Ma, the disk has lost most of its mass through its inner edge.



The evolution of moonlet orbit is now dominated by tidal decay. Moonlets recede back to Mars.



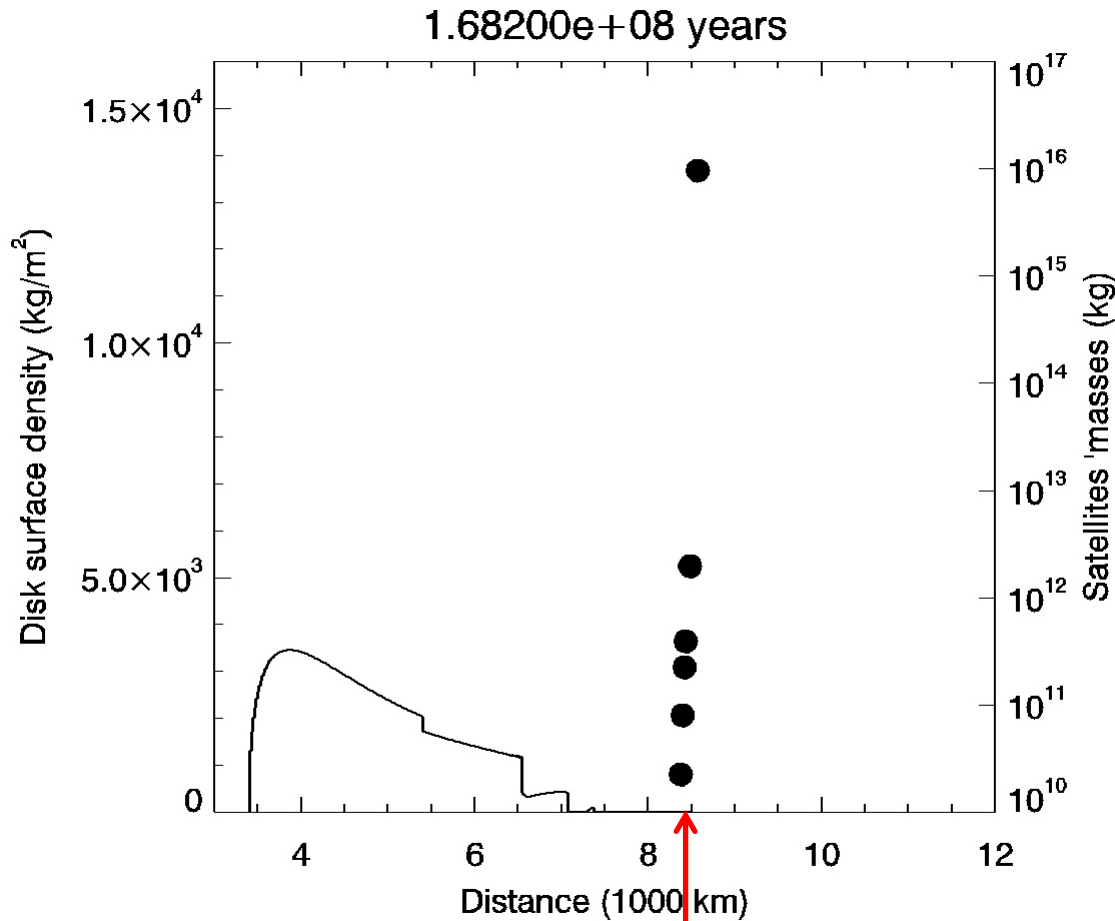
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

The disk is pushed inward by the moonlets receding back to Mars.



Moonlets accrete together, reaching the mass of Phobos (10^{16} kg)



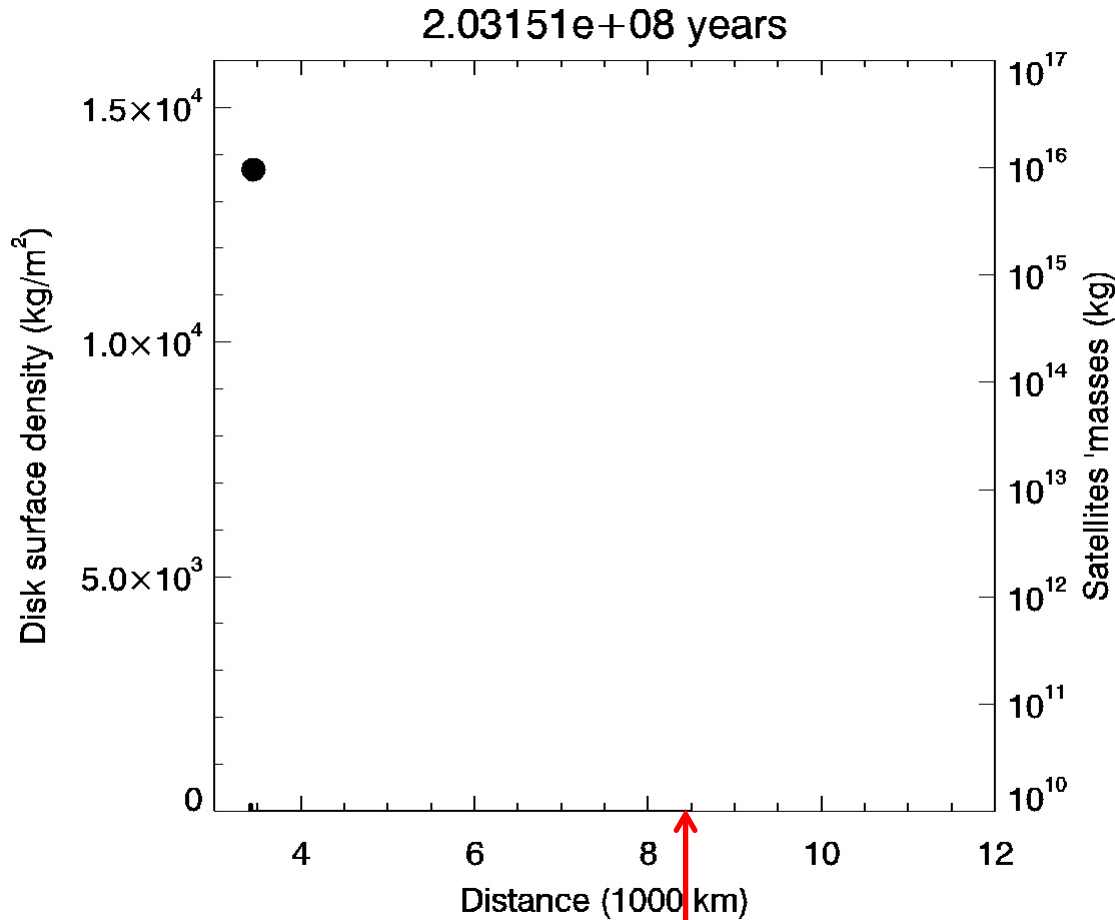
The Roche limit (~ 2.5 Mars' radii)

7. Results of numerical simulations with a disk of ~ 100 Phobos' mass:

In less than 1 Gyr, the disk disappeared and all the moonlets have fallen back to Mars.



Inconsistent with 4 Gy old Martian moon system



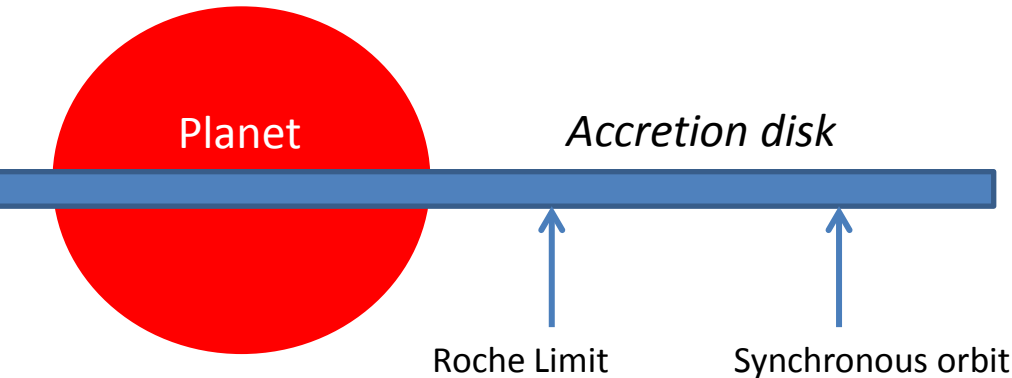
The Roche limit (~ 2.5 Mars' radii)

Main conclusion about the strong tide regime of accretion

- GOOD:
Moonlets with the shape & density of Phobos & Deimos can be formed in a Martian gravitationally unstable accretion disk (**strong-tide regime** of accretion) → **consistent with Craddock (2011)**
- BAD:
This moonlet system is inconsistent with a 4 Gy old Martian satellite system and with Deimos' location (> synchronous orbit)
→ **inconsistent with Craddock (2011)**
- A less massive disk would last longer but would form less massive moonlets and could not still form a Deimos at its current location.
- Needs to explore the **weak tide regime** of accretion (more consistent with Deimos' location, ...)

1. Weak tide regime (far from the planet): *~planet like accretion in a disk*

- ✓ The Hill sphere is much larger than the body size.
- ✓ Accretion is driven by two-bodies encounter.



✓ Hypothesis :

Disk extends to the synchronous orbit (~6 Mars' radii).

Disk NOT gravitationally unstable (→ low dense and hot)

- ✓ Accretion disk surface density profile:

$$\sigma(r) = \sigma_0 r^q$$

with $-5 < q < -0.5$ (inspired from p.p disks, simulation of moon formation)

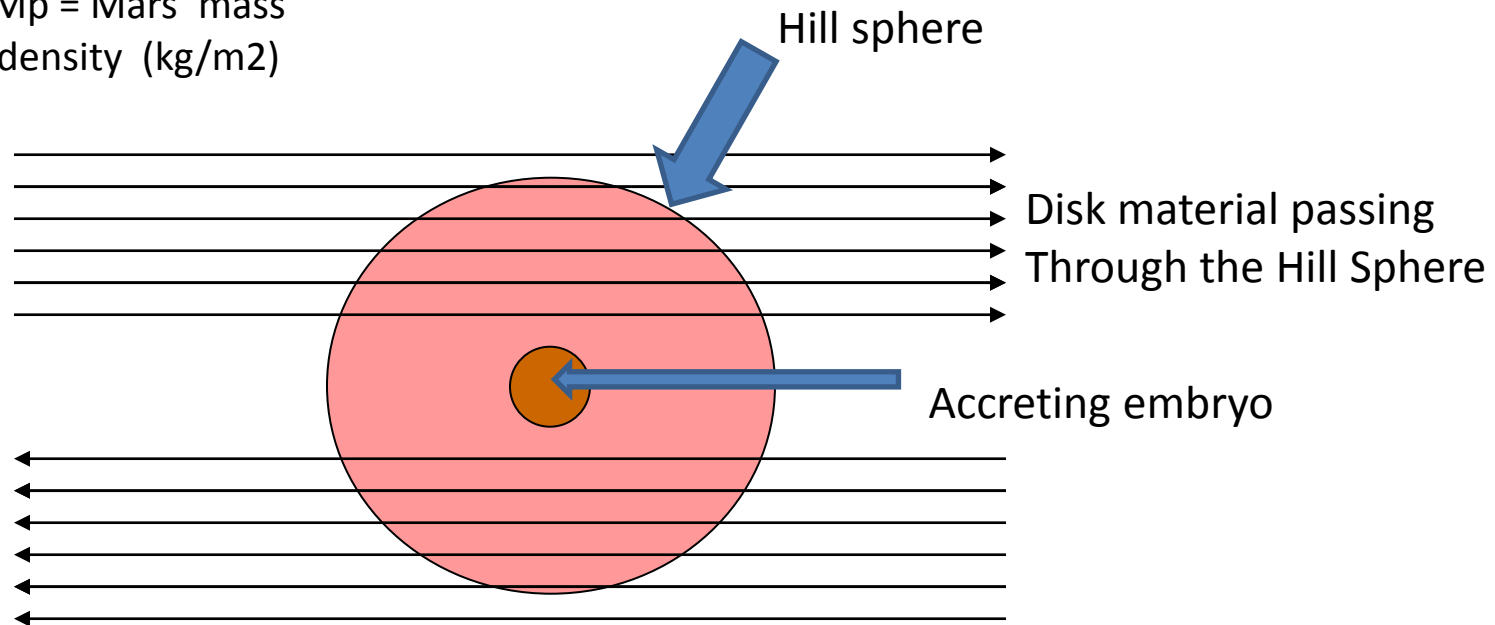
2. Weak tide regime (far from the planet): The Isolation Mass

- The mass of accreted bodies is limited by the amount of material in their feeding region (Hill Sphere).

→ Maximum growth mass = ISOLATION MASS « M_i » (Lissauer, 1987):

$$M_i = (16\pi r^2 \sigma)^{3/2} / (3 M_p)^{1/2}$$

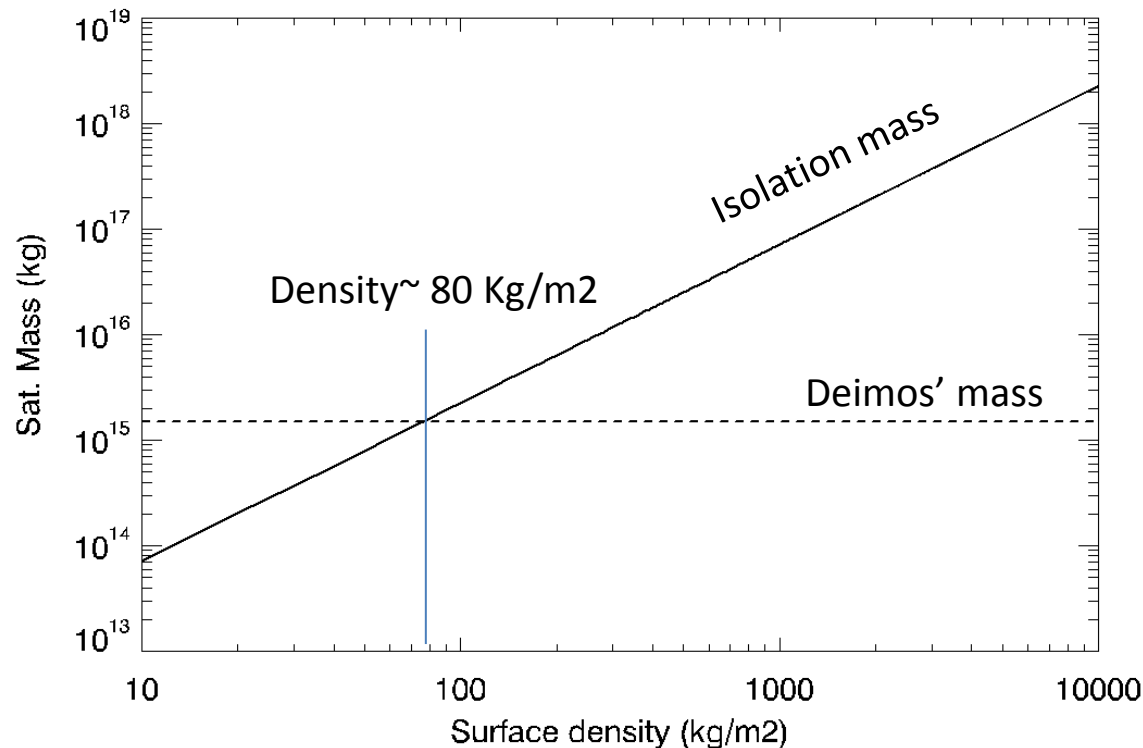
r = distance to Mars, M_p = Mars' mass
 σ = local disk surface density (kg/m²)



- If Deimos formed near its current distance as a « satellite embryo », then we can derive the surface density of the disk @ Deimos location needed to reach the isolation mass.

3. Weak tide regime : Comparison with Deimos' mass

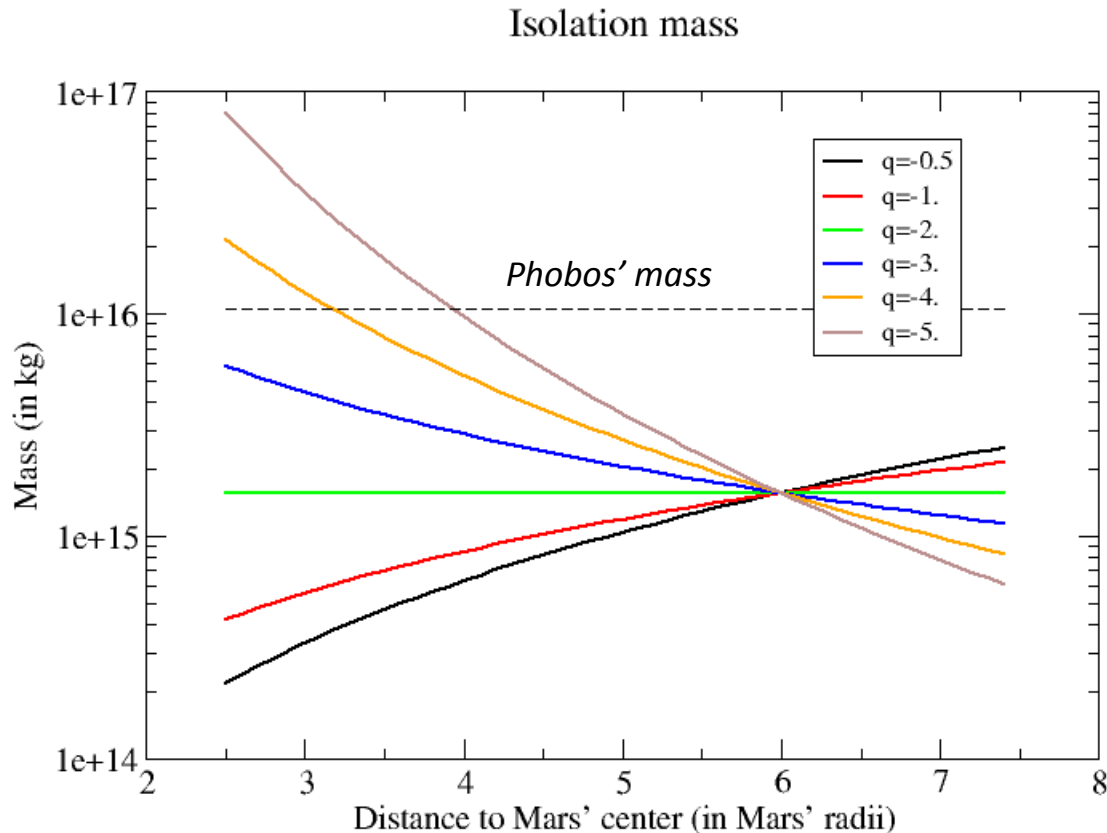
Disk surface density @ Deimos location



- ✓ Density of 80 Kg/m² at 6 Mars' radii is compatible with a disk's mass of:
 - 1.2 10¹⁸ Kg (with q=-3)
 - 3.7 10¹⁸ Kg (with q=-4)
- ✓ Or with a disk of 10¹⁸ Kg extending from 1 Mars Radius to:
 - 70 Mars radii (with q=-1.5)
 - 40 Mars radii (with q=-1)

- Deimos' mass is compatible with the expected accretion disk's surface density at Deimos orbit for a disk's mass of $\sim 10^{18}$ kg (Craddock, 2011).
- Deimos may be a satellite embryo too.
- However, the disk is required to extend beyond the synchronous orbit.

3. Weak tide regime : Comparison with Phobos' mass



- ✓ Isolation mass profile derived from disk's surface density profile (different q values)
- ✓ Each profile provides a mass equal to Deimos' mass at 6 Mars' radii

- Isolation mass can reach Phobos' mass
→ A Phobos' mass embryo can also be formed in the same disk as for Deimos.
- This embryo is formed closer to Mars (< 4 Mars' radii). Thus, expected to rapidly fall back to Mars due to tidal decay of its orbit.
- It requires that a Phobos' mass embryo be formed closer to synchronous orbit.

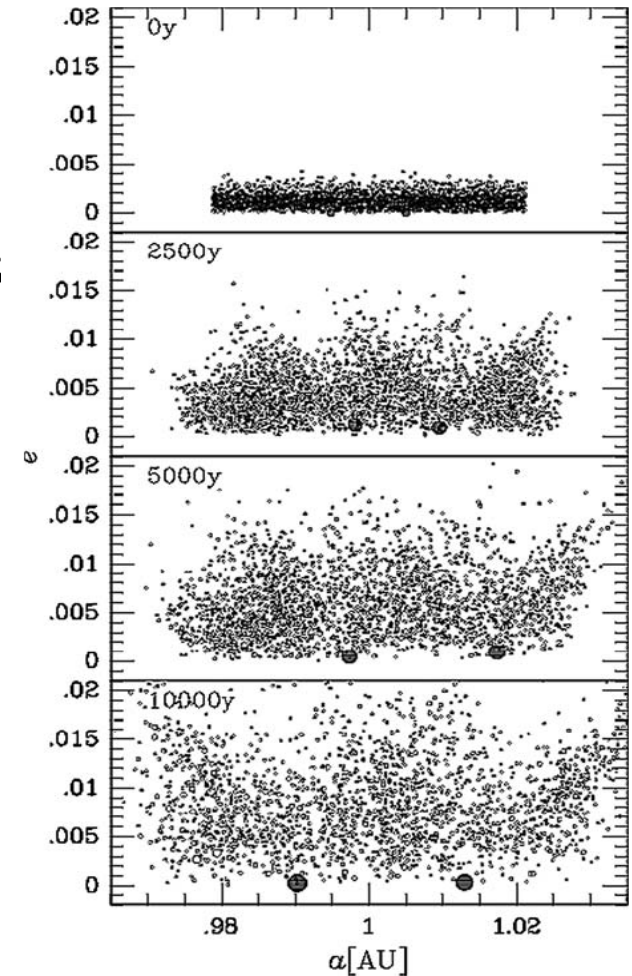
SUMMARY

- Phobos/Deimos formation in an accretion disk seems possible.
- **In a strong tide regime:**
Consistent with shape/density of the moons
Inconsistent with a 4 Gy old Martian satellite system
with Deimos' location (> Synchronous orbit)
- **In a weak tide regime:**
Consistent with Deimos' mass and location.
Problems raised:
 - Phobos' mass body at its current location?
 - It may require specific properties of the disk
(spreading farther from the planet , low density, ...)
- **Further investigations**
Numerical simulations of accretion in a disk extending primarily below the synchronous orbit.

BACKUP SLIDES

4. Low-tide regime (far from the planet): Number of moons formed

- ✓ **After the formation of embryos**
 - Embryos tend to appear on orbits separated by 5-10 x the Hill Sphere radius
 - A population of ~ 50 -100 embryos may appear between \sim Mars' Roche limit and Deimos' orbit
 - Most of them may fall onto the planet (orbital tidal decay)
→ Consistent with Craddock (2011)



Thommes & Duncan 2007

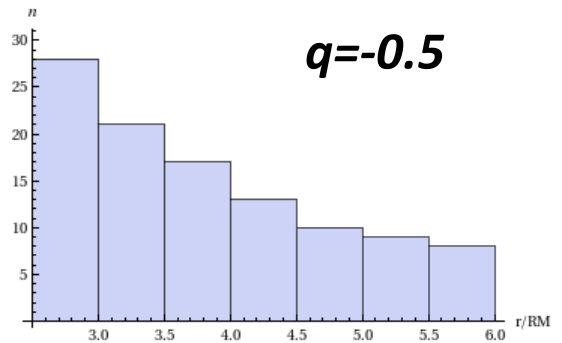
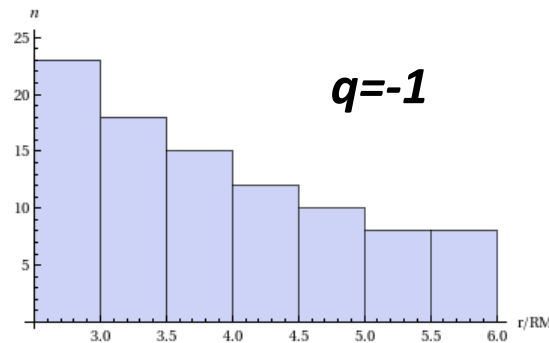
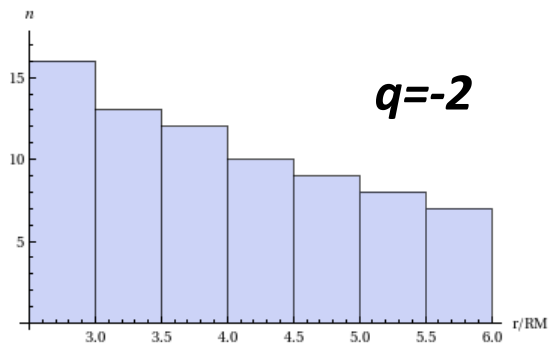
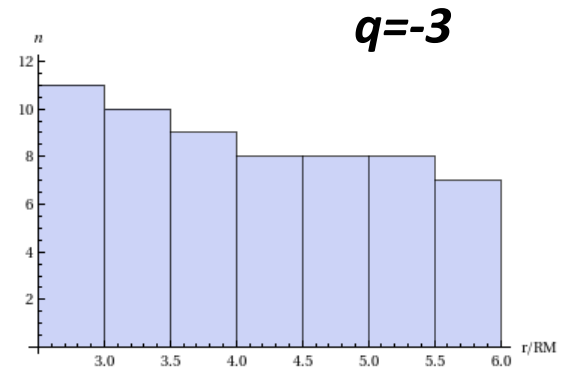
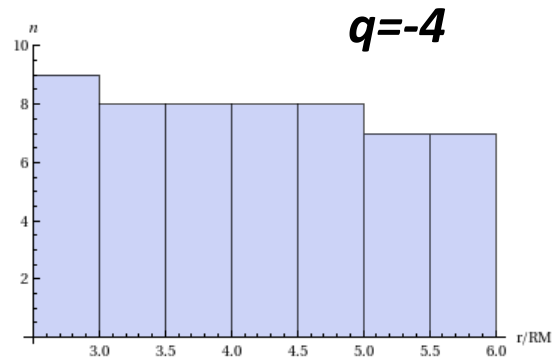
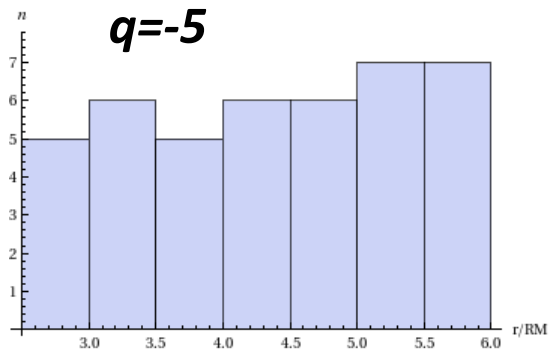
✓ Questions:

Is Phobos the last representent of the falling population?

Is Deimos the single representent of the population $> R_{\text{synchronous}}$?

3. Weak tide regime :

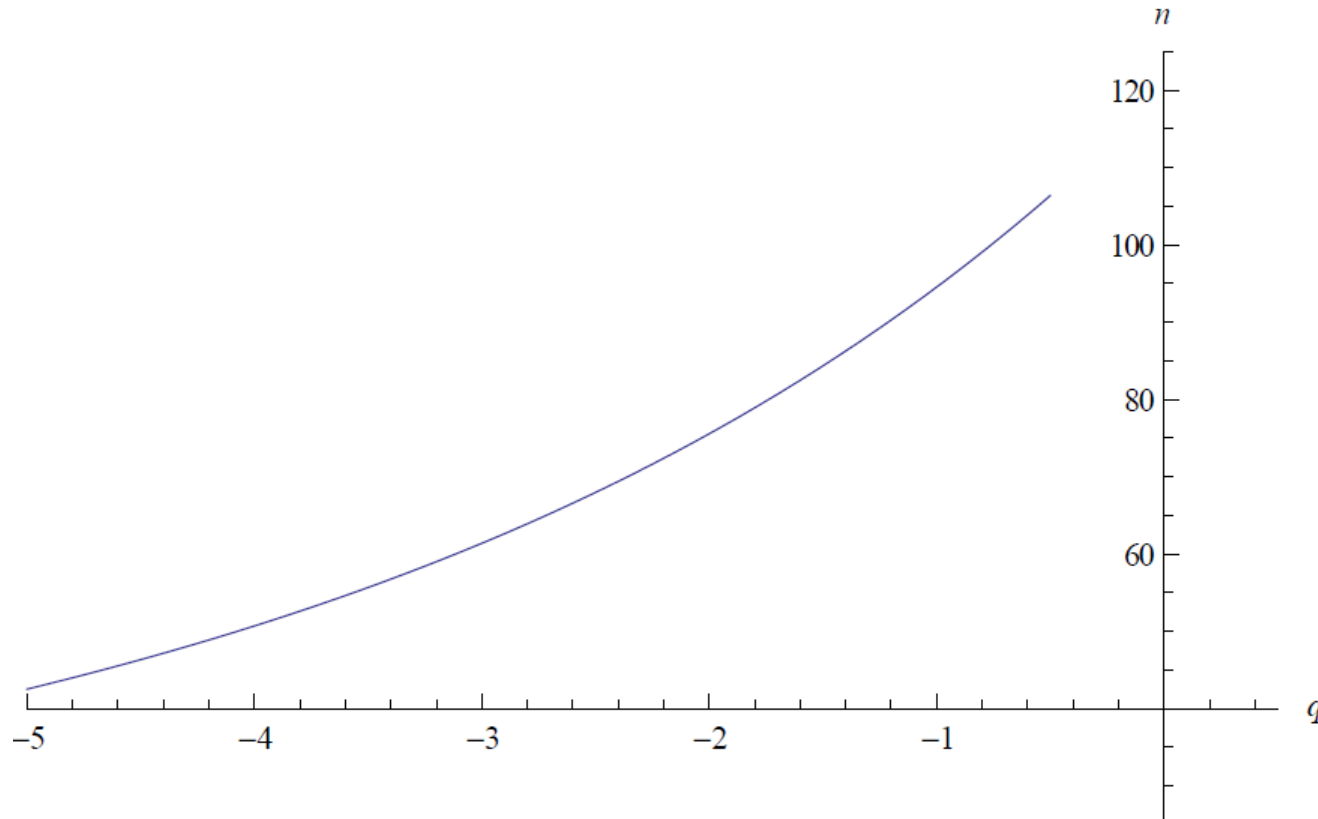
Number of satellite embryos vs distance to Mars



- The number of satellite embryo expected at distances close to the synchronous limit (between 5.5 and 6 Mars' radii) is around **7** !
- As the mass of these embryos is close to Deimos' mass, and as Deimos' mass is **7** times Phobos' mass, then :
 - ➔ Can Phobos be formed by accretion of embryos near the synchronous limit?

3. Weak tide regime :

Number of satellite embryos vs Mars' elongated craters



- The number of satellite embryo is ranging from 50 to 110 depending on the disk' surface density profile (q value).
- This number is just lower than the estimated elongated craters (102 to 174), which could have been formed from decaying moonlets.

Acknowledgements

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