

# MARS ATMOSPHERIC WATER RESEVOIRS

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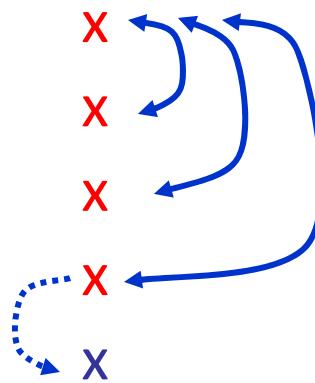
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# Introduction

## Where is the water ?

- Atmospheric vapor
- Atmospheric clouds
- Surface frosts
- Surface ices
- Subsurface ice



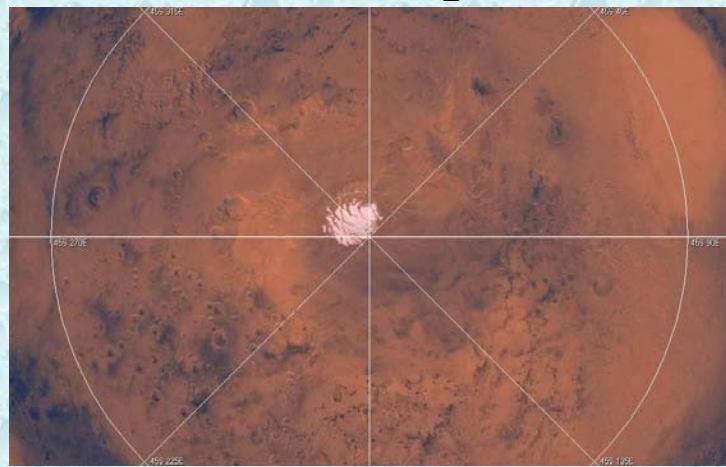
Spring sublimation is accompanied by percolation, which, on timescales of millions years can build icy-rich subsurface layers.

Global atmospheric water vapour amount is small: 10-20 pr  $\mu\text{m}$   $\rightarrow$  1-2  $\text{km}^3$   
But it is directly connected to most other water reservoirs,  
and it can be dependably measured and continuously monitored

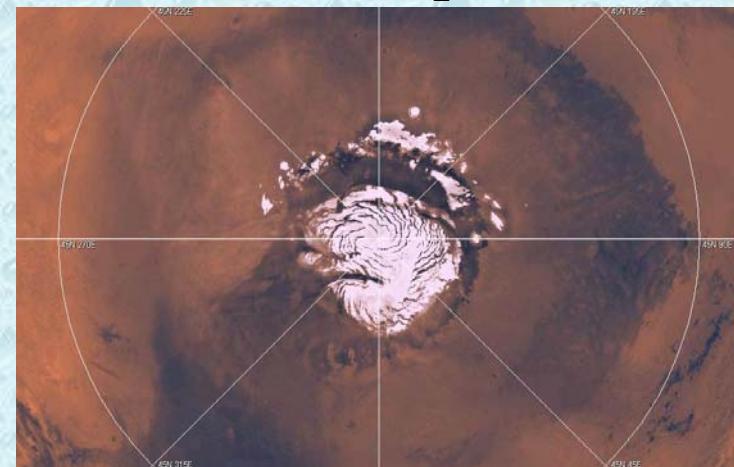
# Introduction: Global asymmetry of Mars climate

## Perennial polar caps

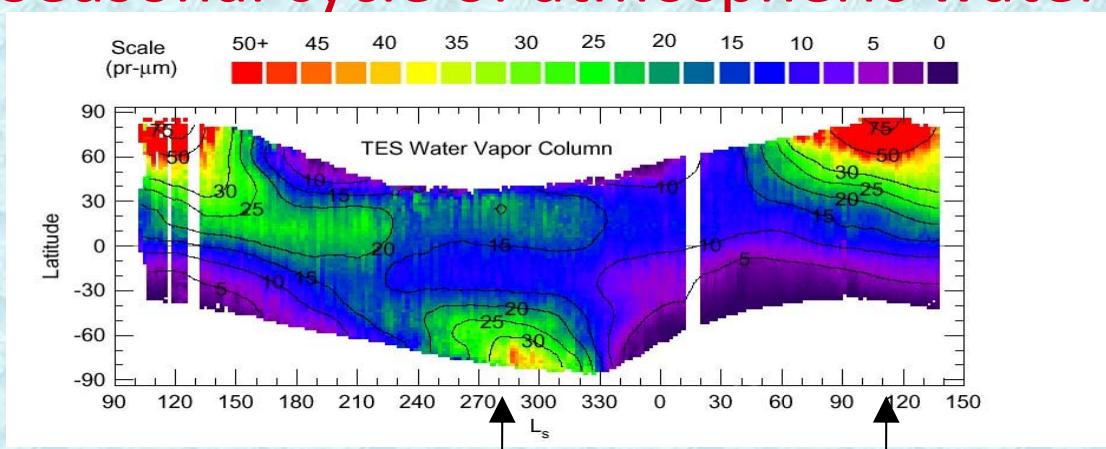
South ( $\text{CO}_2$ )



North ( $\text{H}_2\text{O}$ )



## Seasonal cycle of atmospheric water



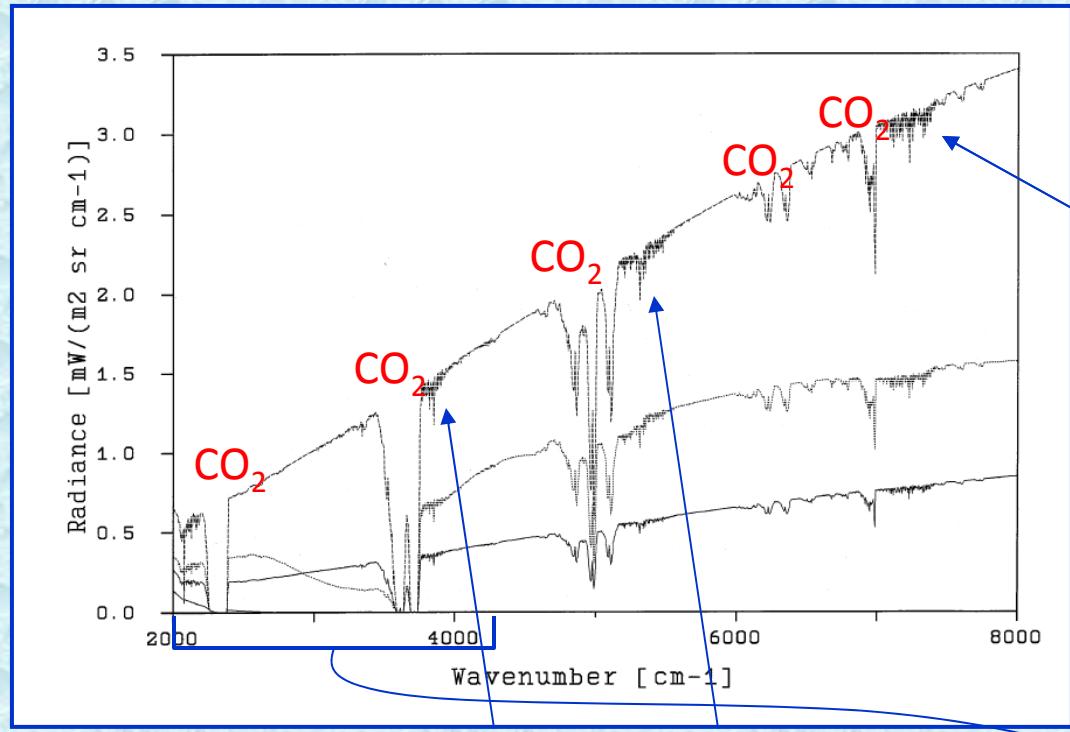
Southern summer (perihelion)

Northern summer (aphelion)

# Spectroscopy is the only mean to detect H<sub>2</sub>O remotely

Band	Comment	When, how measured	References
820-870 nm	abs; weak	Detection, Ground based	Spinrad et al 1963 Barker 1970, 76; Rizk et al 1991; Sprague et al 1996-2005
936 nm	abs; weak	Mars Pathfinder camera	Titov et al 1999
1.1 μm	abs		
1.38 μm	abs	IV/Mars 5; MAWD/Viking; SPICAM/MEX	Moroz, Nadip 1975 Jakosky, Farmer 1982, Fedorova et al., 2004, 2006, 2010
1.9 μm	abs	Auguste/Phobos 88	Krasnopolksy et al 1991 Rodin et al 1997
2.56 μm	abs	OMEGA/MEX; CRISM/MRO PFS/MEX	Melchiorri et al 2007 Maltagliati et al 2008 Smith et al 2008, 2009 Tschimmel et al 2008, Sindoni, Formisano 2010
6.3 μm	abs/emis 3.7; 8μm HDO	IRIS/Mariner 9 IRTF, TEXES ground based	Ignatiev et al 2002 Novak et al 2002 Encrenaz et al 2005, 2008
30-50 μm	emission	IRIS/Mariner 9; TES/MGS; PFS/MEX	Hanel et al 1972; Smith, 2002, 2004, 2008 Fouchet et al 2007
557 GHz	HDO, etc. iso	MIRO/Rosetta fly-by	
183 GHz	203 H <sub>2</sub> <sup>18</sup> O HDO	226 IRAM ground based	Encrenaz et al 1995, 2001
22 GHz	weak	VLA ground based	Clancy et al. 1992

# Spectral bands and spectra: near infrared



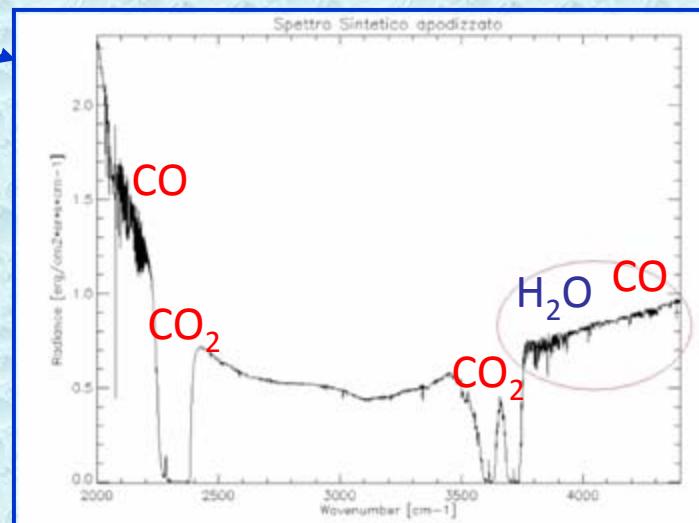
2.56- $\mu\text{m}$  band  
PFS, OMEGA, CRISM

1.87- $\mu\text{m}$  band  
Auguste/Phobos 88

Measured PFS spectrum

Haus&Titov 1999, spectra  
computed to simulate PFS signal

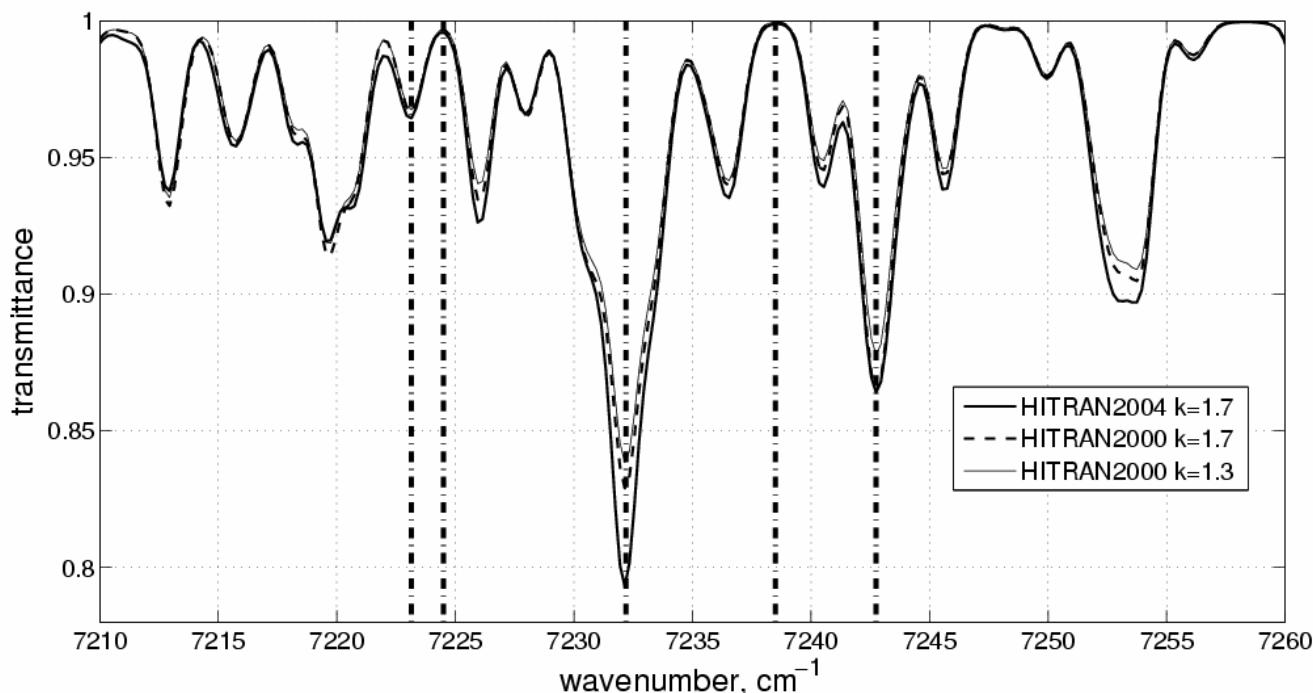
1.38- $\mu\text{m}$  band  
MAWD, SPICAM



# Near-infrared example: MAWD/Viking

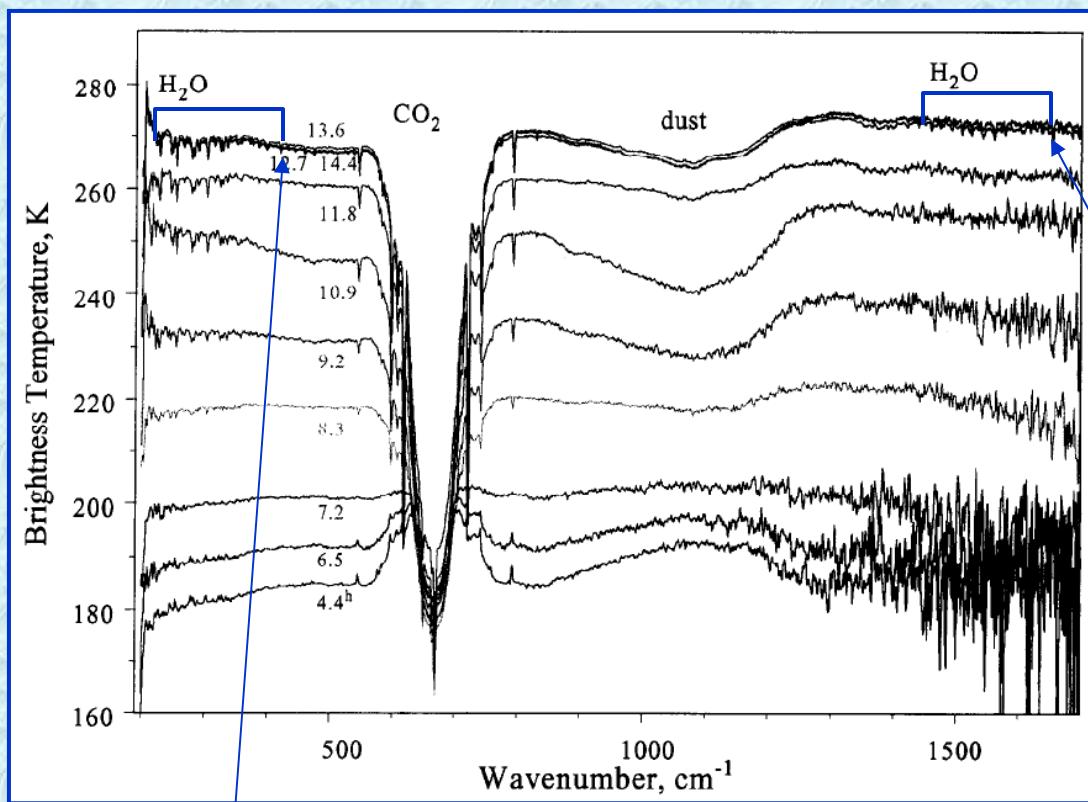
MAWD: grating spectrometer on Viking Orbiter 1, 2 dedicated to H<sub>2</sub>O  
5 channels; spectral resolution= 1.2 cm<sup>-1</sup>  
Farmer, LaPorte 1977; Farmer et al 1977; Jakosky, Farmer 1982  
Account for aerosol: Fedorova et al 2004  
Reprocessing: Fedorova et al 2010

Fedorova et al 2010



Synthetic spectra of water vapour in the 1.38-μm band, and the positions of MAWD channels

# Spectral bands and spectra: thermal infrared



“50-μm” band  
IRIS, TES, PFS LW

Ignatiev et al 2002,  
IRIS/Mariner 9 spectra

“6.3-μm” band  
IRIS

IRIS: Fourier-spectrometer on Mariner 9 (1971-72)

Many spectra, but only a limited part of seasonal cycle ( $L_s=290-340^\circ$ )

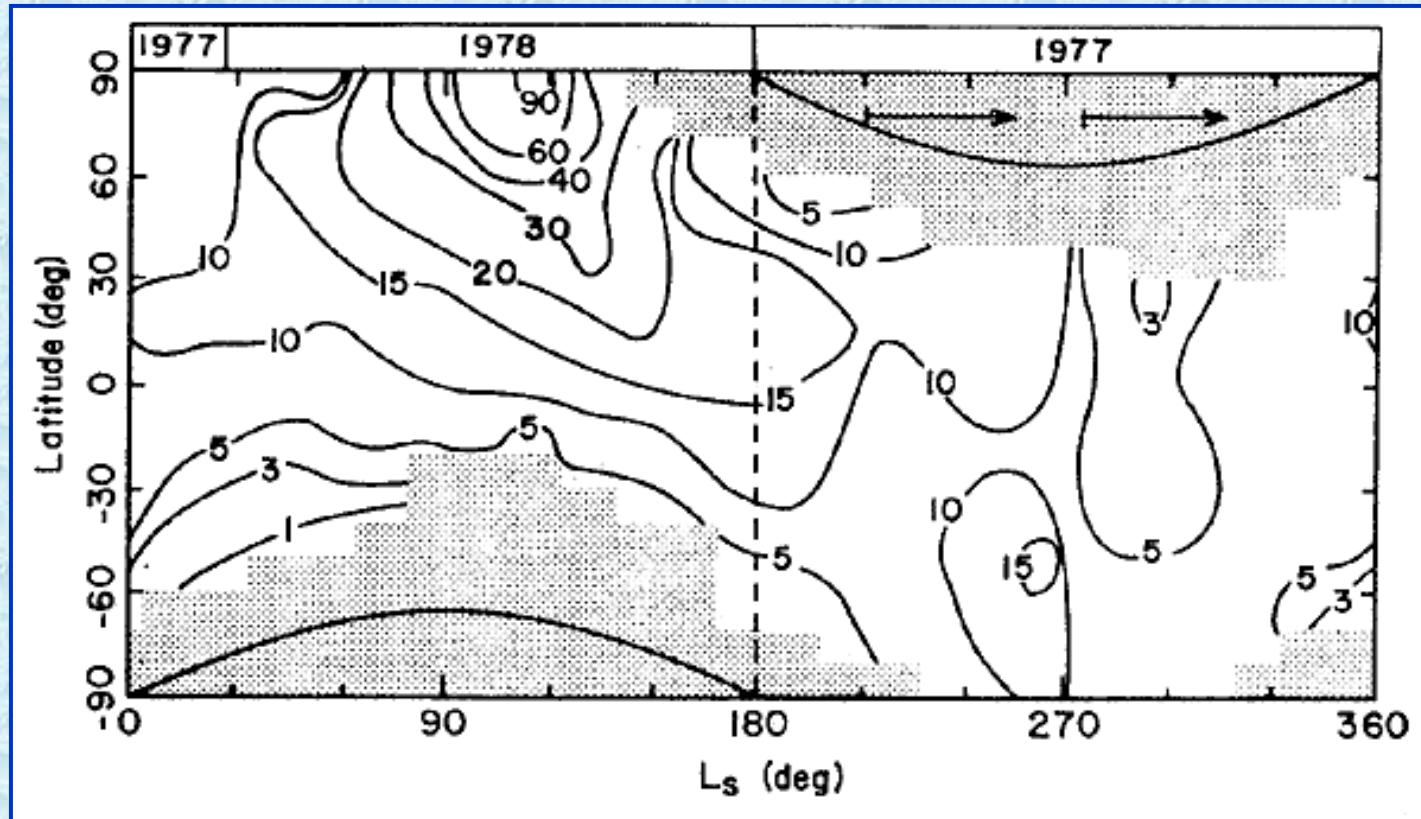
Spectral resolution =  $2.4 \text{ cm}^{-1}$

Water vapour retrieved by Hanel et al., 1972; Ignatiev et al., 2002

# Discovery and early observations of H<sub>2</sub>O on Mars

- Discovery:
  - Spinrad, Münch, Kaplan ApJ 1963: echelle spectrograph with photographic registration in 818–820-nm absorption band, using Doppler shift from telluric lines : 10 pr μm. Long follow-up with similar measurements, with photoelectric, then linear array registration.
- First spacecraft measurements:
  - Hanel et al. 1972: IRIS/Mariner 9, 50-μm (200-300 cm<sup>-1</sup>) emission band
  - Moroz, Nadjip 1975: IV/Mars 3, Mars 5, 1.38-μm absorption band
- First systematic measurement:
  - Jakosky, Farmer 1982: MAWD/Viking orbiters, 1.38-μm absorption band

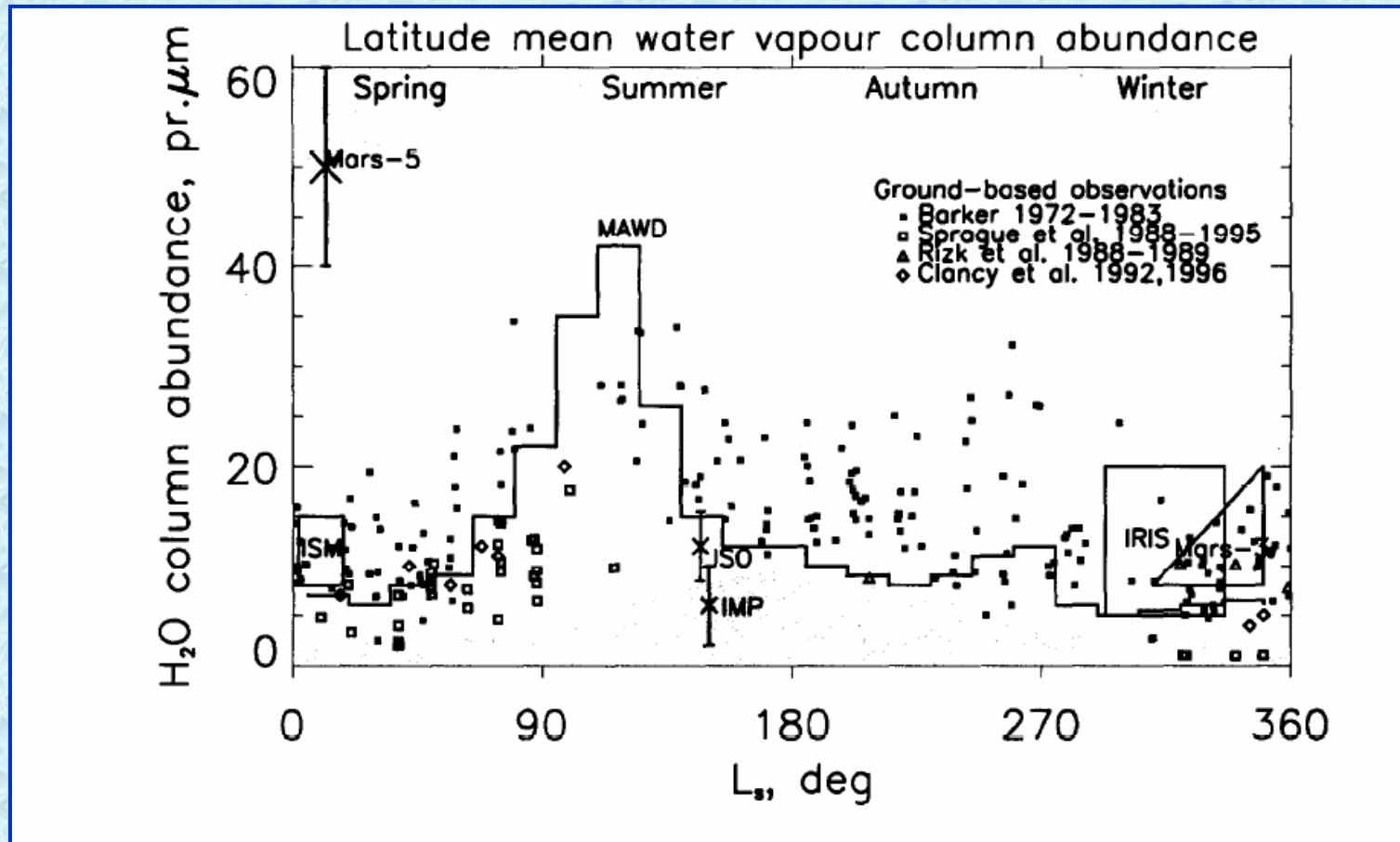
# MAWD Viking measurements of seasonal cycle



- Northern summer polar maximum (NPM)
- Asymmetry of the water cycle
- Basis for atmospheric models (e.g. Haberle, Jakosky 1991)
- Supposed exchange with regolith (e.g. Houben et al 1997; Jakosky et al 1997)
- Discarded exchange with regolith (Richardson, Wilson 2002)

# Comparison of early water vapour datasets

- Jakosky, Barker 1984
- Titov, 2002



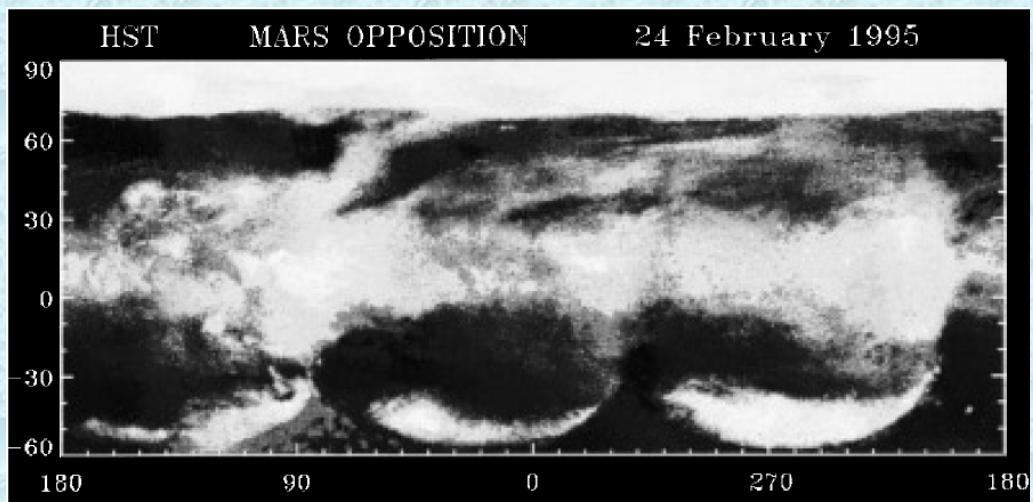


# Clancy et al., Icarus 1996:

- Clancy et al. monitored atmospheric temperature from microwave measurements of CO lines, 1980, 1982, 1989, 1991, 1993, 1995; H<sub>2</sub>O and clouds on some of these oppositions.
- They retrieved temperature 15-20K colder than Viking in aphelion seasons, low water amounts, and low saturation of water
- This associated with persistent low-latitudes clouds (aphelion cloud belt, ACB)

➔ “Clancy hypothesis”: Asymmetry of water cycle explained by orbit eccentricity

Aphelion cloud belt  
observed by Hubble  
Space Telescope (Clancy  
et al. 1996)



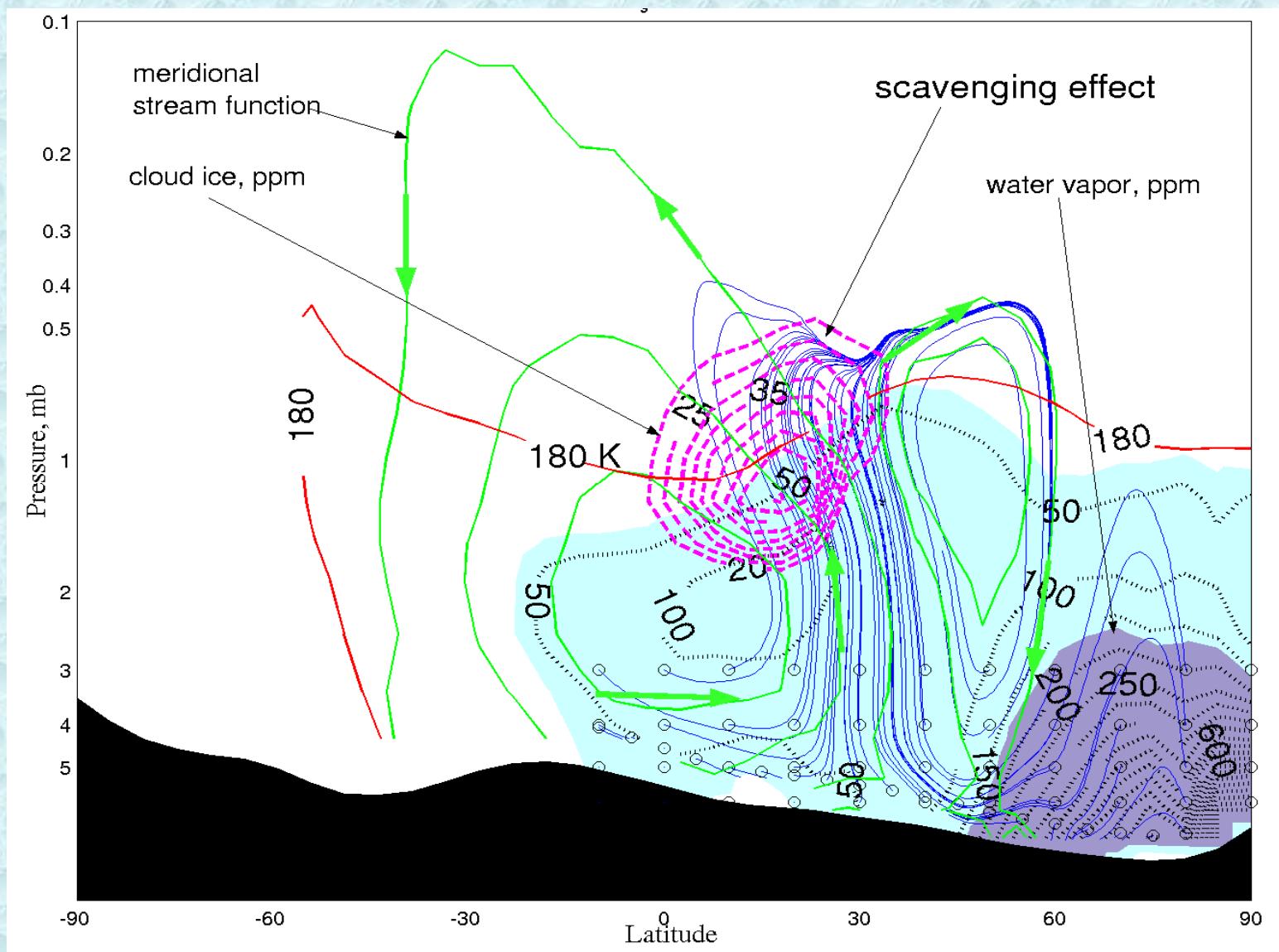
# Clancy hypothesis

- Unlike on Earth, Hadley circulation cell on Mars transports air from summer to winter hemisphere, and extends up to polar latitudes
- In cold aphelion (Northern summer), global equatorial cloud belt blocks water in the northern hemisphere.
- In warm perihelion (Southern summer) no clouds form, and water is distributed evenly

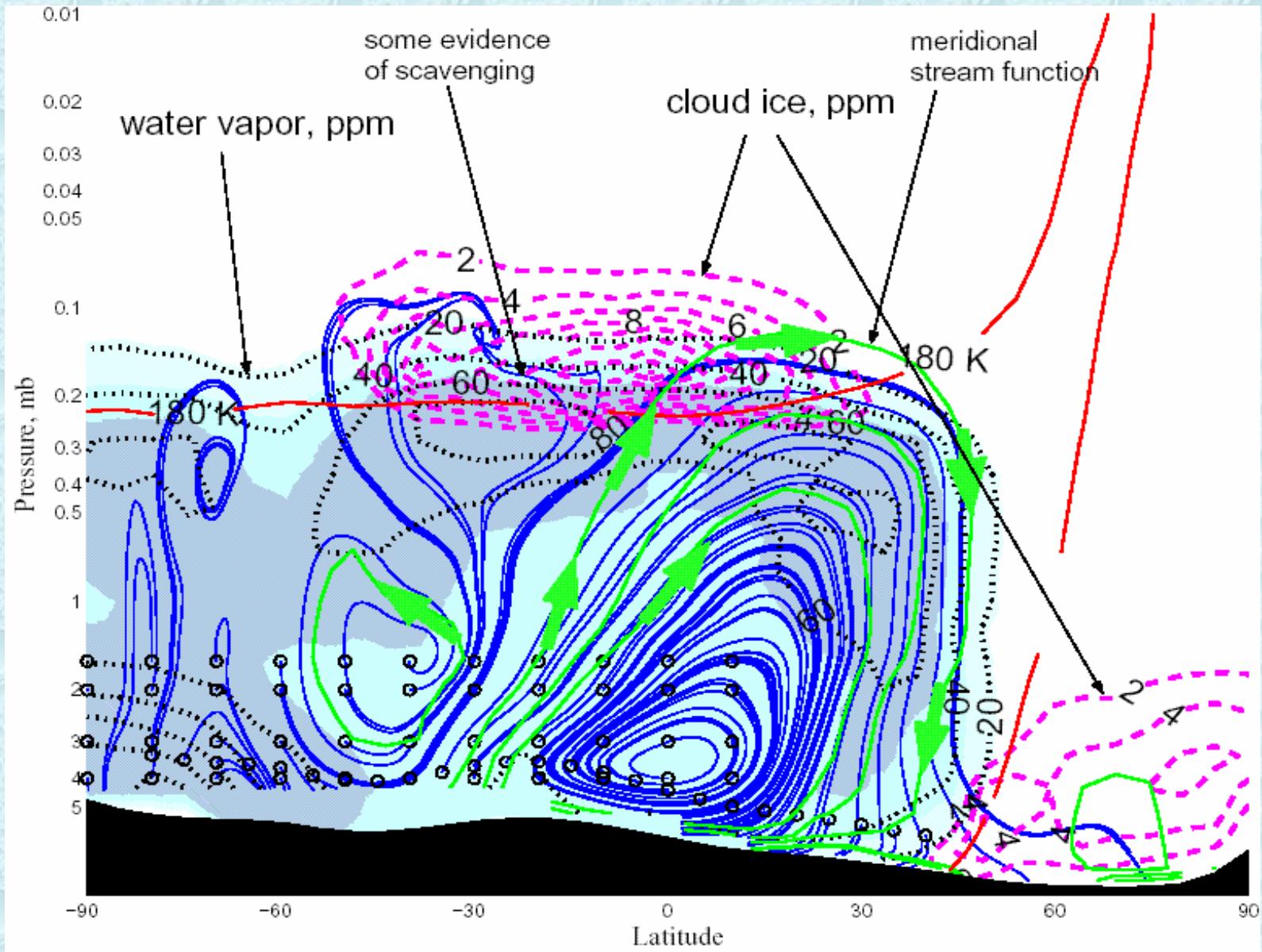
## Supporting facts:

- Colder Viking climate (correction of IRTM results, Wilson, Richardson, 2000)
- Discovery of water ice on perennial South polar cap (Bibring et al., 2004; inversion of perihelion is estimated at the scale of 50 000y, cf lecture by F. Forget)
- Dryer Viking climate (correction of MAWD results, Fedorova et al., 2010; see below)

# Aphelion atmospheric water cycle: MGCM



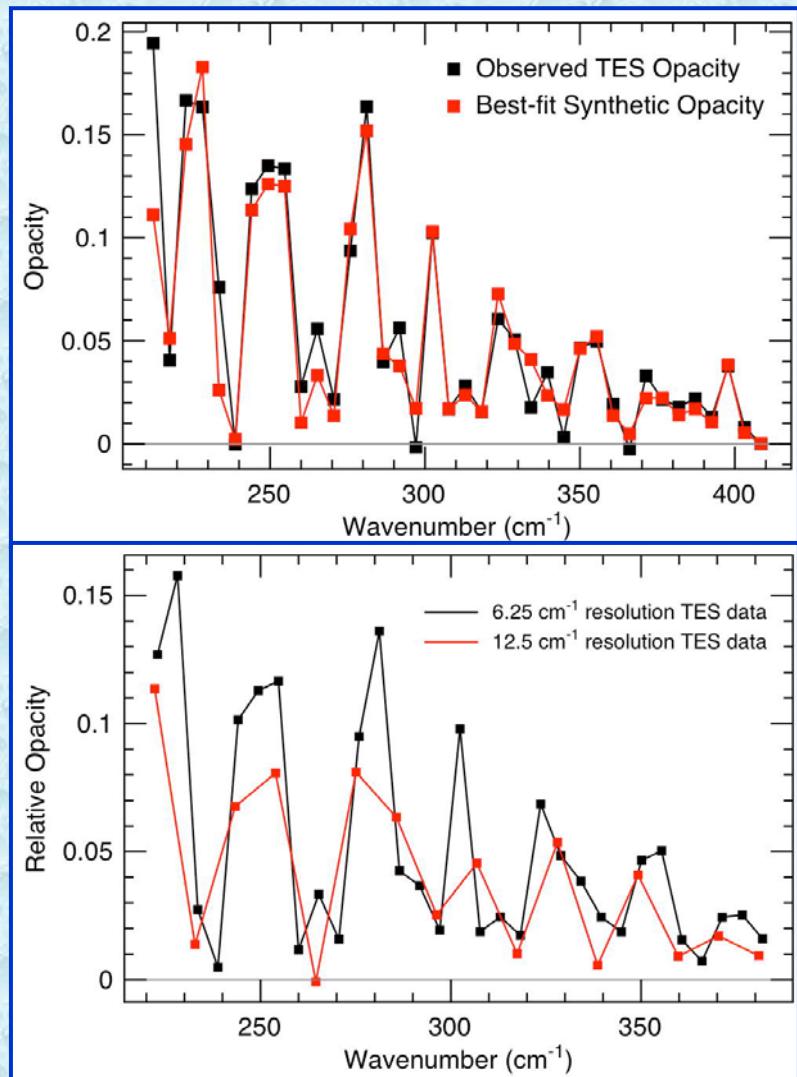
# Perihelion atmospheric water cycle: MGCM



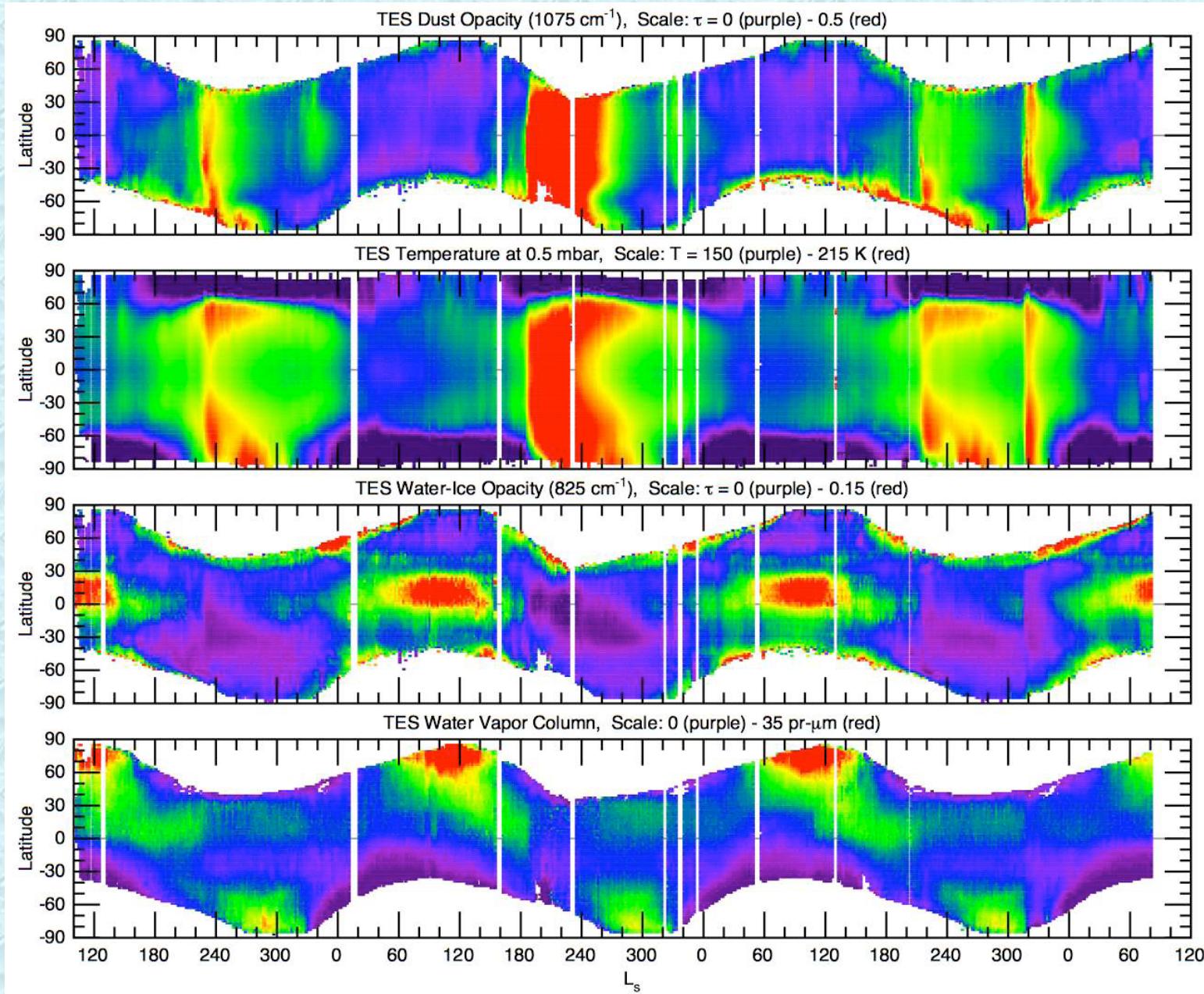


# Recent measurements: I. TES/MGS

- TES: Fourier-spectrometer on MGS (1999-2004)
- Spectral resolution=  $6.25$  or  $12.5 \text{ cm}^{-1}$
- Nearly polar orbit, 12 orbits/day,  $\sim 2:00$  AM and 2:00 PM
- Measures temperature, dust opacity, water ice clouds opacity, and water vapor (MD Smith, 2002, 2004, 2008)



# MGS TES: dust, temperature, clouds and water vapour

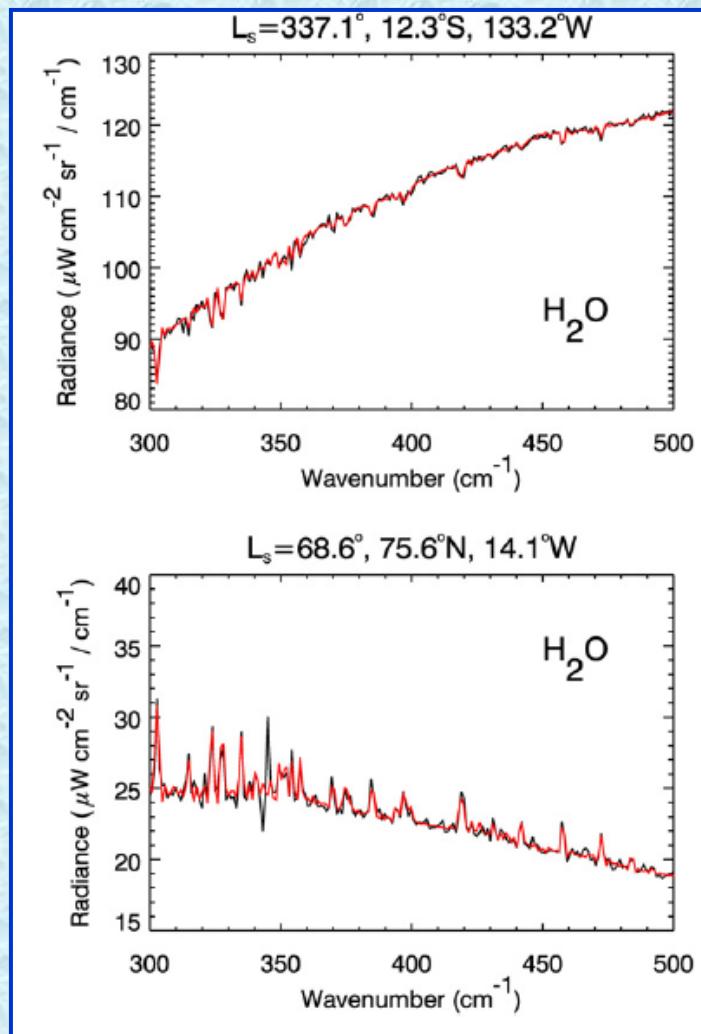
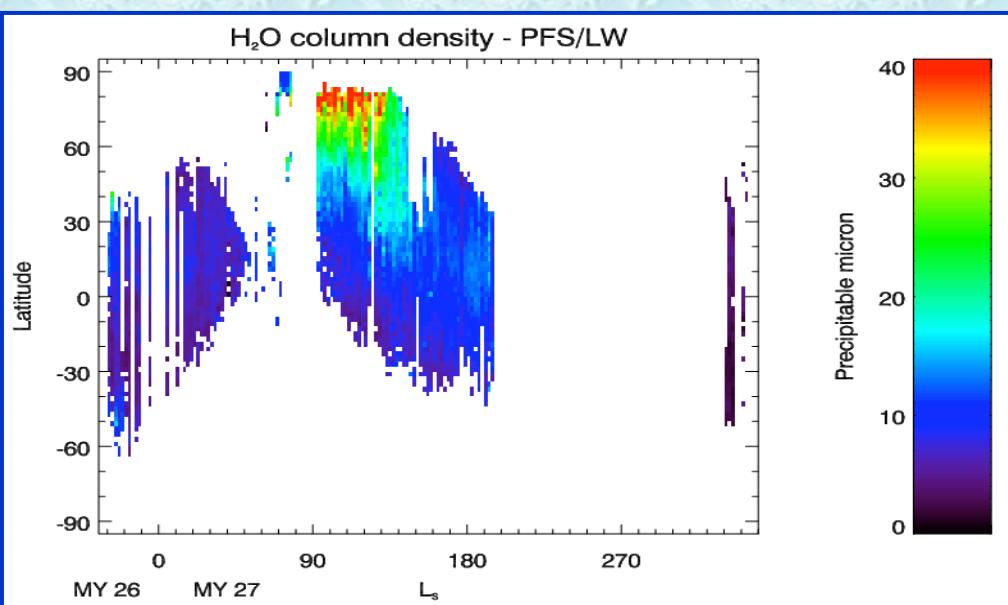


# Recent measurements: II. Mars Express

- PFS spectral resolution  $1.4 \text{ cm}^{-1}$ 
  - LW 25-35  $\mu\text{m}$ : processed by 2 groups: Fouchet et al., Icarus 2007
  - SW 2.56  $\mu\text{m}$ : processed by 3 groups: Tschimmel et al., Icarus 2008, Encrenaz et al. A&A 2009; Sindoni, Formisano, 2010.
- OMEGA 2.56  $\mu\text{m}$  low spectral resolution, but mapping. Processed by 2 groups
  - Encrenaz et al. A&A 2005, 2008, 2009; Melchiorri et al. PSS 2007, Icarus 2009
  - Maltagliati et al., 2008
- SPICAM 1.37  $\mu\text{m}$  spectral resolution  $3.5 \text{ cm}^{-1}$  processed by
  - Fedorova et al., JGR 2006, Icarus 2010

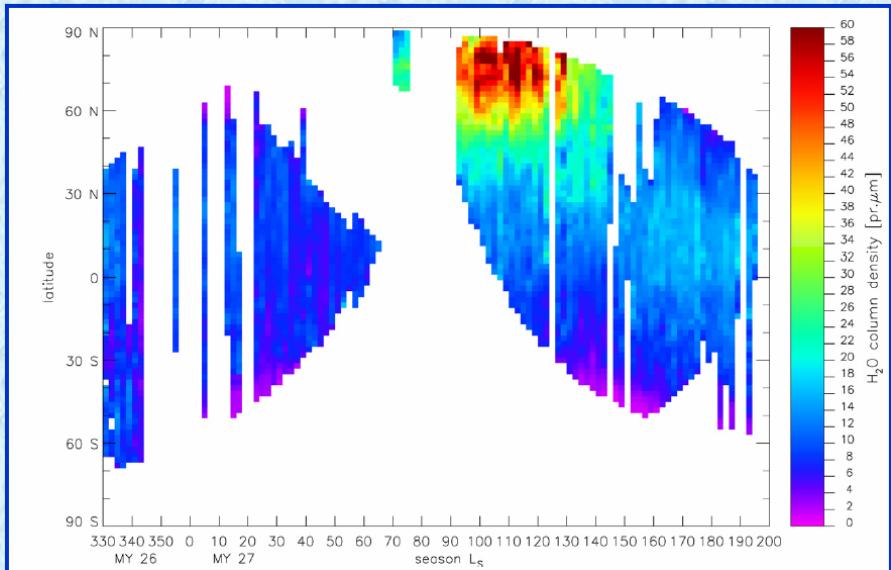
# PFS LW/ Mars Express

PFS-LW: Fourier-spectrometer  
Spectral resolution=  $1.4 \text{ cm}^{-1}$   
Only a fraction of data available  
 $\text{H}_2\text{O}$  retrieved by Fouchet et al 2007



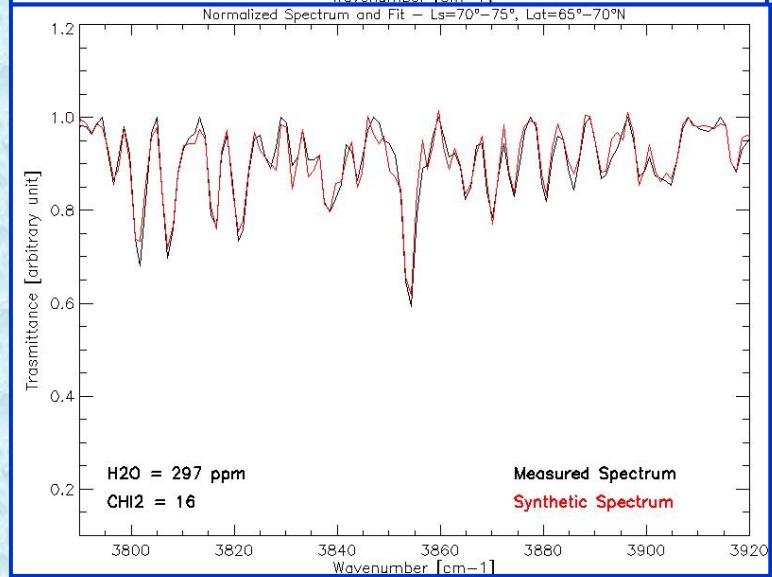
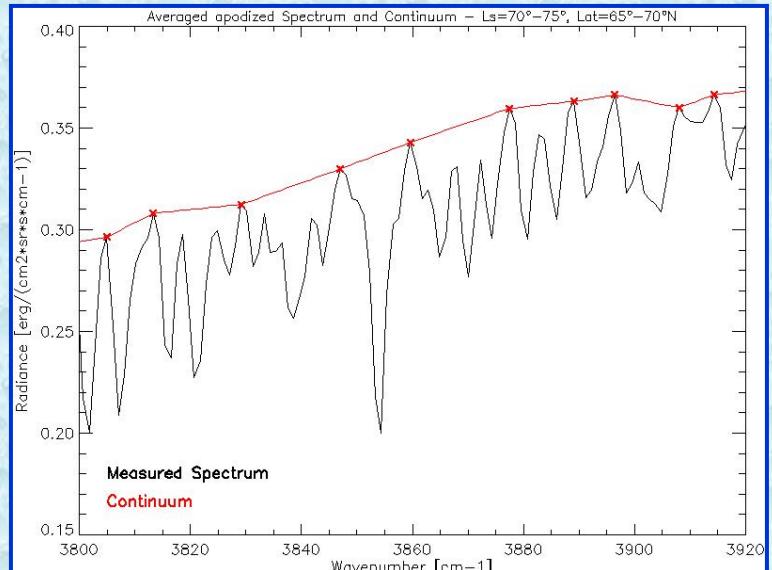
Fouchet et al., Icarus 2007

# PFS-SW/Mars Express



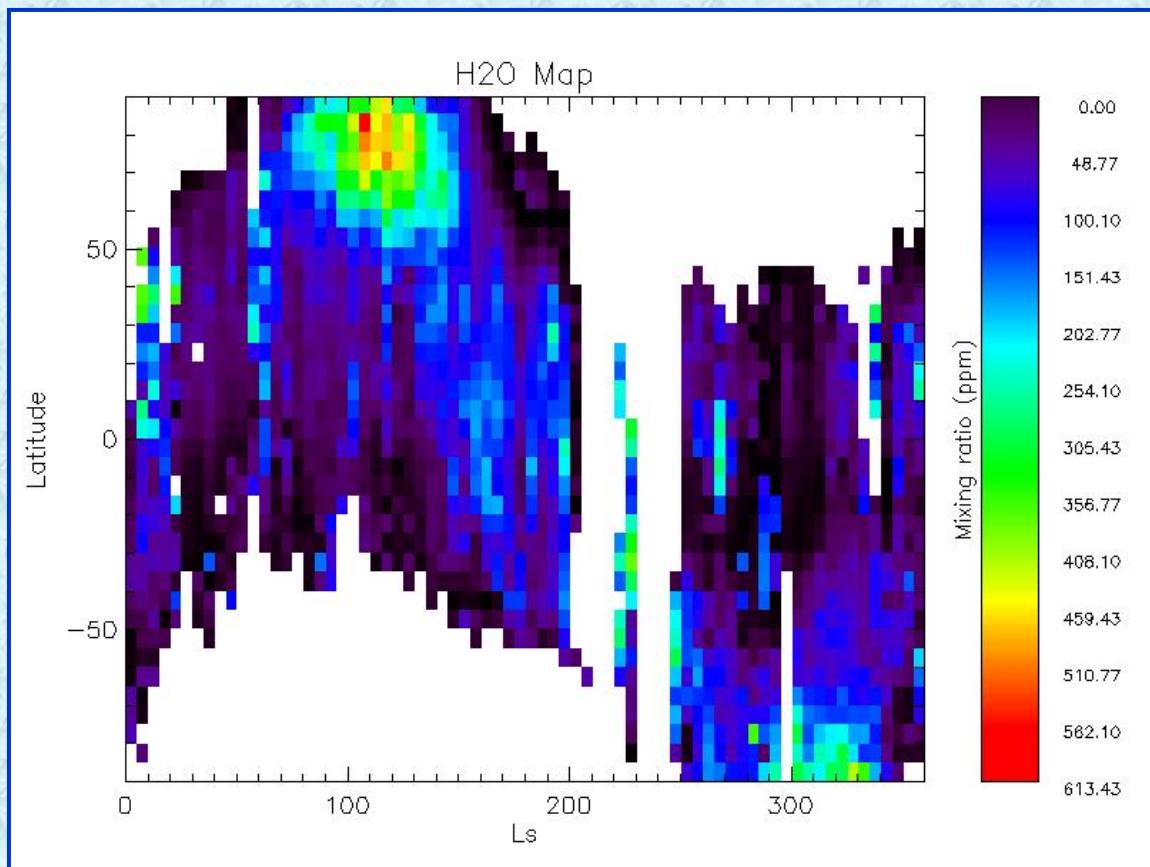
Tschimmel et al Icarus 2008

- PFS-SW: Fourier-spectrometer
- Spectral resolution=  $1.4 \text{ cm}^{-1}$
- Problems with spectral continuum; work in progress
- $\text{H}_2\text{O}$  retrieved by Tschimmel et al 2008;  
Sindoni, Formisano 2010



Sindoni, Formisano 2010

# PFS LW/ Mars Express



- PFS-SW by Sindoni, Formisano 2010 : full set of data on one seasonal graph

# OMEGA: spectra

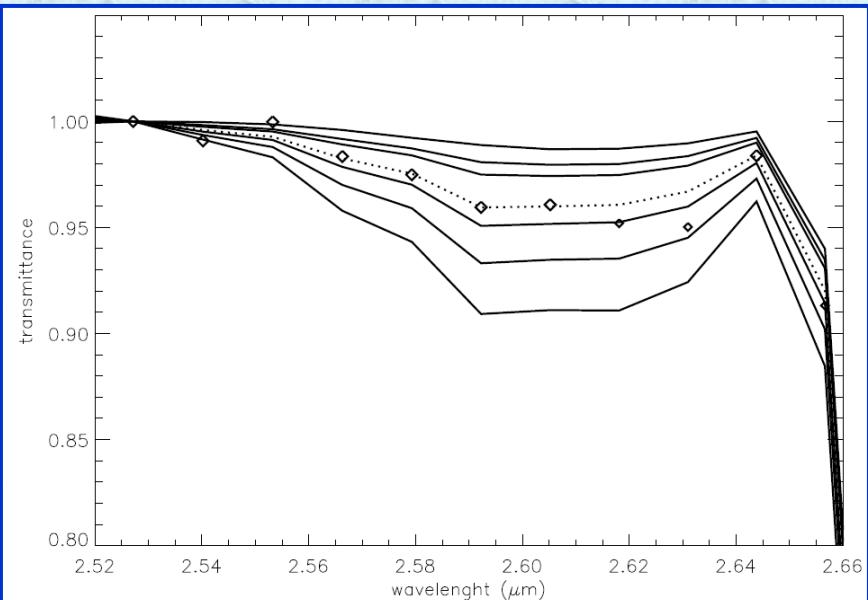
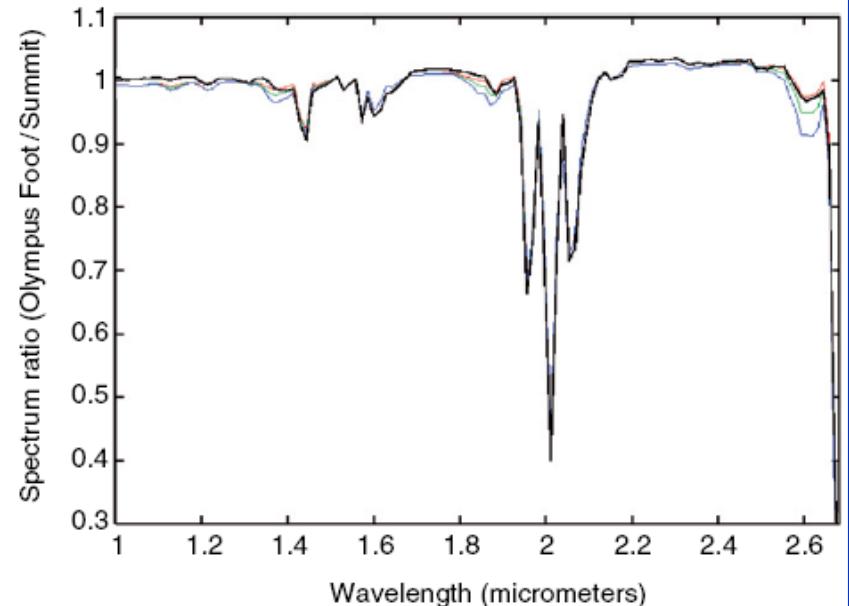
OMEGA: grating spectrometer

Resolving power  $\sim 300$

Mapping capability

Correction for aerosol is an issue

Water vapour retrieved by  
Meudon and MPS teams



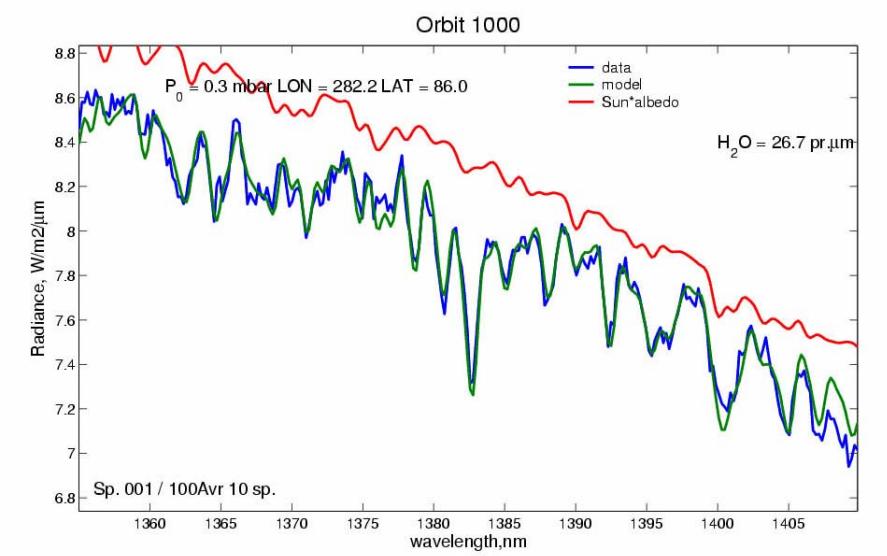
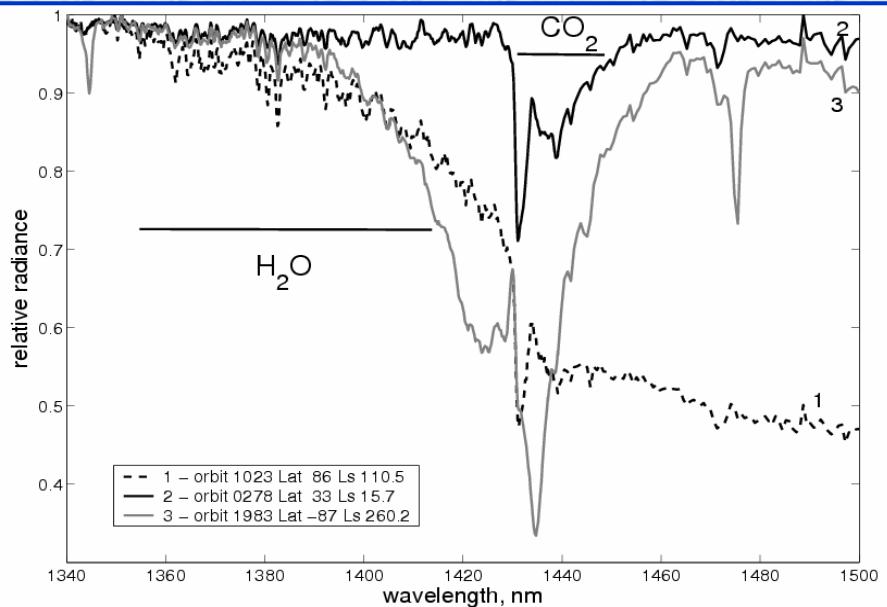
# SPICAM/Mars Express

SPICAM: AOTF spectrometer

Spectral resolution=  $3.5 \text{ cm}^{-1}$

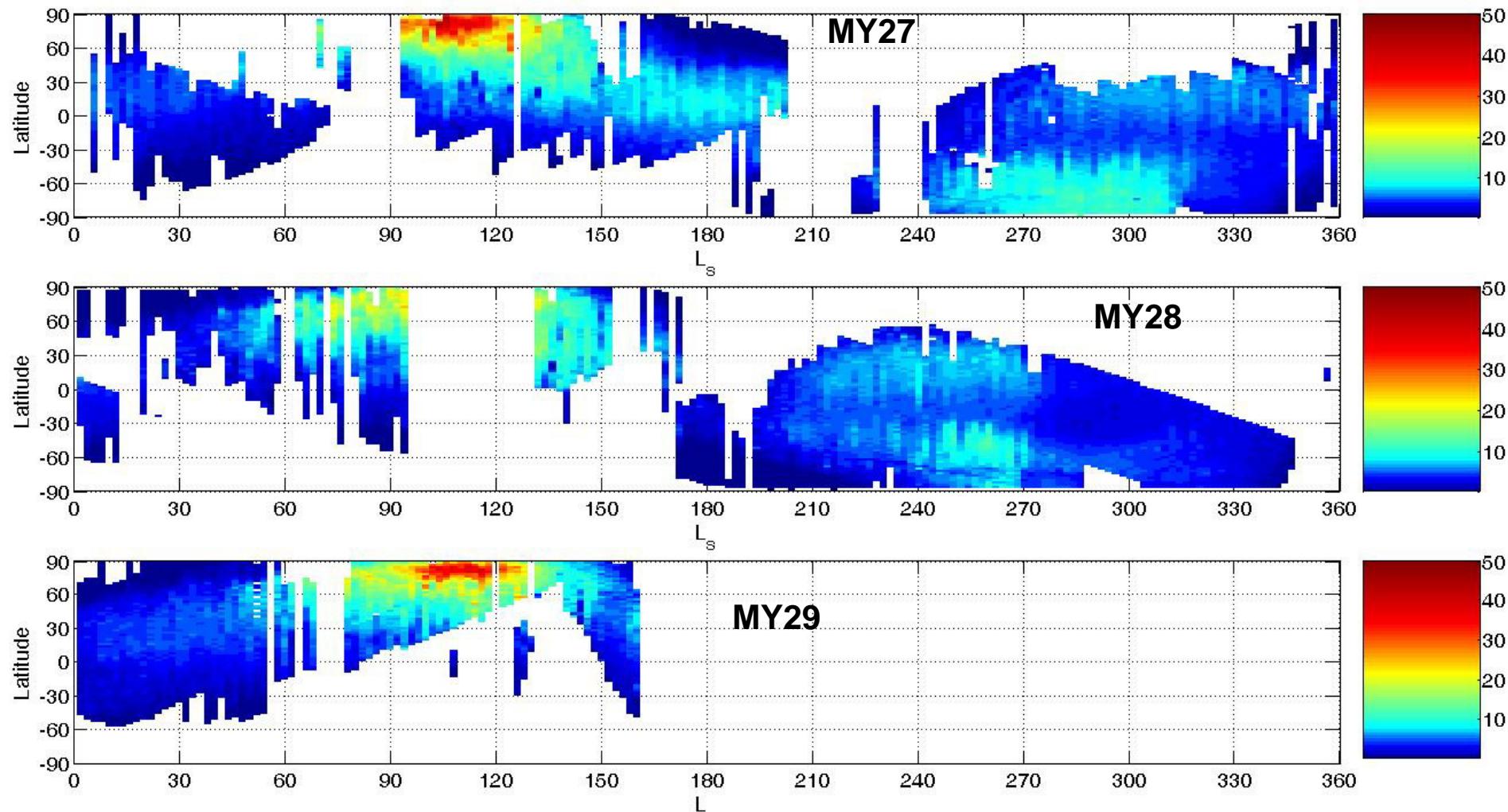
Correction for aerosol is an issue

Water vapour retrieved by  
Fedorova et al., 2006, 2010



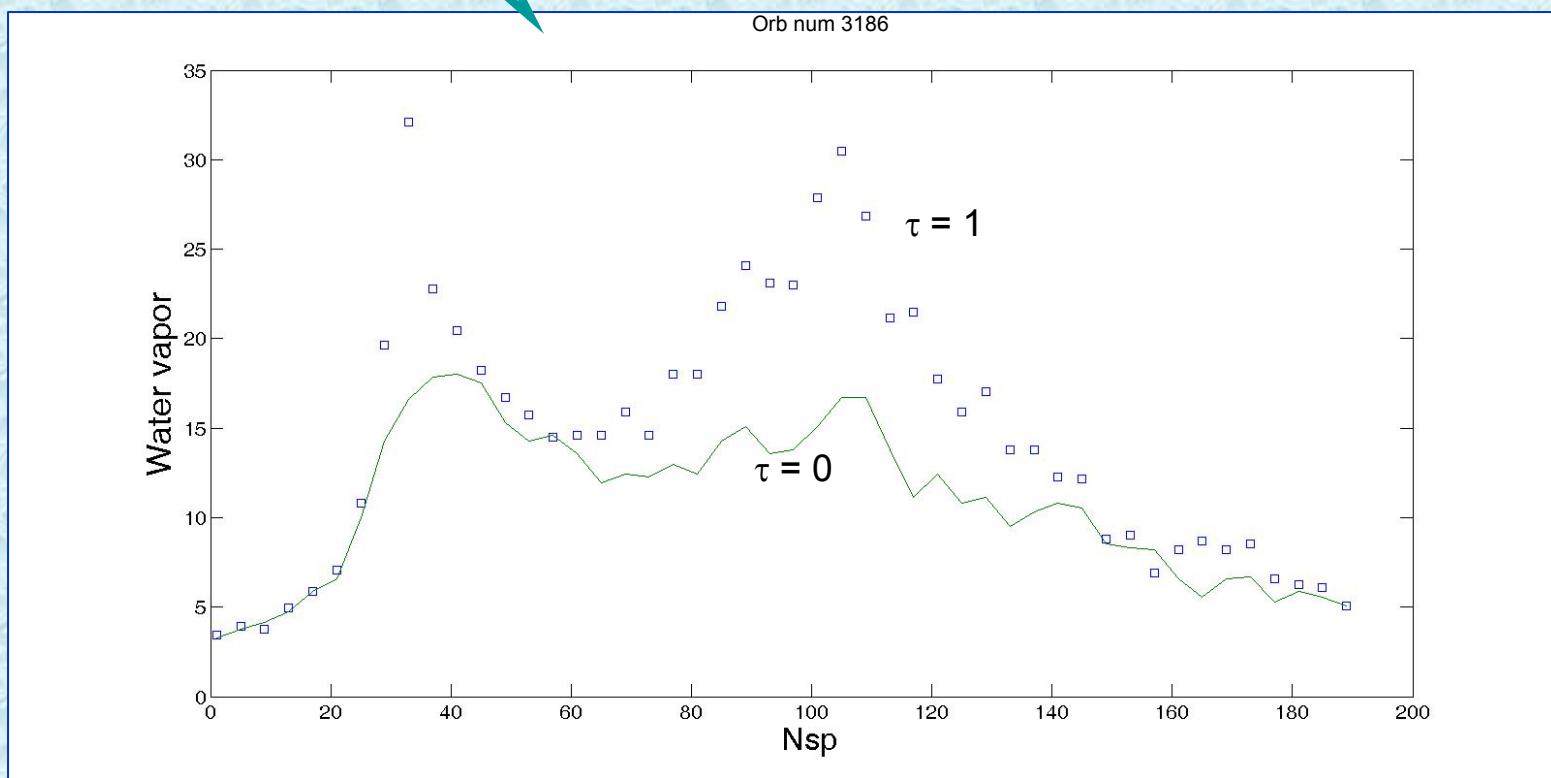
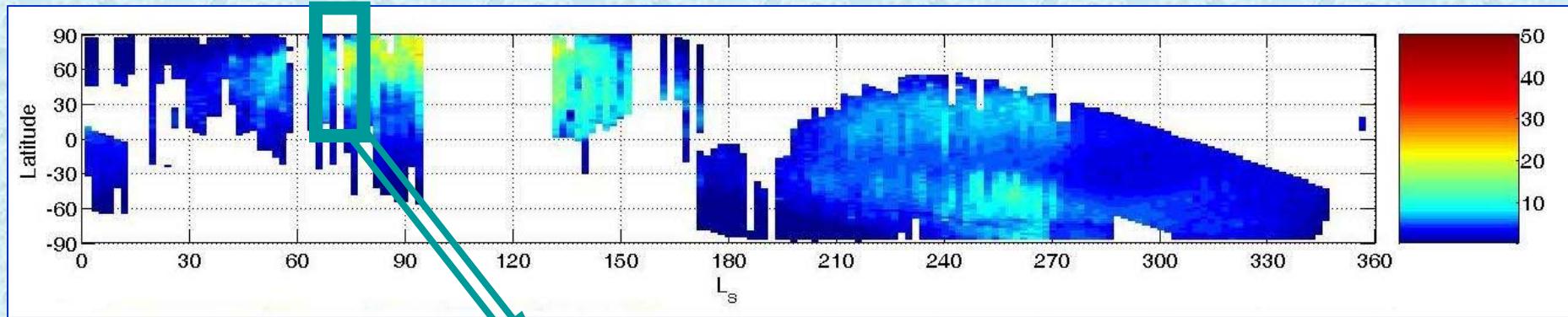
# SPICAM: seasonal distribution of H<sub>2</sub>O

H<sub>2</sub>O, pr.μm



Water vapor band 1.38 mm, HITRAN2004, Martian Climate Database V4.2 Trokhimovskiy et al 2009 EPSC

# Aerosol bias on SPICAM H<sub>2</sub>O retrieval: worst case





# **WATER IN MARS ATMOSPHERE: COMPARISON OF RECENT DATA SETS (2005-2006)**

O. Korablev, JL Bertaux, RT Clancy, T. Encrenaz, A. Fedorova, V. Formisano, T. Fouchet, N. Ignatiev, E. Lellouch, L. Maltagliati, R. Melchiorri, F. Montmessin, A. Rodin, MD Smith, A Sprague, D. Titov, A. Trokhimovskiy, M. Tschimmel, L. Zasova,

**Data of 3 Mars Express instruments, TES/MGS, and ground-based  
compared and validated**

# Assumptions and methods for H<sub>2</sub>O retrieval in MEX experiments.

Instrument	Spectral data, broadening	Synth spectrum comp.	Temp. profile	Saturation altitude	Pressure definition, footprint	PFS data processing	Dust-aerosol	Solar spectrum	Comment
PFS LW (NI)	HITRAN 04 Gamache95	LUT from LBL step 0.002 15T x10p	D. Grassi ~60 layers, adaptive like EMCD	From T mixed below	EMCD, MOLA with 0.5° averaged along PFS track (>10sp)	Calibrated by Giuranna apod	no	N/A	
PFS LW (TF)	GEISA 03 Gamache93	LUT from LBL step 0.01 12T x45p	15-μm inv. 45 layers, 12mb	From T mixed below supersaturation allowed	Same 9 sp	Calibrated by Giuranna unapod	no	N/A	
TES	GEISA 97 1.5	Correlated k, LUT for H <sub>2</sub> O 10Tx10p	15-mm inv (Conrath 2000)	From T mixed below	MOLA+ VL seasonal	n/a	Derived from TES 9-μm	N/A	

Instrument	Spectral data, broadening	Synth spectrum comp.	Temp. profile	Saturation altitude	Pressure definition, footprint	PFS data processing	Dust-aerosol	Solar spectrum	Referecne
PFS SW	HITRAN 04 Gamache95	Same as LW	EMCD	From T mixed below	15 sp	Uncalibrated, no phase corr, apod.	no	Fiorenza &F. 2005	Nicolay Ignatiev, personal communication
PFS SW	GEISA 03 Gamache93	Same as LW	EMCD	From T mixed below supersaturation allowed	16 sp	Uncalibrated, no phase corr, unapod.	no	Fiorenza &F. 2005	Fouchet et al., Icarus 2007
PFS SW	HITRAN 04 Gamache	LUT from LBL 20T	EMCD 0-98 km step 2	From T Mixed OR CONFINED below	15 sp with MOLA; press EMCD	Calibrated, no phase corr, apod.	no	Fiorenza &F. 2005	Tschimmel et al., Icarus 2008
SPICAM	HITRAN 04 1.5	LUT from LBL adaptive grid, min 0.001 16T x13p	LMD, grid is kept	From T 0 water above, mixed below	EMCD (3.5x5°), 6 or 10 sp	n/a	no	Fiorenza &F. 2005	Fedorova et al., JGR 2006; corrected in Icarus 2010, in press
OMEGA	GEISA 03 Old Gamache (1993)	Curves of growth, varied p0	single profile	From a chosen T profile	"Pres0" (0.1°, up to 1/32°) Omega data at 0.1x0.1°	n/a	no	Relative to Olympus top	Encrenaz et al. A&A 2005, 2008 Melchiorri et al. PSS 2007, Icarus 2009
OMEGA	HITRAN 04 Gamache	LUT from LBL 20T	EMCD, 5 profiles per cube	From T mixed below	EMCD	n/a	no	Fiorenza &F. 2005	Maltagliaty et al., Icarus 2008

# Side effects of MEX water vapour comparisons

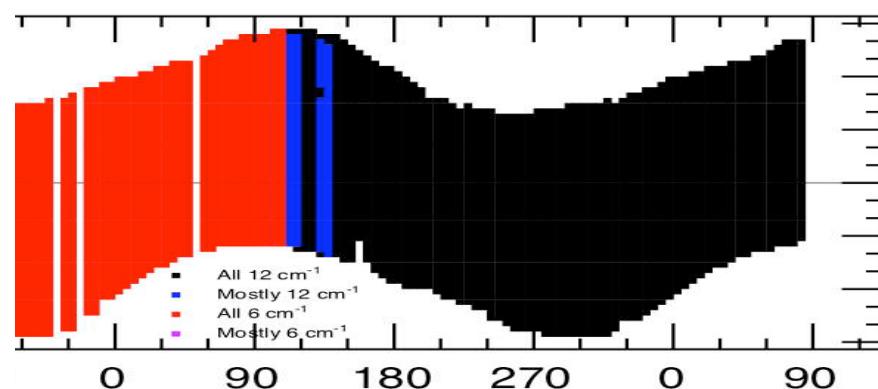
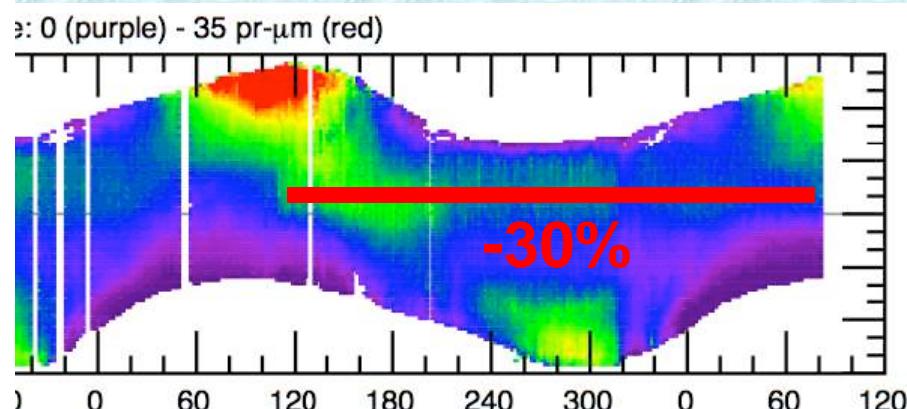
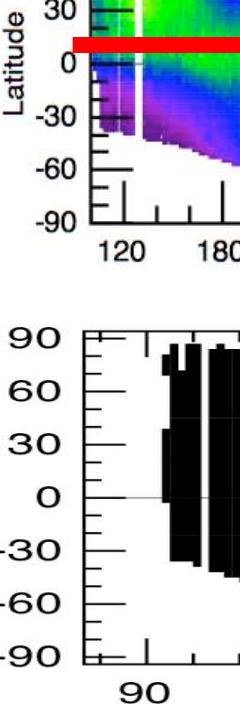
- TES/MGS database modification
  - Bug in processing low resolution ( $12.5\text{ cm}^{-1}$ ) portion of data
  - Reduction of  $\text{H}_2\text{O}$  content in these data by **30%**
- MAWD/Viking dataset modification
  - Reprocessing of MAWD with new spectroscopic database (HITRAN 2004)
  - Reduction of the entire dataset by **60%**

# TES database modification

## Modification to TES Water Vapor Retrieval:

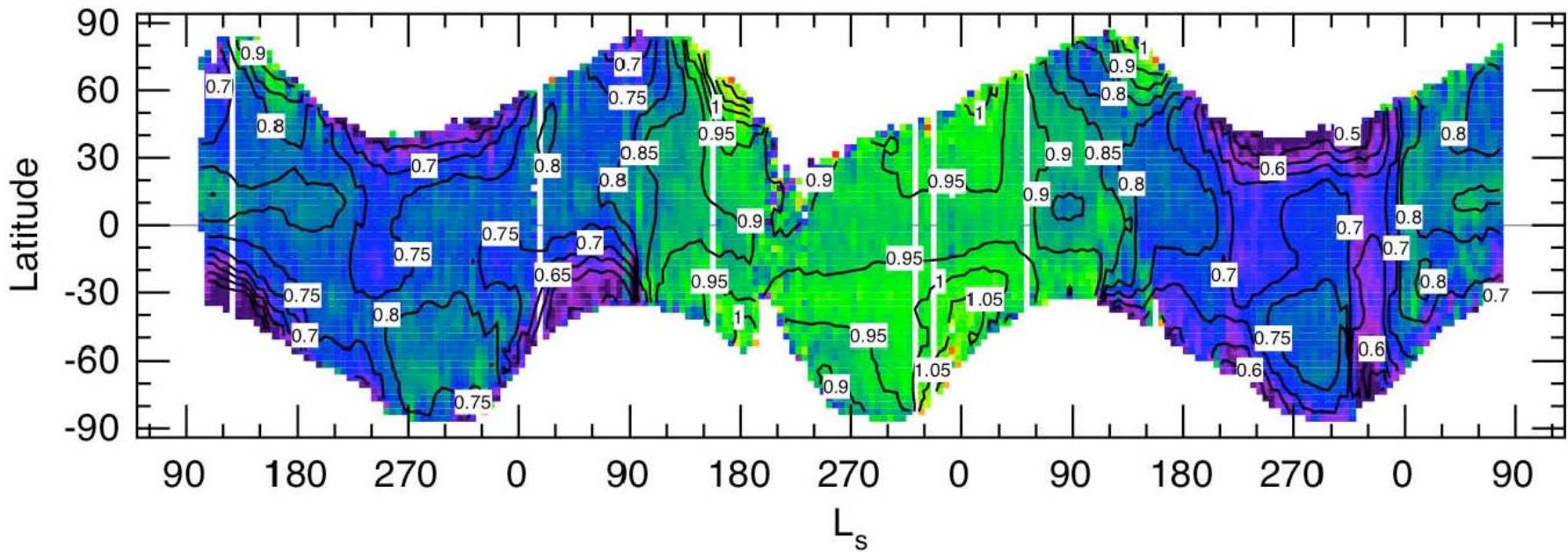
A reanalysis of the TES water vapor retrieval was performed in early 2006 when a comparison against concurrent retrievals from PFS/LW spectra (at the same wavelength as TES) showed that the TES retrievals gave significantly higher water vapor abundances than the corresponding PFS/LW retrievals (*Fouchet et al.* 2007). The analysis showed that for the spectra taken at the TES higher spectral resolution ( $6.25\text{ cm}^{-1}$ ), retrievals using each of the four water vapor bands gave consistent results, but for spectra taken at the lower spectral resolution ( $12.5\text{ cm}^{-1}$ ), retrievals using the two weaker bands ( $300$  and  $330\text{ cm}^{-1}$ ) were significantly higher and not consistent with other observations. Using only the two stronger bands at  $250$  and  $280\text{ cm}^{-1}$ , retrieved water vapor abundances are consistent between the higher and lower spectral resolution TES observations, and are much closer to the values retrieved using PFS/LW spectra.

This correction to the TES water vapor retrieval results in a reduction in derived water vapor abundance by about 30% in the TES lower resolution data. The TES higher resolution data are not affected by this correction. Lower resolution data were taken at the beginning (MY 24,  $L_s=103^\circ$  to MY 25,  $L_s=90^\circ$ ) and end of the TES dataset (MY 26,  $L_s=120^\circ$  to MY 27,  $L_s=80^\circ$ ), while higher resolution data were taken during the middle of the TES dataset



# TES database modification

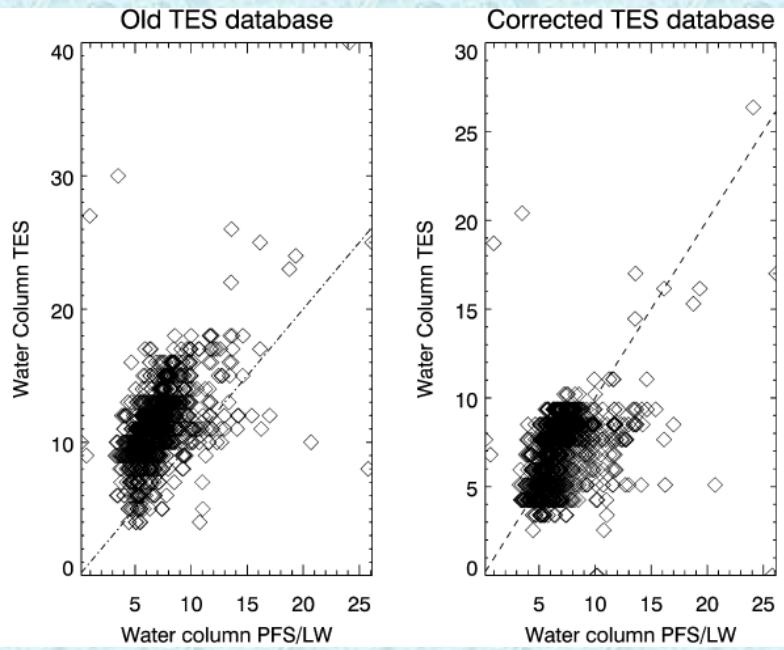
## Ratio of “New” / “Old” abundance



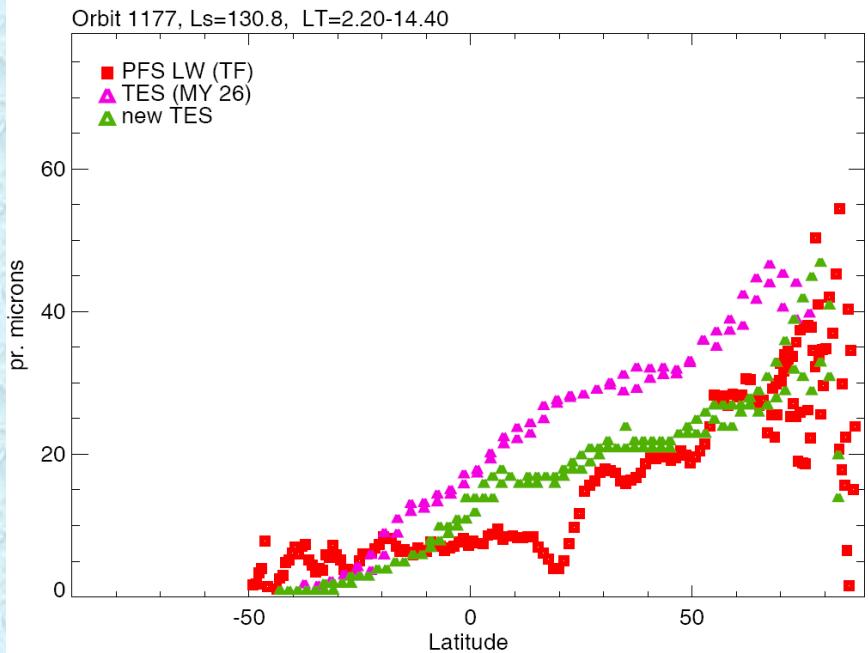
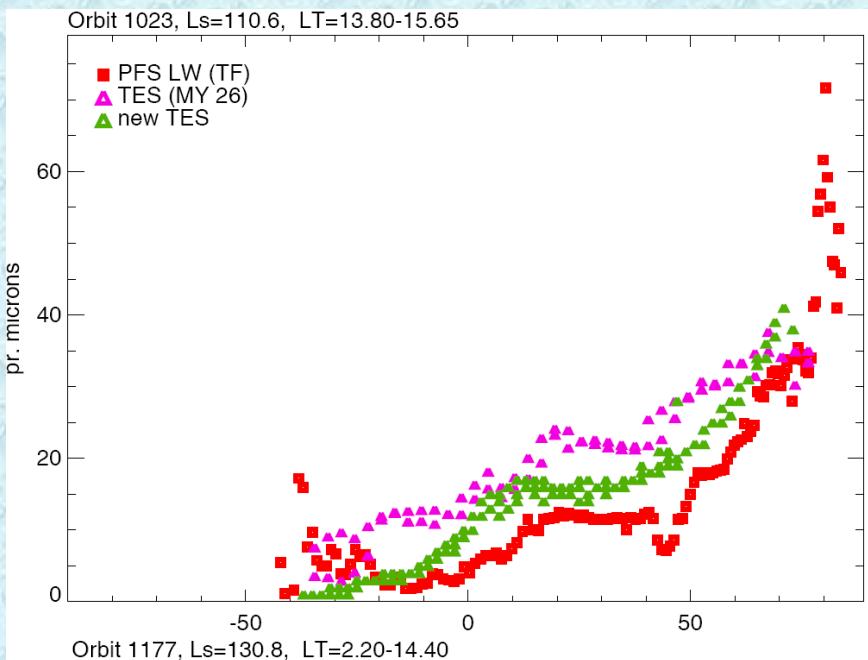
- Higher resolution observations: decrease of ~5%
- Lower resolution observations: decrease of ~20-30%
- More consistency from year to year (from higher to lower resolution data)

# PFS LW vs TES

Good agreement after  
correction of TES

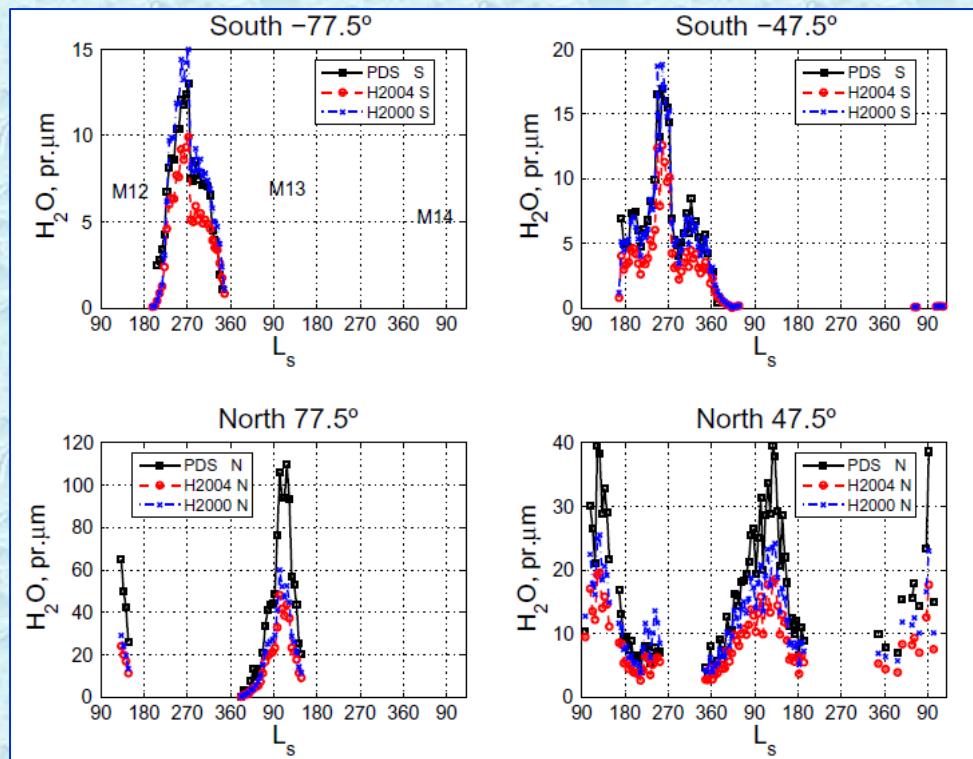


Fouchet et al., Icarus 2007



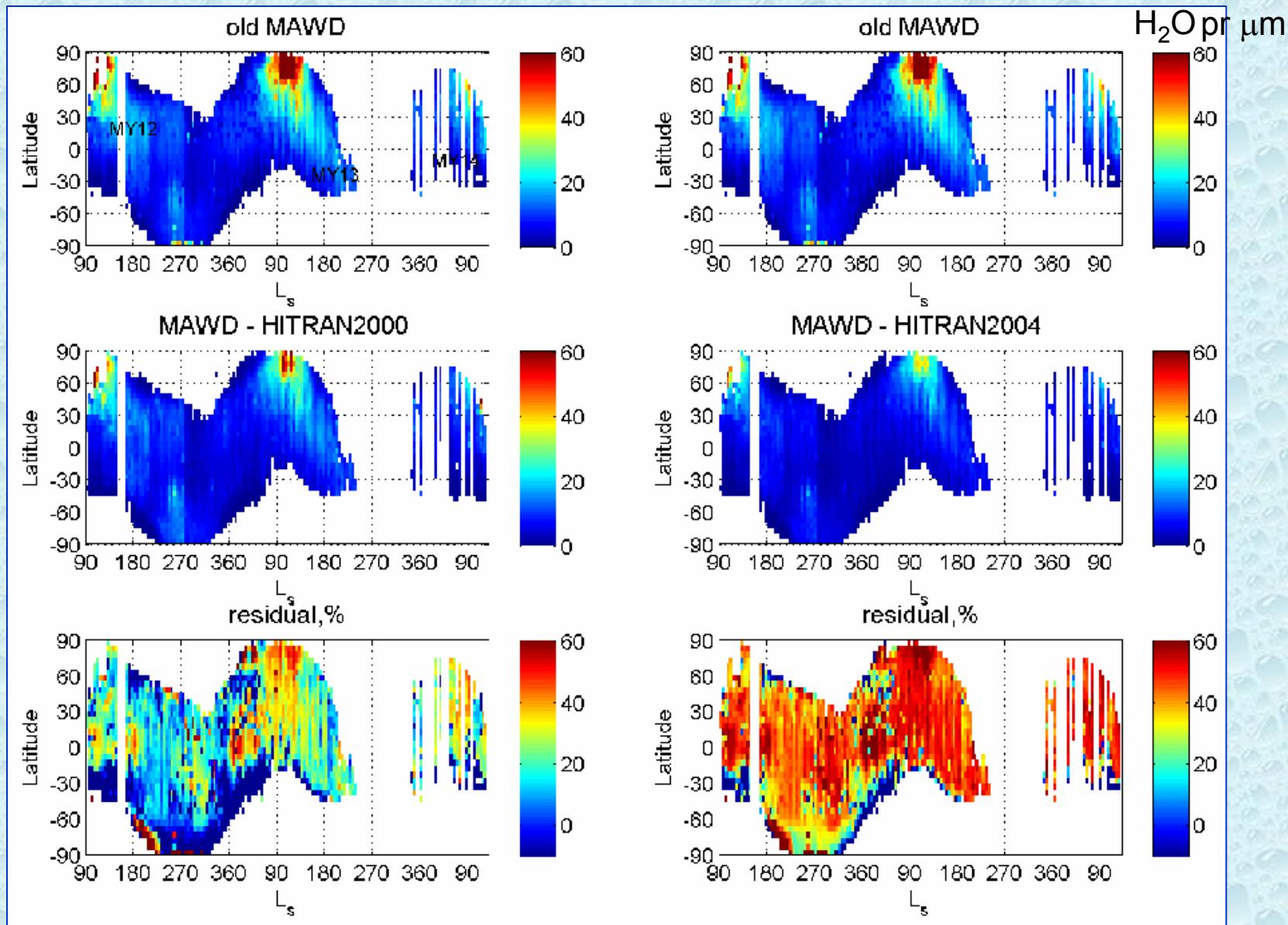
# MAWD database modification

- Sensitivity to spectroscopic database
  - HITRAN 2004 (and later): x2 more H<sub>2</sub>O lines in the 1.38-μm band
  - Different line parameters
  - More apparent absorption → less retrieved water
- More realistic retrieval assumptions
  - EMCD 4.2 atmospheric profiles
  - Lower saturation altitude → less retrieved water
  - Accurate account for solar spectrum
  - MOLA topography

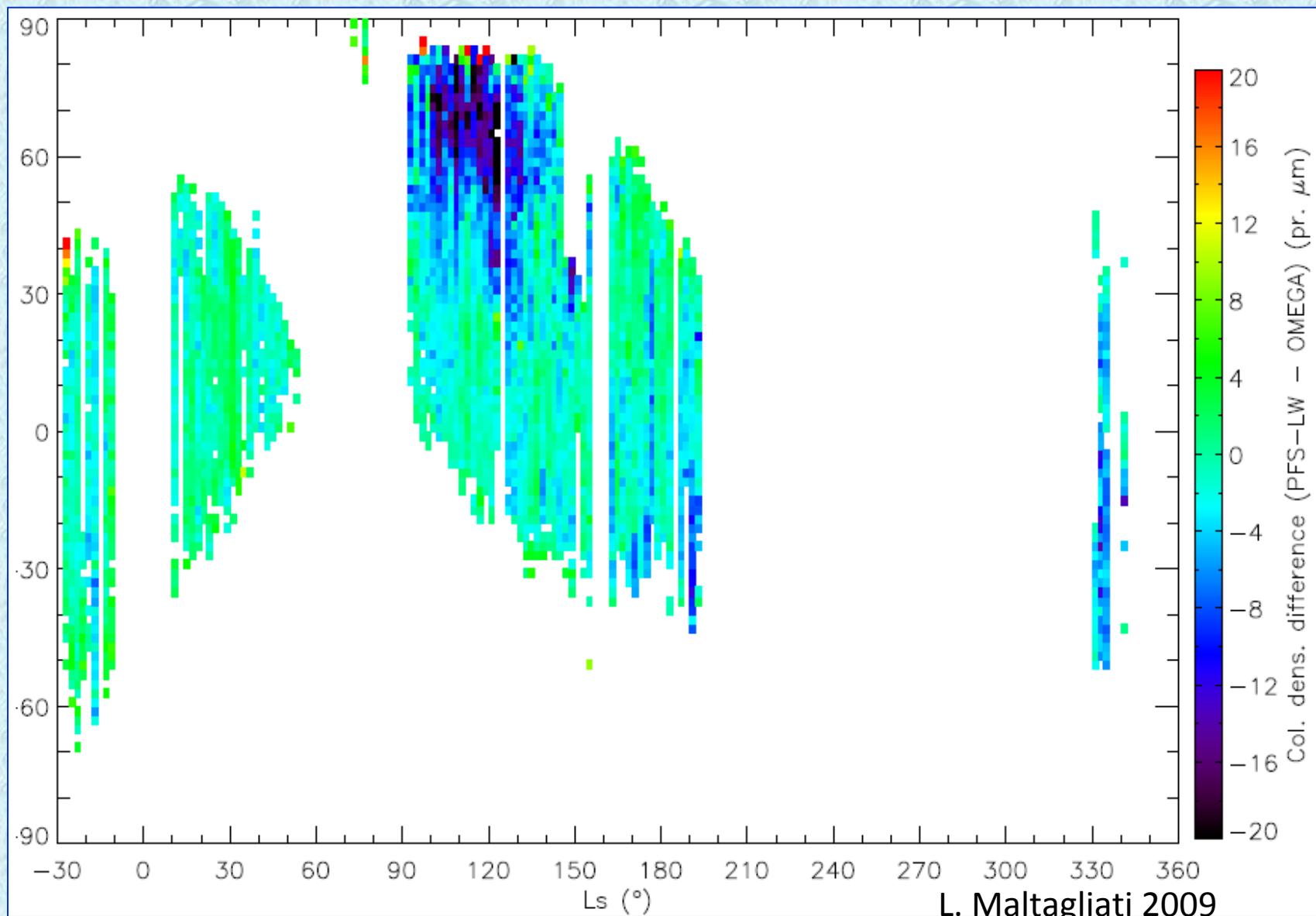


Fedorova et al., 2010

# MAWD database modification: results

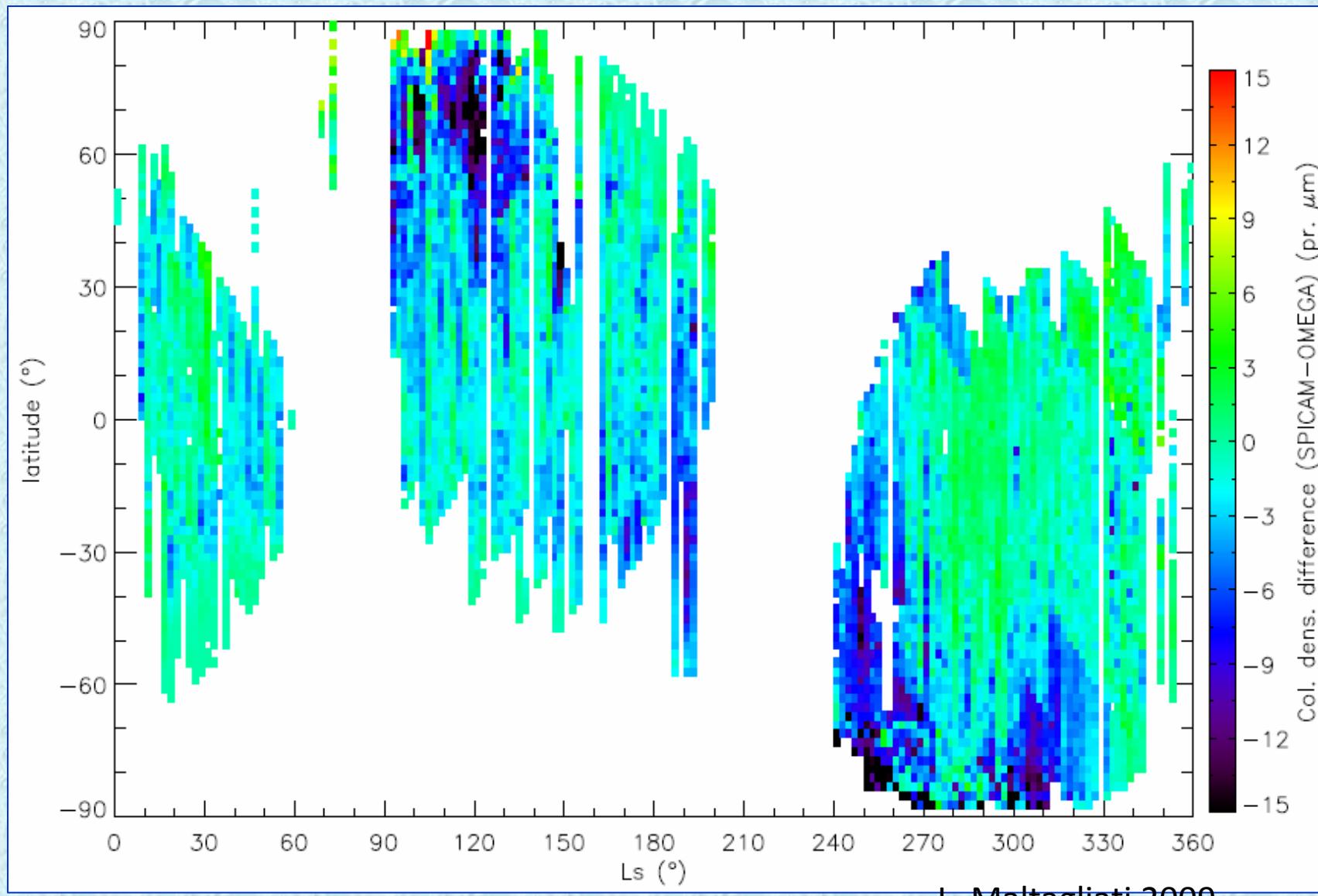


# PFS LW-OMEGA

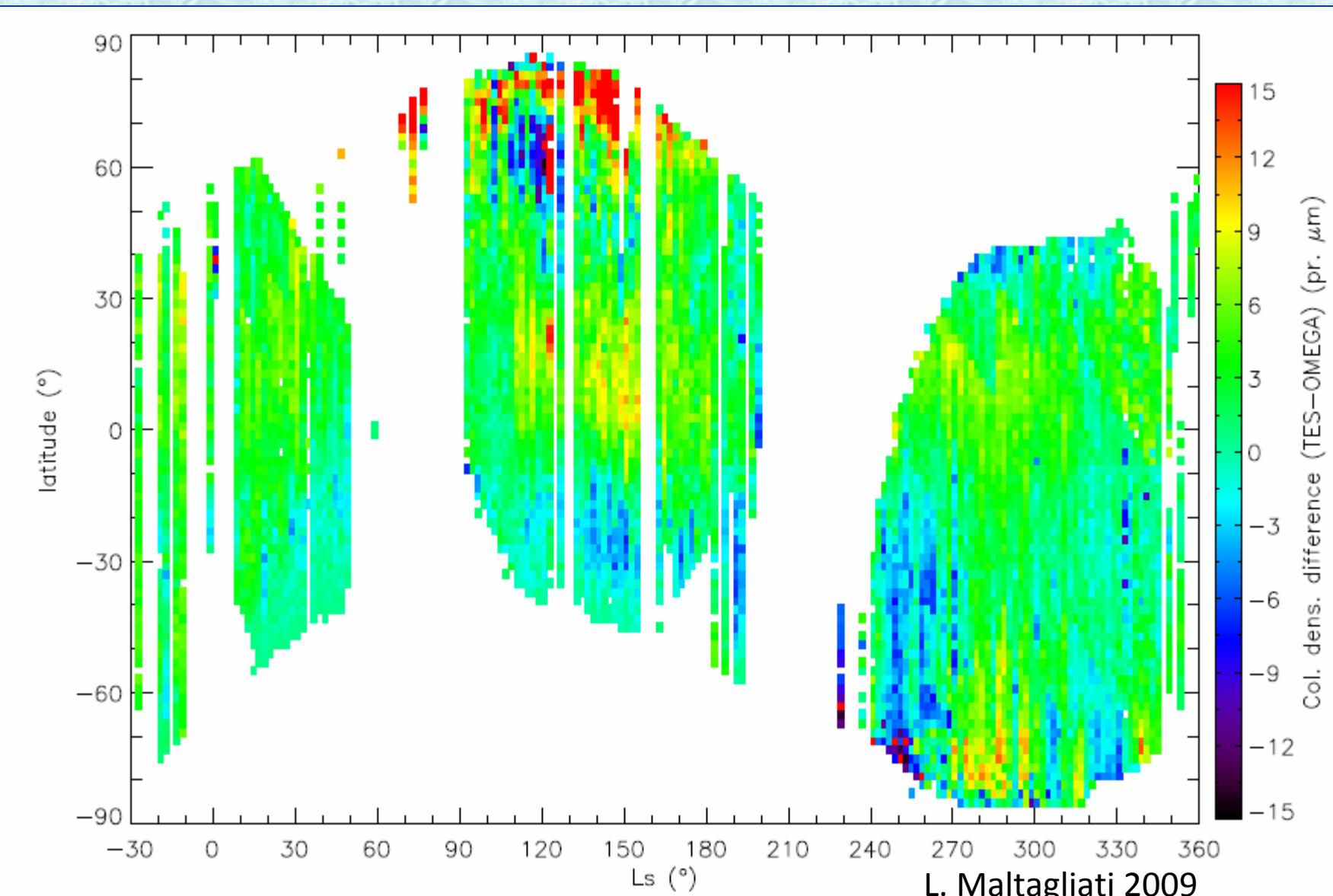


L. Maltagliati 2009

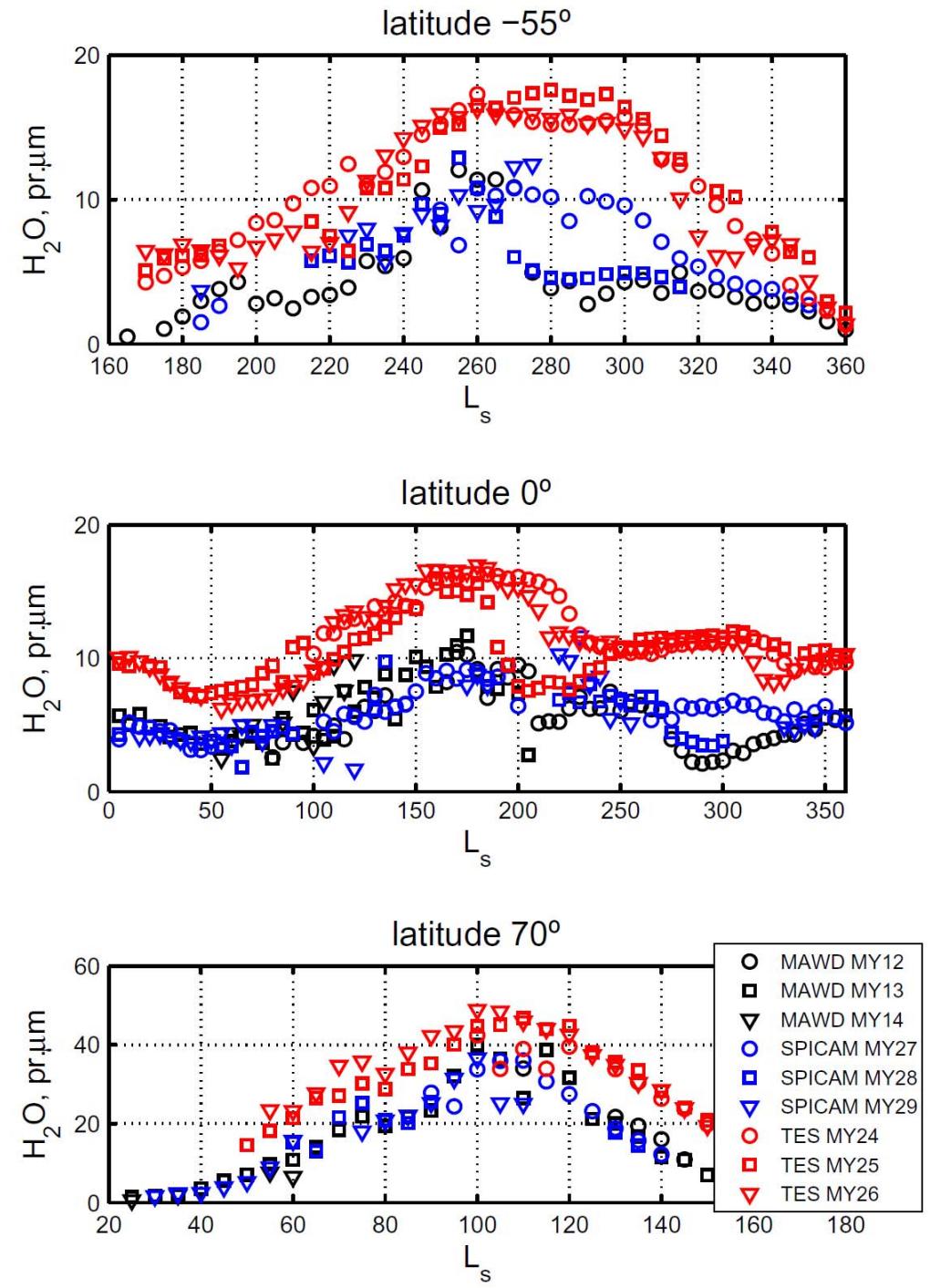
# SPICAM-OMEGA



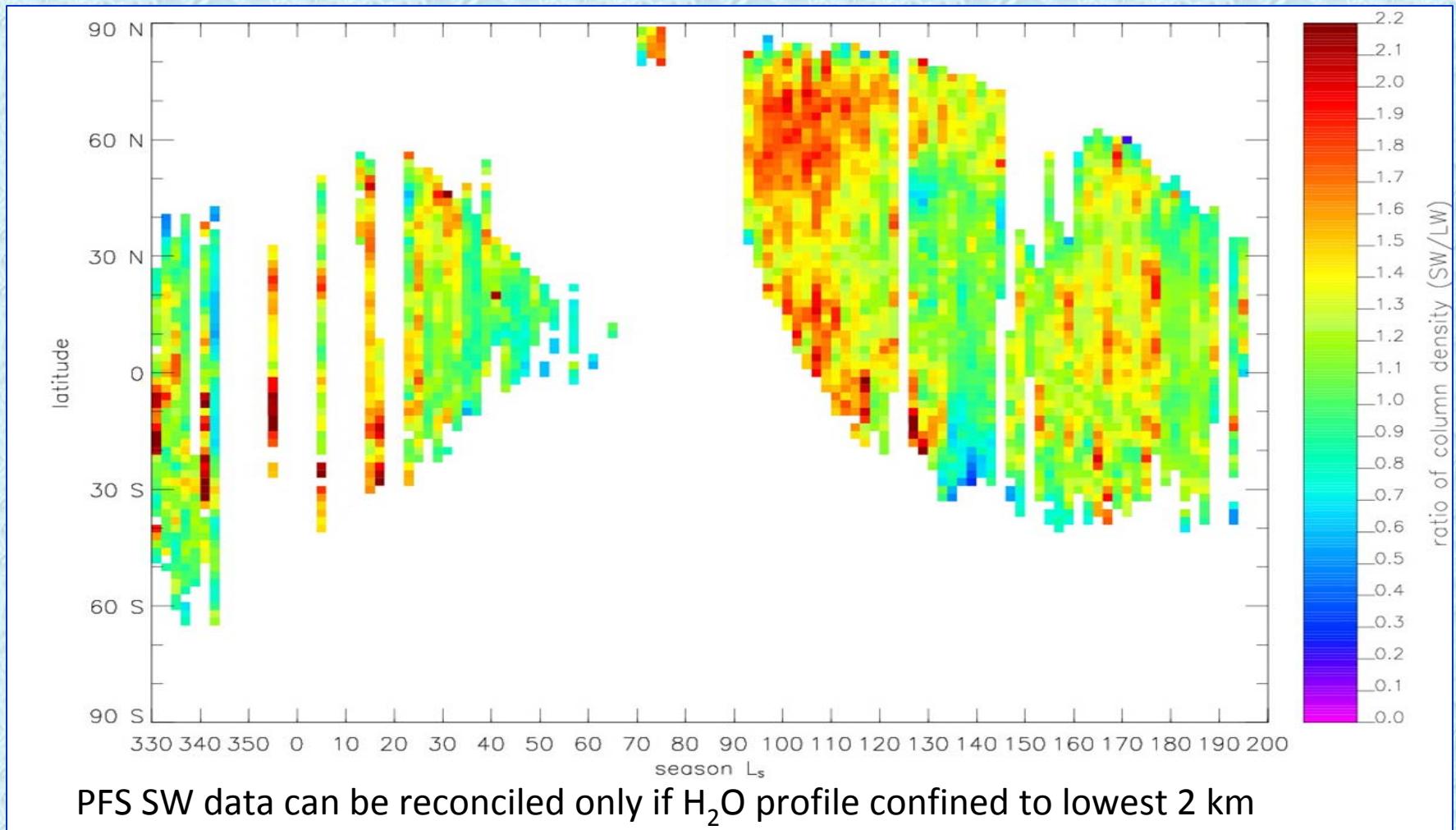
# TES-OMEGA



# MAWD-TES-SPICAM



# PFS SW/PFS LW





# Conclusions

- Reasonable agreement of Mars Express data sets, TES, and recently MAWD has been achieved

$$\text{MAWD}^* \approx \text{SPICAM} \leq \text{PFS LW} \leq \text{OMEGA} \leq \text{TES}^* \leq \text{PFS SW}$$

- \* Correction of TES (–30% for 60% of data) → new TES database (2006) available from M. Smith
- \* Correction of MAWD PDS dataset (–50-60% global; no aerosol accounted for; Fedorova et al, 2010)
- Generally dryer water cycle
- Much less seasonal variations

# Conclusions: MEX

- Thermal IR (PFS LW)
  - Most trustworthy (low contribution of aerosol)
  - Vulnerable to line broadening approximation in CO<sub>2</sub>
  - Short dataset of PFS
- OMEGA (2.56 μm)
  - Spectra of low resolution, vulnerable to contamination by H<sub>2</sub>O ice and other surface features
- PFS SW (2.56 μm)
  - Unexplained overestimation of H<sub>2</sub>O
- SPICAM (1.38 μm)
  - Most vulnerable to aerosol scattering

# Terrestrial remote sensing parallels

- Montoux et al., ACP 2009:
- The reference is balloon TDLAS (5-10% accuracy)
- HALOE displays the best precision (2.5%), followed by SAGE II (7%), MIPAS (10%), SAOZ (20–25%) and SCIAMACHY (35%), all of which show approximately constant H<sub>2</sub>O mixing ratios between 20–25 km. Compared to HALOE of ±10% accuracy between 0.1–100 hPa, SAGE II and SAOZ show insignificant biases

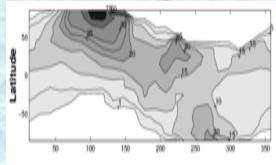
MIPAS is wetter by 10% and SCIAMACHY dryer by 20%.

- MIPAS (6.1, 10.5, 12.4 μm) may be considered as an analogue to TES and PFS LW
- SCIAMACHY (1.38 μm) is similar to SPICAM and MAWD

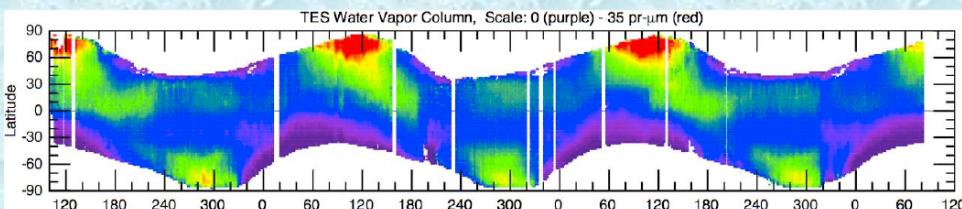
→ We must use “proxies” like on Earth

# What do we have for water vapor

MAWD MY 12-13



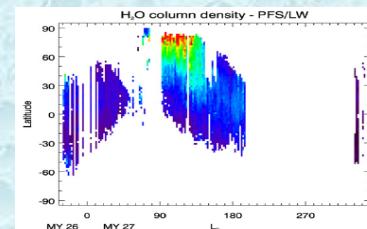
Jakosky and Farmer 1982,  
Reprocessed by  
Fedorova et al, 2004, 2010



M.D. Smith 2004, 2006, 2008

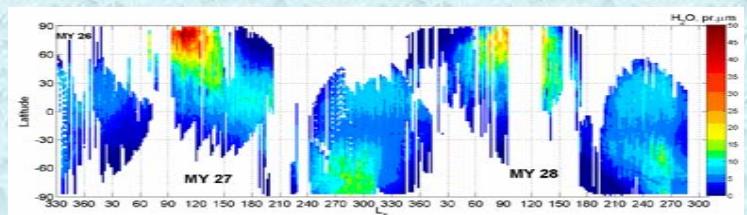
Fouchet et al., 2007  
Tschimmel et al., 2008  
Sindoni, Formisano 2010

TES MY 24-26



PFS MY 27

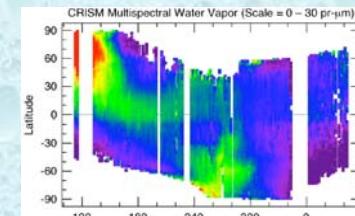
Fedorova et al., 2006  
Trokhimovskiy et al 2008  
Maltagliati et al 2008



SPICAM, OMEGA MY  
27-28

M.D. Smith et al., 2008, 2009

- Reasonable convergence of Mars Express data sets and TES is achieved
- Lessons for MAWD/ Viking:  
decrease of Viking results  
[Fedorova et al., Icarus 2010]



CRISM MY 28

# Vertical structure of atmospheric water

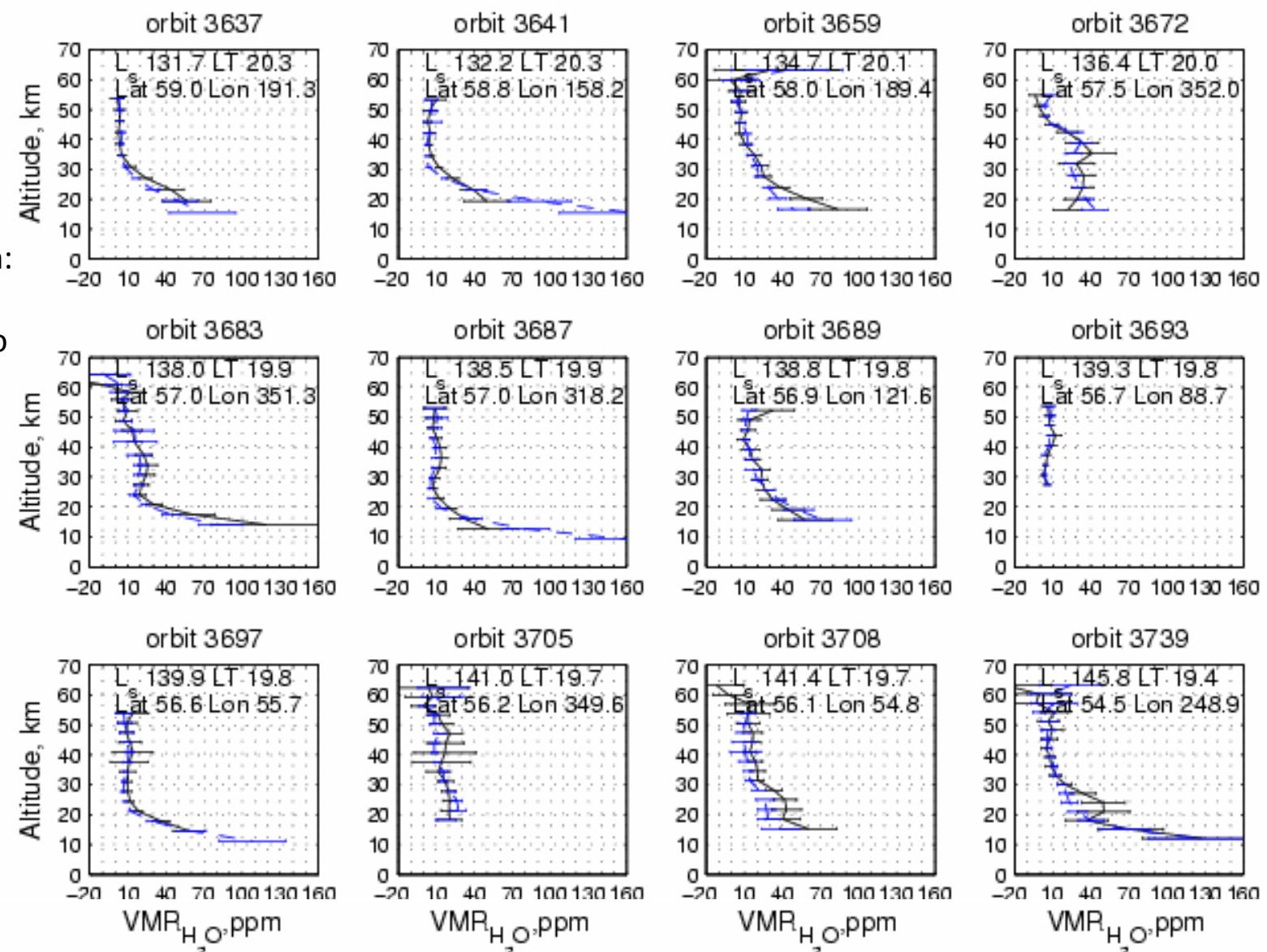
- Constraints from temperature and column measurements
- Constraints from microwave measurements (line profile allows to determine saturation altitude)
- Solar occultation measurements: both water and aerosols (difficult to disentangle ice and dust)
  - Auguste/Phobos 88 (Krasnopol'sky et al., 1991; Rodin et al., 1997)
  - SPICAM/MEX (Fedorova et al., Icarus 2009)
- Limb radiation profiles (MCS/MRO expected)

# H<sub>2</sub>O profiles by SPICAM/MEX

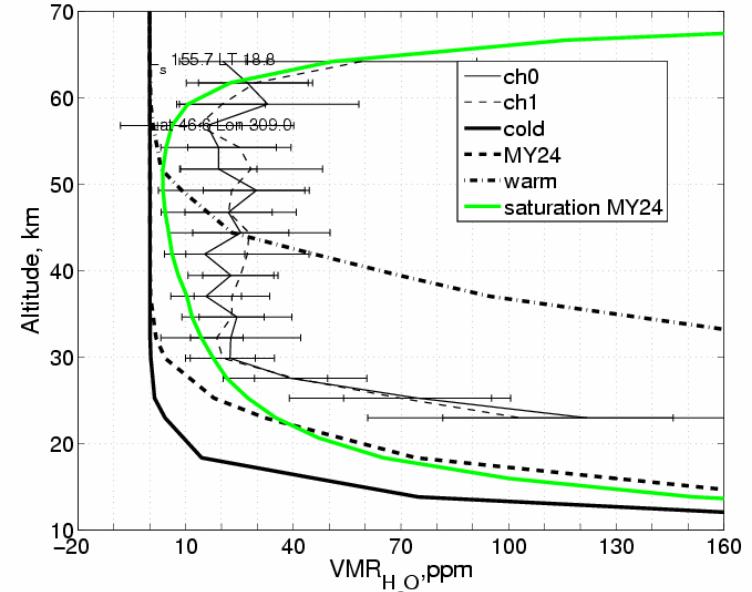
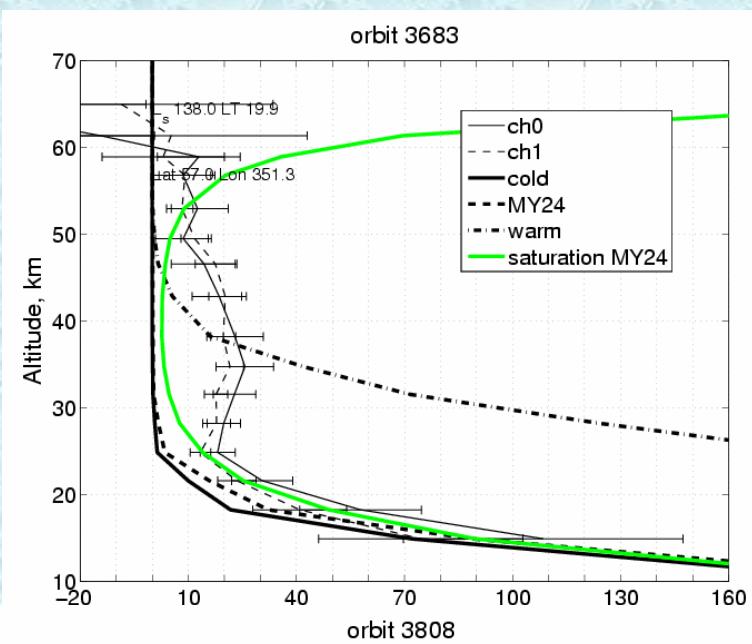
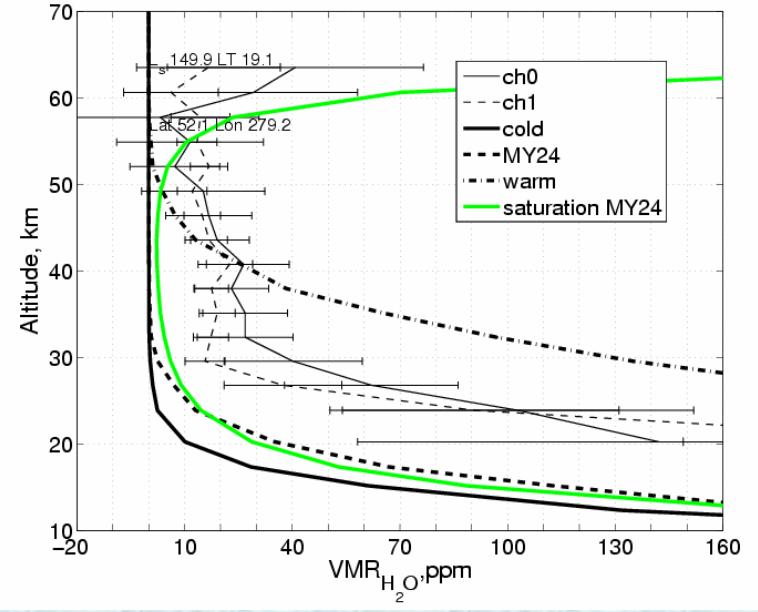
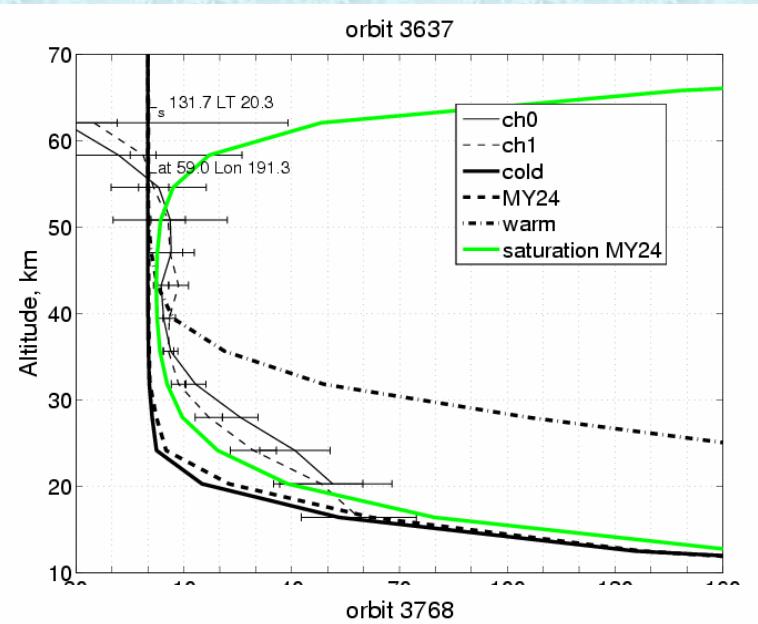
25 to 45 km:

5-15 pmm

Below 20-25 km:  
fast increase of  
H<sub>2</sub>O mixing ratio  
to  
50-100 ppm

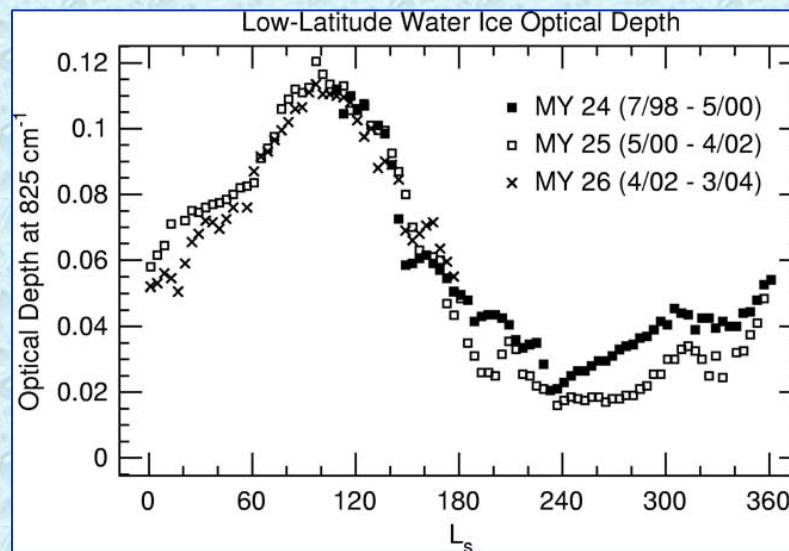
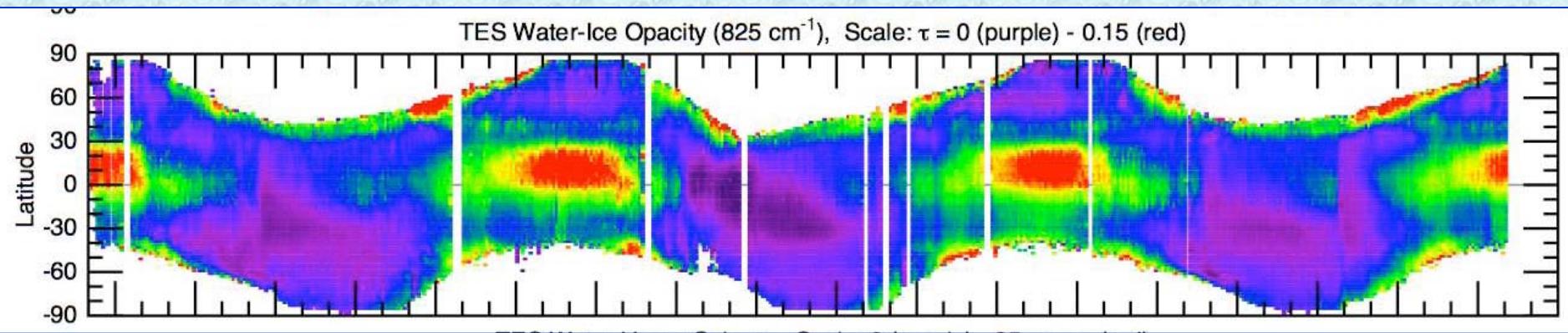


# Comparing SPICAM profiles with MGCM



# Water ice clouds

- Telescopic observations (e.g. HST)
- TES (and PFS) 12- $\mu\text{m}$  measurements
- IR mapping by OMEGA and CRISM (good for local mapping)
- UV measurements by SPICAM



# Future work

- Need to account for aerosol !
  - No regular aerosol database for Mars Express
  - Very difficult from NIR CO<sub>2</sub> bands
  - Scale coefficient =? when using SPICAM UV opt. depth
  - Could be only checked with EPFs
  - **Expected PFS aerosol database**
- Future measurements (ExoMars Orbiter 2016; surface network; future supporting orbiters):
  - Need continuity in monitoring of water vapour
  - Explore all types of orbits to characterize different cycles
  - IRIS-type instrument may be optimal (versatile “weather” instrument)
  - Possibly 6.3- $\mu$ m band useable (simplifies FTS)
  - Supporting NIR instrument (spectral resolution about MAWD)
  - Accurate measurements near the surface to characterize surface-atmosphere exchange: high sensitivity needed (night, southern winter): TDLAS
  - Profiling of the low scale height : bistatic lidar?